TECHNICAL REPORT ON THE RED LINE PROPERTY

SKEENA MINING DIVISION

NORTHWESTERN BRITISH COLUMBIA, CANADA

Latitude: 56° 58' 51.9" N

Longitude: 130° 32' 33.5" W

406,250 m E, 6,316,340 m N

(Universal Transverse Mercator Zone 9; 1983 North American Datum)

prepared for

COPPERHEAD RESOURCES INC.

607-1750 Davie Street

Vancouver, B.C.

V6G 1W3

by

P. Metcalfe, Ph.D. P.Geo. FGS

Palatine Geological Ltd. (1003060)

P.O. Box 289, Gabriola, B.C. V0R 1X0

4th November, 2022

1.0 SUMMARY

The 2416.9 hectare (ha) Red Line property ("the Property") comprises eight electronic mineral tenures, without financial, royalty or environmental encumbrance, located 990 km north-northwest of Vancouver, Canada and 18 km west of Bob Quinn airstrip, British Columbia. These tenures are owned entirely by McLymont Mines Inc., a wholly-owned subsidiary of Romios Gold Resources Inc. ("Romios"). The Property is the subject of an option agreement dated April 6th, 2022 between Romios Gold Resources Inc. and Copperhead Resources Inc. ("Copperhead") whereby Copperhead can acquire a 75% ownership interest in the Property from Romios by:

1. Spending a total of CAD$325,000 in exploration on the Property over a 36 month period;
2. Issuing an initial 1,000,000 common shares to Romios (completed);
3. Issuing 500,000 common shares to Romios on or before 6th April 2025;
4. Paying $75,000 in cash on or before 6th April 2025 and:
5. Entering into a joint venture agreement with Romios.

The Property lies near the eastern edge of the Boundary Ranges of the Coast Mountains, an area of severe topography with elevations varying from 300 m above sea level (a.s.l.) to more than 2000 m a.s.l., deeply incised by glaciers and streams. Valley bottoms and lower slopes are covered with fluvioglacial deposits. Vegetation on the valley bottom and sides is a mix of coniferous trees and understory species. Above an alp at an elevation of roughly 1,000 m a.s.l., vegetation thins to subalpine trees and shrubs, thinning upslope to heather and tundra. The climate is cool, moderated by Pacific air; consequently annual precipitation is high, estimated to be in excess of 1200 mm, roughly 2/3 of which falls as snow between the months of October and April. The operating season, owing to snowpack, begins in July and ends as late as October.

Resource roads run north and south of the Property, although the last few kilometres must be traversed using a helicopter. Bob Quinn airstrip is situated on British Columbia Highway 37 which extends from the interprovincial Yellowhead Highway (16) to the Alaska Highway in Canada's Yukon Territory. A spur of Highway 37 (37A) leads to Stewart, Canada's most northerly ice-free deepwater port, 205 km from Bob Quinn airstrip. All roads carry mining traffic. The immediate area supplies two operating mines with personnel, although supplies must be brought in from Stewart or from Terrace, further to the south. B.C. Hydro's Northwest Transmission Line passes to the east of the property and the Bob Quinn substation on the line is 15 km from the Property boundary; a spur of this line comes as close as 6 km.

The Property is located in the northwest of Stikinia, an ancient island arc accreted to continental North America in the Jurassic period. Stikinia hosts the area of the Golden Triangle which, as presented herein, extends northward from the historic Anyox mining camp to the operating Red Chris Mine at its northeast corner and to the developed prospect of the Galore Creek deposit in the northwest. This area encloses more than 1,000 mineral occurrences, including two producing and sixteen past-producing mines. Historic production of precious metals is approximately that of the Klondike.

The oldest rocks observed in Stikinia are of Devonian age, although it is likely that the microcontinent basement is older. Deep basement structures, almost as old as the microcontinent and with north-south and east-west orientation have controlled emplacement of magma and channelled hydrothermal fluids. In addition, Latest Triassic regional faults with northwest-southeast trend have similarly channelled mineralization. Structural breaks of both types transect the Property.

The layered supracrustal rocks and related intrusions in this part of Stikinia have been interpreted as products of no fewer than five episodes of arc magmatism:

1. Late Triassic (212-203 Ma) potassic undersaturated volcanic and intrusive rocks (Galore Creek and Copper Canyon porphyry deposits);
2. Latest Triassic to earliest Jurassic (*ca.* 205-196 Ma) alkaline porphyry-related, deformed mesothermal Ag-Au veins (Red Mountain);
3. Early Jurassic Texas Creek Plutonic Suite (*ca.* 196-187 Ma) alkaline plagioclase porphyry-related epithermal, transitional and mesothermal Ag-Au veins and base and precious metal deposits (Silbak Premier, Scottie Gold, Kerr, Sulphurets, Mitchell, Iron Cap, Valley of Kings, SNIP, Stonehouse, Bronson Slope);
4. Latest Early Jurassic (*ca.* 185-183 Ma) small, poorly mineralized porphyry intrusions; and:
5. Middle Jurassic (*ca.* 175-172 Ma) calc-alkaline arc and tholeiitic back-arc magmatism and syn- and epigenetic, stratabound base and precious metal deposits (*e.g.*: Eskay Creek deposit) related to the back-arc basin formation.

At least one period of deformation, in the latest Triassic/earliest Jurassic, formed northeast-vergent folds whose axial planes became loci for compressional shearing, followed by dextral transpressional displacement, dextral transtension and, lastly, extension. This deformation is at least in part synchronous with Galore Creek and Texas Creek volcanism, magmatism and precious metal mineralization.

Stratified rocks in the Upper Triassic Stuhini Group and the uppermost Triassic to Middle Jurassic Hazelton Group formed from deposits typical of volcanic terrains, commonly deposited on deeply eroded palæosurfaces. The necessity of correlating changes in palæosurface, depth of volcanic pile and depth of mineralization for any given exploration target require that regional geological mapping carried out by government be augmented by detailed property-scale mapping to provide an adequate framework for mineral exploration.

The Property lies across the valley of Downpour Creek, to the east of the Forrest Kerr fault trace which separates Palæozoic sedimentary and plutonic rocks to the west from a section of prospective Triassic, Lower Jurassic and Middle Jurassic volcanosedimentary rocks to the east. This area, bounded by Forrest Kerr Creek and the Iskut River, is entirely within the interpreted northerly projection of the Eskay Rift and exposes rocks from all the prospective metallogenic episodes of Mesozoic magmatism and mineralization.

Mineral deposit types considered for the Property therefore include, without limitation and in no preferred order: porphyry Cu-Au; intrusion-related precious metal (gold) veins; epithermal deposits of any sulphidation type and sediment-water (seafloor) exhalative deposits of any type. The reader is also encouraged to consider the possibility of the existence of a skarn deposit, but is enjoined to consider the principles of metal source and transportation pathway, ground preparation, depositional trap mechanism and subsequent preservation of mineralization, rather than being constrained by conventional models of mineral deposit formation.

Fieldwork in 2022, constrained by the property boundaries, confirmed the presence of a folded, fine-grained siliciclastic sedimentary sequence in the northern and north-central part of the

Property, overlying a coarse fragmental unit with exclusively volcanic clasts, albeit with little to no evidence of hot emplacement. The observed section is consistent with a Middle Jurassic Downpour Creek siliciclastic unit unconformably overlying volcaniclastic rocks of the Lower Hazelton Group. The south-central part of the Property is underlain by fine-grained clastic sedimentary rocks correlated with the Upper Triassic Stuhini Group. Fieldwork located a single subcrop or proximal float occurrence of K-feldspar megacrystic porphyry on its southwestern ridge. Its similarity to metallogenic intrusive rocks at Galore Creek and other locations in the vicinity was noted at the time. These are late Triassic intrusions and mineralizing systems which predate Rhaetian-Hettangian erosion, therefore any mineralized system of this age would lie entirely beneath the Jurassic strata on the Property and would, therefore, be a blind target.

Previous assessment work noted two southeast-trending geophysical anomalies crossing the southwest ridge where a moderate arsenic (As) anomaly was returned both from 1990 geochemical sampling and from current fieldwork. Geochemical and geophysical anomalies are spatially coincident with a southeast-striking, southwest dipping fault or fracture zone mapped during inspection of the Property.

Recommendations for future fieldwork proceed from field observations and historic data:

1. Consultation with independent experts in geophysics and in exogenic geochemistry, for the purposes of recommending optimal techniques for the location of blind precious metal deposits of exhalative, porphyry and intrusion-related precious metal vein deposits under glacial and bedrock cover;
2. Planning and execution of the recommended deep-penetration geophysical survey;
3. Isotopic age measurement and geochemical analysis of a sample of K-feldspar porphyry where located in outcrop, using U-Pb laser ablation mass spectrometry on zircon to allow for determination of the primary oxidation state of the magma (and thereby potential fertility);
4. Creation of a test line for evaluation of exogenic geochemical methods for locating blind porphyry, vein and Eskay Creek-style exhalative targets;
5. Detailed mapping of the Property, with particular attention to intrusive rocks and any associated alteration, accompanied by intensive prospecting and:
6. Planning and execution of the recommended exogenic geochemical survey.

This exploration is recommended in two phases to allow for the timing of the upcoming operating season. Phase 1 comprises items 1-4 at a cost of CD$100,000. Phase 2, contingent on encouragement from Phase 1, comprises items 5-6, at a cost of CD$255,600.

TABLE OF CONTENTS

LIST OF FIGURES

LIST OF TABLES

2.0 INTRODUCTION

The author was engaged by Copperhead Resources Inc. ("Copperhead"), a private British Columbia (B.C.) corporation, to prepare a technical report compliant with the guidelines of National Instrument 43-101 (NI43-101) on the Red Line property "the Property". The Property is the subject of an April 6th 2022 option agreement between Romios Gold Resources Inc. ("Romios") and Copperhead whereby Copperhead can acquire a 75% ownership interest in the Property by:

1. Spending a total of CAD$325,000 in exploration on the Property over a 36 month period;
2. Issuing an initial 1,000,000 common shares to Romios (completed);
3. Issuing 500,000 common shares to Romios on or before 6th April 2025;
4. Paying $75,000 in cash on or before 6th April 2025 and:
5. Entering into a joint venture agreement with Romios.

The technical report was prepared to assess the Property as a Qualifying Property for a listing of Copperhead's shares on a recognized Canadian stock exchange. The author conducted a Property inspection between the 9th and 19th July, 2022, spending a total of nine days in the field. Sources for the report include the Geological Survey of Canada, British Columbia Geological Survey, U.S. Geological Survey, British Columbia's MINFILE, Mineral Titles Online (M.T.O.) and Assessment Report Indexing System (ARIS) and publications in various scientific journals.

3.0 RELIANCE ON OTHER EXPERTS

Any technical report based on a brief site visit must, of necessity, rely on information acquired by previous workers. In the case of the Red Line property, this information is derived in part from research carried out by both the Geological Survey of Canada and the British Columbia Geological Survey Branch and in part from exploration carried out by owners of previous mineral tenure on ground composing the present Property. Much of this research and exploration was carried out prior to the institution of National instrument 43-101 and, therefore, prior to any definition of "qualified person".

This author has worked for both federal and British Columbia surveys and for the University of British Columbia during the course of his career, is professionally acquainted with nearly all researchers previously active in the area and is entirely satisfied with the quality of their research. These sources and others, where used, will be referenced at the appropriate place in the text. Notwithstanding these sources, this author takes full responsibility for his interpretation of data from this and previous studies and for the content of this report.

Units used in this report are metric. Positional data, including maps, are presented in a Universal Transverse Mercator (UTM) projection (Zone 9), using the 1983 North American Datum (NAD'83), unless otherwise stated (*e.g.*: the insert on Figure 1).

4.0 PROPERTY DESCRIPTION AND LOCATION

The 2416.9 hectare (ha) Red Line property comprises eight electronic mineral tenures located on the eastern edge of the Coast Ranges of British Columbia 990 kilometres (km) north-northwest (NNW) of Vancouver, Canada. The Property, elongated north to south, is centred on latitude 56°

58' 51.9" N and longitude: 130° 32' 33.5" W (406,250 m E, 6,316,340 m N), 18 km due west of the Bob Quinn airstrip; this airstrip is adjacent to B.C. Provincial Highway 37 (Figure 1).

Mineral rights are the purview of his Majesty the King in right of the Province of British Columbia. Presently, mineral rights are held through three distinct types of tenure. The first and oldest of these is through a Crown granted mineral claim with a maximum dimension of 457 m (1,500 feet) on a side; the practice of granting such claims was discontinued in 1960. The second type of mineral rights is found in legacy mineral tenures, wherein mineral rights to any available, unprotected area were secured by physical staking of a two-post or four-post mineral tenure, each claim unit with a maximum dimension of 500 m on a side and enclosing 25 ha. In excess of 5,000 legacy mineral tenures and 7,000 Crown granted mineral claims survive to the present day and have a potential impact on mineral rights acquired in any subsequent cospatial electronic mineral tenure. However, neither type survives in the area of the Property.

Acquisition of mineral tenure through physical staking was discontinued in 2005 and the present system of electronic tenure acquisition began. Tenures are acquired by selection of cells on a series of predefined grids covering the entire province. The cost of acquisition is by area.

Mineral tenures composing the Property are all of the last, electronic type, are listed in Table 1 and are shown in Figure 2. The owner of record for each tenure is McLymont Mines Inc., a wholly-owned subsidiary of Romios Gold Resources Inc. To the best of the author's knowledge, none of the tenures carries any environmental liability, royalty, back-in right, required payment or other encumbrance.

Table 1. Mineral tenures composing the Property.
Data from Mineral Titles Online, 28th August 2022.

Tenure no.	Claim Name	Tenure type	Issue date	Expiry date	Area (ha)	Free Miner no.	Owner (100%)
1078227	NEWMONTEAST	Mineral Cell Title Submission	27Aug2020	27Aug2023	440.7898	146096	McLymont Mines Inc.
1078228	NEWMONTEAST1	Mineral Cell Title Submission	27Aug2020	27Aug2023	35.2584	146096	McLymont Mines Inc.
1078229	NEWMONTEAST2	Mineral Cell Title Submission	27Aug2020	27Aug2023	405.8476	146096	McLymont Mines Inc.
1078230	NEWMONTEAST3	Mineral Cell Title Submission	27Aug2020	27Aug2023	17.641	146096	McLymont Mines Inc.
1078231	NEWMONTEAST4	Mineral Cell Title Submission	27Aug2020	27Aug2023	352.2826	146096	McLymont Mines Inc.
1089834	RED LINE 1	Mineral Cell Title Submission	21Jan2022	21Jan2023	600.0016	146096	McLymont Mines Inc.
1096973	RED LINE 2	Mineral Cell Title Submission	06Aug2022	06Aug2023	353.1513	146096	McLymont Mines Inc.
1096974	RED LINE 3	Mineral Cell Title Submission	06Aug2022	06Aug2023	211.8936	146096	McLymont Mines Inc.

Five of the mineral tenures are in their third year, while the remaining three were acquired in this calendar year. To maintain mineral tenure in British Columbia, assessment work must be carried out during the first year after staking and, thereafter, whenever appropriate. The value of the assessment work carried out is applied to extend the expiry date of any mineral tenure in the contiguous group on which the assessment work took place. To maintain a tenure or group of tenures for the first two years, the value of the work carried out on the group must equal or exceed five dollars per hectare per year. Thereafter, for the third and fourth years, the required expenditure is $10 per hectare per year; for the fifth and sixth years, $15 per hectare per year and, subsequently, $20 per hectare per year.

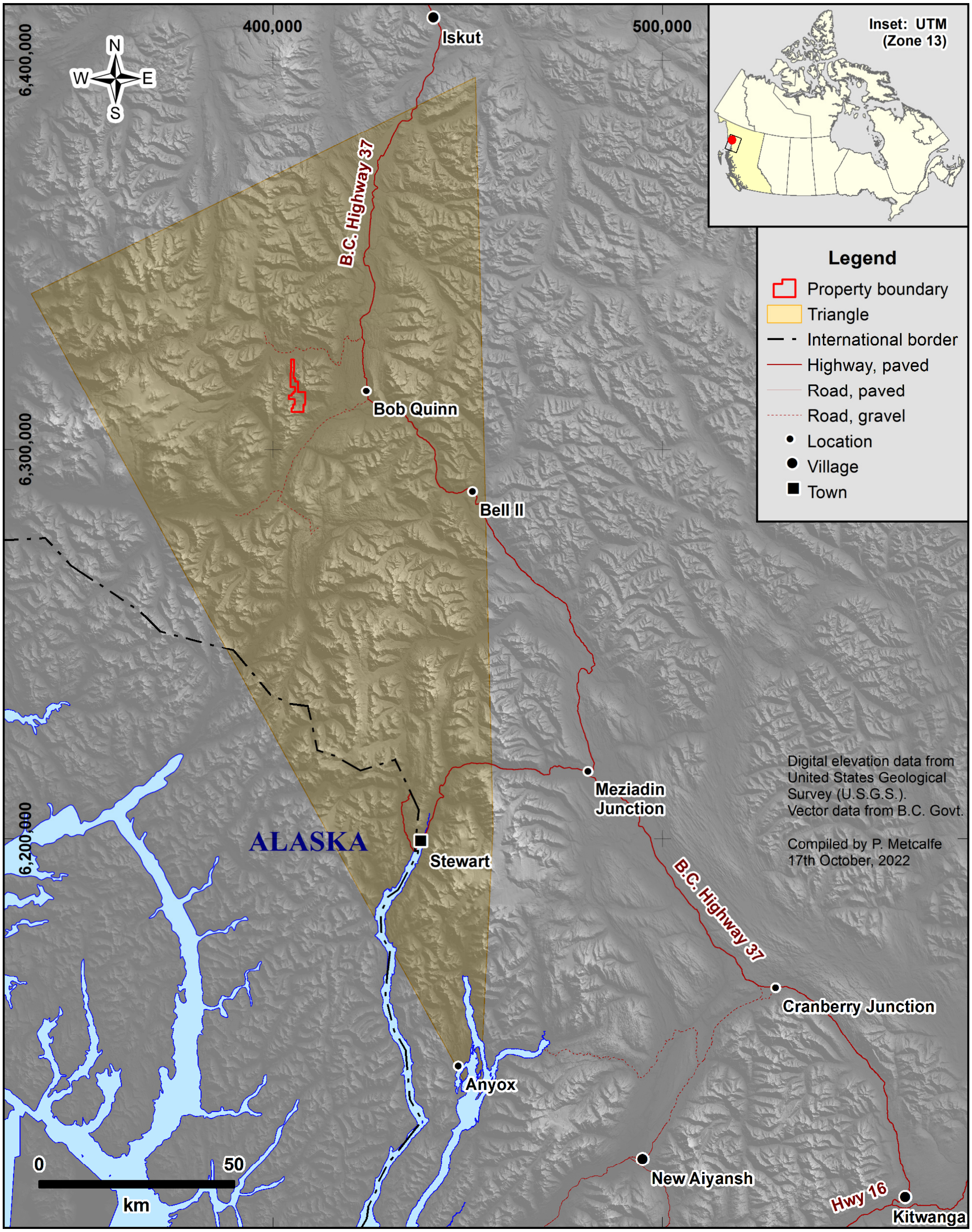

Figure 1. Property location and access.

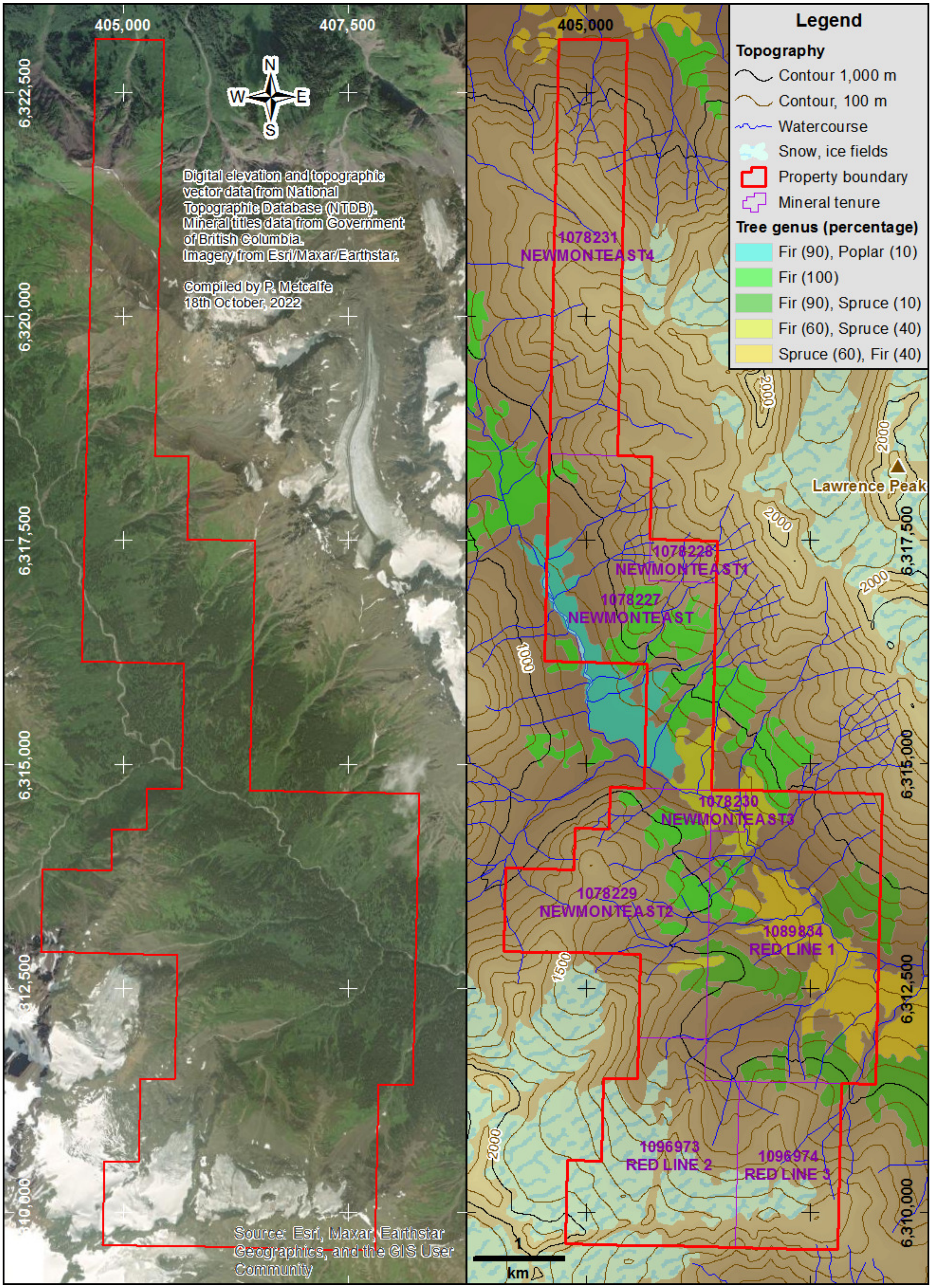

Figure 2. Property mineral tenure, physiography and vegetation.
Data from National Topographic Database, B.C. Mineral Titles and B.C. Forest Service.

Surface rights in British Columbia are independent of mineral tenure. A permit from the government of British Columbia is required for camp construction, substantial removal of timber or other disturbance. Non-intrusive exploration on the Property (*e.g.*: geological mapping, geochemical sampling) does not require permitting. Intrusive work, such as trenching, diamond drilling and induced polarization (IP) geophysical surveys, do require a permit. Under such circumstances, the application for such a permit is referred by the British Columbia government to interested parties, including local communities, for comment prior to issuing the permit.

5.0 ACCESSIBILITY, CLIMATE, LOCAL RESOURCES, INFRASTRUCTURE AND PHYSIOGRAPHY

5.1 Topography, elevation, and vegetation

The Property lies in the eastern Boundary Ranges of the Coast Mountains of British Columbia, an area characterized by mountainous terrain deeply incised by glaciers and rivers. Mountain peaks in the immediate vicinity are in excess of 2000 m above sea level (a.s.l.). The glaciated valley of More Creek to the north of the Property is at an elevation of 400 m a.s.l. and the valley of the Iskut River to the south and east of the Property is at an elevation of 330 m a.s.l. Elevations on the Property vary between 610 m a.s.l. on the Property's eastern edge and 2140 m a.s.l. at its southwestern corner. The valleys of Southmore and Forrest Kerr Creeks form the western boundary of a divided *massif* which includes Lawrence Peak and Faisal Peak.

The Property itself straddles the valley of Downpour Creek (see Section 5.4: Climate), which divides the area defined above. The creek rises in the icefields to the west of the Property, flowing northeast to a point just west of the Property's northwestern edge, then turns sharply to flow in a southeasterly direction across two segments of the Property. As the creek begins its second transect of the Property, it enters its canyon, formed by fluvial overdeepening of the glacial valley. This canyon greatly limits the availability of locations in the valley bottom of the southern Property (Red Line 1) where a helicopter may be safely landed. To north and south, the steep, vegetated valley sides preclude landing below the lowest alp, at roughly 1000 m a.s.l., although landing is possible on gravel bars in the creek above its canyon.

Figure 2 shows the disposition of tree species on the Property. The subalpine floor of the northern valley is host to a mixture dominated by fir, with minor poplar. On the sides of the upper valley, fir becomes exclusive, admixed on the downstream slopes with spruce and progressing to a spruce-dominant mix on the lower valley bottoms to the southeast.

A particular feature of the valley bottom is the presence of superannuated, dying trees, which occur far less commonly in the managed forestry regions in more accessible parts of British Columbia. Defoliated and fallen trees have created gaps in the forest canopy akin to gaps created by slope failure on the valley sides. Here and on the avalanche and landslide slopes on the valley sides, verdant growth of alder, willow, salmonberry and, most noticeably and inevitably, devil's club significantly hinders progress on foot, particularly across-slope.

Tree line, above the lowermost alp, is a transition zone between stunted subalpine fir and a zone of perennial and annual alpine flowering plants. Upslope, the vegetation follows the progression common to the alpine of northwestern British Columbia, passing through a zone of white mountain heather and sparse grass to tundra.

5.2 Means of access to the Property

The Property lies 18 km to the west of Bob Quinn airstrip, located on B.C. Highway 37. This paved highway extends north from its junction with the interprovincial Yellowhead Highway (Hwy. 16) immediately south of Kitwanga (Figure 1) northward to the Alaska Highway near Upper Liard in the Yukon Territory. Stewart, Canada's northernmost ice-free deepwater port, is accessible *via* a branch (37A) of Highway 37, 205 km from Bob Quinn. Both highways have been upgraded in recent years and presently handle heavy traffic from the Red Chris Mine.

The remaining 18 km are more problematic, as noted in the previous section. Two gravel roads, both constructed to carry heavy traffic, depart the highway 3 km south and 8 km north of the airstrip, respectively. The road departing south of the airport is the old road to the Eskay Creek mine, with private access only. It runs along the left (south) bank of the Iskut River, passing less than 5½ km from the southern boundary of the Property. Despite proximity, there is no crossing of the Iskut River at that point; the nearest such lies more than 30 km southwest of the Property.

The more northerly route is the proximal part of the mining road to the Galore Creek project, operated by the Galore Creek Mining Corporation. This road crosses the Iskut River, but runs on the north side of More Creek, with no crossing point to facilitate the last 2 km to the Property. Absent any change, access (by helicopter) is, logistically, easier from Bob Quinn airstrip.

5.3 Proximity of the Property to a population centre and nature of transport

Bob Quinn airstrip, the recommended departure point for helicopter access to the Property, is located between kilometres 294 and 296 on Highway 37. The communities of Iskut, Telegraph Creek and Dease Lake lie to the north within a radius of roughly 200 km and a significant part of the populations are workers in mineral exploration and mining. Notwithstanding these communities, nearly all supplies must be sourced from the south, *via* the Port of Stewart (205 km), the towns of Terrace (394 km) and Smithers (411 km) and, ultimately, Vancouver.

The Property is well-situated for access to hydroelectric power. Its boundary is less than 15 km west of the Bob Quinn substation on BC Hydro's Northwest Transmission Line and only six kilometres from the spur line from the Forrest Kerr 195 MW run-of-river hydroelectric generating station.

5.4 Climate and length of the operating season

The author is not a climate scientist and his experience is confined to observations of the weather in the Iskut River region during a number of time periods between 1988 and 2022. There is a paucity of climate data for the region, mainly with regard to continuity, owing to the ephemeral nature of the mines and airstrips operational in the area during this time period. The greatest longevity of any station in the area is that at Bob Quinn airstrip which produced data for the Federal Ministry of the Environment from 1971 to 2000 (Government of Canada 2022).

The reader should note that Bob Quinn lies in the rain shadow of the Coast Ranges at an elevation of 610 m a.s.l., somewhat below the minimum elevation on the Property. Based on his experience, the author considers the temperature averages of -8.5° in January to 14.1° in July to be crudely representative of the lower elevations of the Property and the estimated annual rainfall of 641.6 mm an overly optimistic number. The reader is referred to a publication from the Brucejack Mine project, roughly 60 km south-southeast of the Property (Rescan 2013), wherein a model originated by Wang *et al.* (2006) was used to estimate an elevation-dependent annual precipitation of between 1200 and 1600 mm at Brucejack Lake.

Notwithstanding average monthly temperatures, snow is possible at any time of year at nearly any elevation and roughly two-thirds of the annual precipitation falls as snow. In years of heavy winter precipitation, snow-pack from the previous year might hinder exploration at higher elevations until as late in the year as September. A case in point was the 2022 field season, when the snowpack in July was the deepest and latest-melting the author had seen. Despite this, all major and many subsidiary drainages flow throughout the year, except at alpine elevations. Fieldwork at higher elevations is usually possible until October.

6.0 HISTORY

6.1 Regional history

The history of exploration in the immediate area of the Red Line property is best viewed within the context of the regional exploration of the area which came to be known as the Golden Triangle. Exploration of the coastal areas began before Confederation. Souther (1971) described the first discovery of fine placer gold on the Stikine River in 1873 and subsequent prospecting along that river throughout the Cassiar and Klondike gold rushes. Souther also described a shift in emphasis around the turn of that century, from placer to bedrock exploration.

In 1909, evidently following colours up the Iskut River, a tributary of the Stikine, prospectors staked two groups of mineral claims, now Crown-granted, on the lower slopes of Johnny Mountain, near the confluence of Bronson Creek with the Iskut. However, as Souther (*ibid.*) noted: "It was not until 1955 that systematic exploration of the more remote parts of the map-area was begun"; two World Wars had intervened. Souther continued:

> *"In [1955] saturation prospecting employing large helicopter-supported parties was initiated by Hudson Bay Mining and Smelting Company. Since that time many parts of the map-area have undergone intensive exploration by geophysical and geochemical as well as conventional prospecting methods. A large number of prospects have been explored by trenching and diamond drilling and at least two of these (Galore Creek and Liard Copper) have proven reserves in excess of 200 million tons of +0.50 per cent copper equivalent[1]."*

The immediate consequence of corporate interest in the area was Operation Stikine, one of the first helicopter-based mapping projects to be carried out by the Geological Survey of Canada (1957). The map produced covers six 1:250,000 scale map sheets and includes the Property area near its eastern margin.

The 1971 US default on gold may have been one of the contributing factors in the decline of major corporate investment in mineral exploration, culminating in a recession in the early 1980s. When interest resumed in the Iskut River area, it was at the hands of small, investor-funded companies active on the Vancouver Stock Exchange. This renewal of junior-funded activity led

---

[1] The author regrets the necessity of stating the obvious: that this estimate is "non-compliant and should not be relied upon". A reliable number (Henderson 2018) is: 528 Mt Proven & Probable Reserves at 0.585% Cu, 0.32 g/t Au, 6.02 g/t Ag, considerably in excess of Souther's estimate.

to the discovery and subsequent exploitation of the Stonehouse and SNIP deposits on Johnny Mountain, 45 km to the southwest of the Property. The term "Golden Triangle" began to be used (candidly, for marketing purposes) to describe the crudely triangular area extending from Anyox in the south to the past-producing Sulphurets placers in the northeast and to the Johnny Mountain area in the northwest. This area, bounded on its north side by the Iskut River, crudely resembled that defined geologically by Grove (1971, 1986) as the Stewart Complex.

The increase in exploration activity during the late 1980s to early 1990s engendered increased research activity on the part of the Geological Survey of Canada (Anderson 1989, 1993), the Geological Survey of British Columbia (Alldrick and Britton 1988, 1992, Alldrick *et al.* 1989, Alldrick *et al.* 1990, Fletcher and Hiebert 1990, Lefebure and Gunning 1989, Alldrick 1989) and the Mineral Deposit Research Unit (*e.g.* Macdonald *et al.* 1993, Macdonald *et al.* 1996).

Similar studies were carried out to the east, to include the area of the present Property north of the Iskut River (Logan *et al.* 1990, Logan *et al.* 1997). The cause of this was the 1989 discovery of the world-class Eskay Creek deposit, which lay outside the bounds of the "original" Golden Triangle. More recently, mapping by Alldrick *et al*. (2005a, 2005b) succeeded in tracing the host stratigraphy of the Eskay Creek deposit north of the Iskut River (Nelson *et al*. 2018).

The development of the Red Chris deposit into an operating mine led to a variety of creative interpretations of the original topology of the "Golden Triangle". The author's own reconstruction is presented as a compromise in Figure 1, incorporating the original Anyox and Stewart mining camps in its southern part, the world-class Brucejack Lake, Eskay Creek and SNIP deposits in its central part and, north of the Iskut River, the northern Triangle, including Red Chris, Galore Creek and Schaft Creek.

6.2 Property history

The history of ownership of ground covered by the present Property is considerably shorter than that of the region as a whole. Post-1970 exploration interest in this ground, on the eastern edge of the Mesozoic outcrop area, began in 1988, owing to the presence of the mineralized Forrest Kerr Fault system extending north from the Iskut River to More Creek. Previous mineral tenures, initially held by Ecstall Mining Corporation and Omega Gold Corporation, are extinct.

This technical report includes previous work on these now-extinct mineral tenures, referred to in assessment reports as the FOR and NING claim groups. Five previous exploration projects were carried out over this ground (Figure 3). Four comprised geochemical rock and stream sediment sampling (Nicholson 1990a, 1990b, Gal and Walker 1990, Walker and Moore 1990). As a consequence, Kennecott Canada Inc. acquired an option to earn an interest in the For property in 1990 and, in 1992, carried out an airborne geophysical survey comprising four-frequency electromagnetic, high sensitivity magnetometric and two-frequency very low frequency electromagnetic (VLF-EM) measurements. The target was an Eskay Creek type stockwork-massive sulphide deposit (Fields 1992). Despite moderate encouragement, no further work was carried out and the tenures lapsed in 1994. Holders of subsequent mineral tenures on land now enclosed by the Property carried out no work on the land; these tenures are also extinct.

The Property itself has a limited history of ownership. The component tenures were acquired by Romios through McLymont Mines Inc. between 2020 and 2022 (Table 1). A single short exploration program, carried out in 2021, was severely curtailed by weather (Biczok 2021).

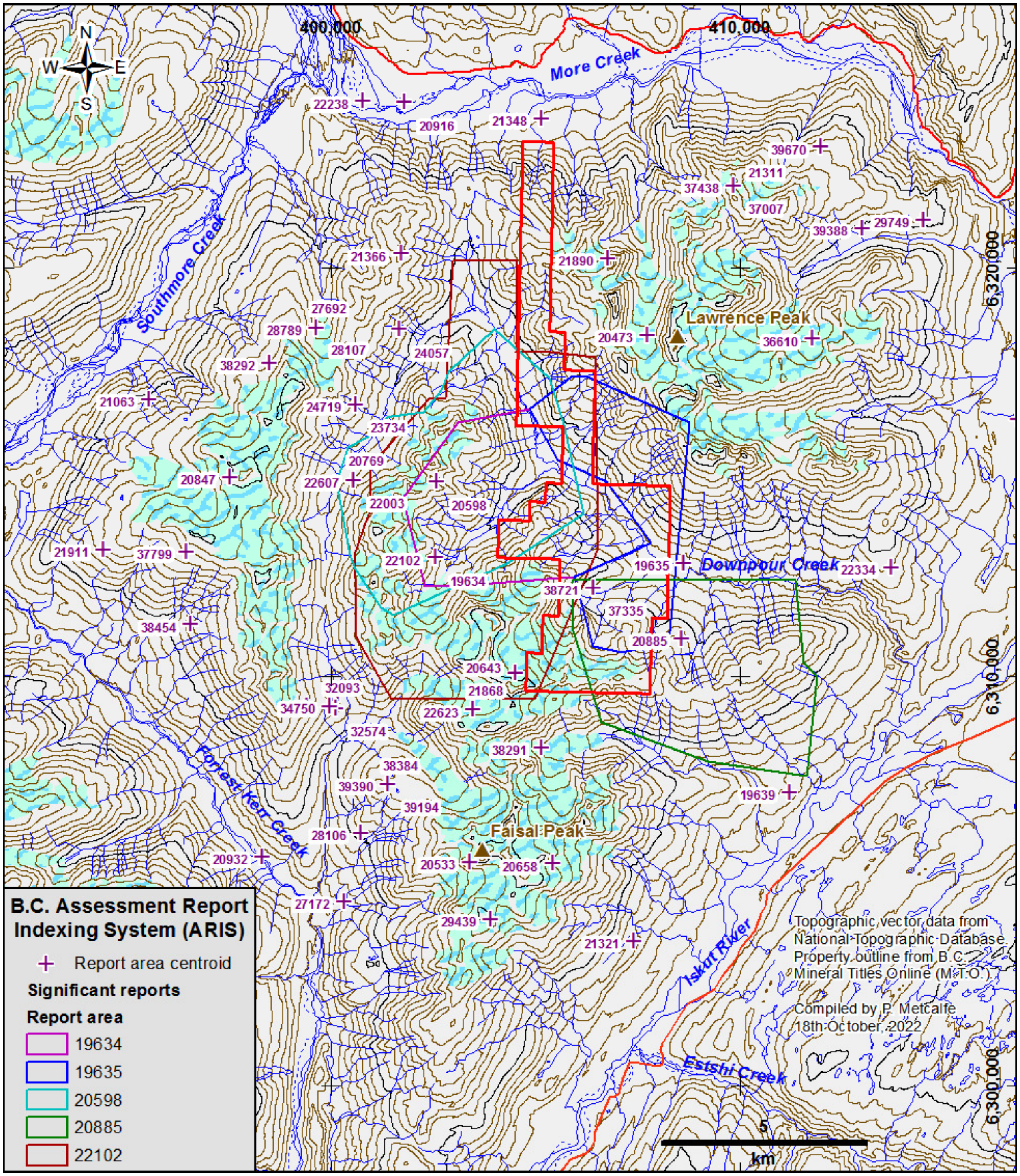

Figure 3. Local topography and previous exploration in the Property area.

The Property occupies a polygonal highland area bounded by the northeasterly-flowing Southmore Creek, the easterly-flowing More Creek, the southerly, then southwesterly-flowing Iskut River and the southeasterly-flowing Forrest Kerr Creek. Coloured polygons represent the exploration footprints of B.C. Assessment Reports on ground now enclosed by the Property. Polygons representing work on adjacent ground are not shown.

7.0 GEOLOGICAL SETTING AND MINERALIZATION

7.1 Regional geology

The Property is located within the Intermontane Belt of the Canadian Cordillera near the western margin of the Stikine terrane. This terrane represents the remnants of the microcontinent of Stikinia, which collided with and was accreted to North America in the Jurassic period. More specifically, the Property lies within an area of Stikinia extending north and northwest from a southern apex at the old mining camp of Anyox (Figure 1) and which hosts more than 1,000 mineral occurrences of dominantly precious metal vein type, with related skarn, porphyry and massive sulphide occurrences. The area encompasses metamorphic and plutonic rocks of the Coast Plutonic Complex on the west, is dominated by Stikinia and includes part of the western margin of the Bowser Basin (Evenchick 1991a, 1991b) to the east (Figure 4).

Grove (1971, 1986) defined the Stewart Complex as the area of Mesozoic rocks exposed south of the east-west structure defined by the valley of the Iskut River, in a corridor narrowing southward towards the tidewater mining camps of Stewart and Anyox. This geological definition was cospatial, for a time, with the name "Golden Triangle" beloved of promoters, until the size and economic viability of the Red Chris deposit were realized. Perforce, the Triangle expanded to include not only Red Chris but also the large Au-bearing porphyry deposits of Schaft Creek and Galore Creek.

The Stewart Complex is not truly a metamorphic complex *sensu strictu*. The Mesozoic rocks, while commonly folded, are rarely metamorphosed beyond lower greenschist facies. Similarly, zones of extreme deformation are not common. Considerable work, after the initial discoveries at Johnny Mountain (SNIP and Stonehouse) and at Eskay Creek, was carried out in the northern (Iskut River) area (*e.g.*: Anderson 1989, 1993, Anderson *et al.* 2003, Macdonald *et al.* 1996), in an attempt to resolve the stratigraphy and its relationship to regional and local mineralization.

The oldest strata exposed in Stikinia are Devonian in age and it has been inferred that the basement of the microcontinent is of approximately that age. This has been tested recently by George *et al.* (2021) with the discovery of a small but significant population of pre-Devonian zircons from central Stikinia.

Anderson *et al.* (2003) described northwestern Stikinia as underlain by rocks of at least five Palæozoic to Cenozoic tectonostratigraphic packages. The three lower assemblages comprise multiple, overlapping Late Palæozoic and Early Mesozoic arc assemblages, of which the Late Triassic Stuhini Group is the latest product. These assemblages form a base for the Jurassic arc and basinal assemblages. The Jurassic and older rocks were intruded by the Palæogene post-kinematic granitoid intrusions of the Coast Plutonic Complex.

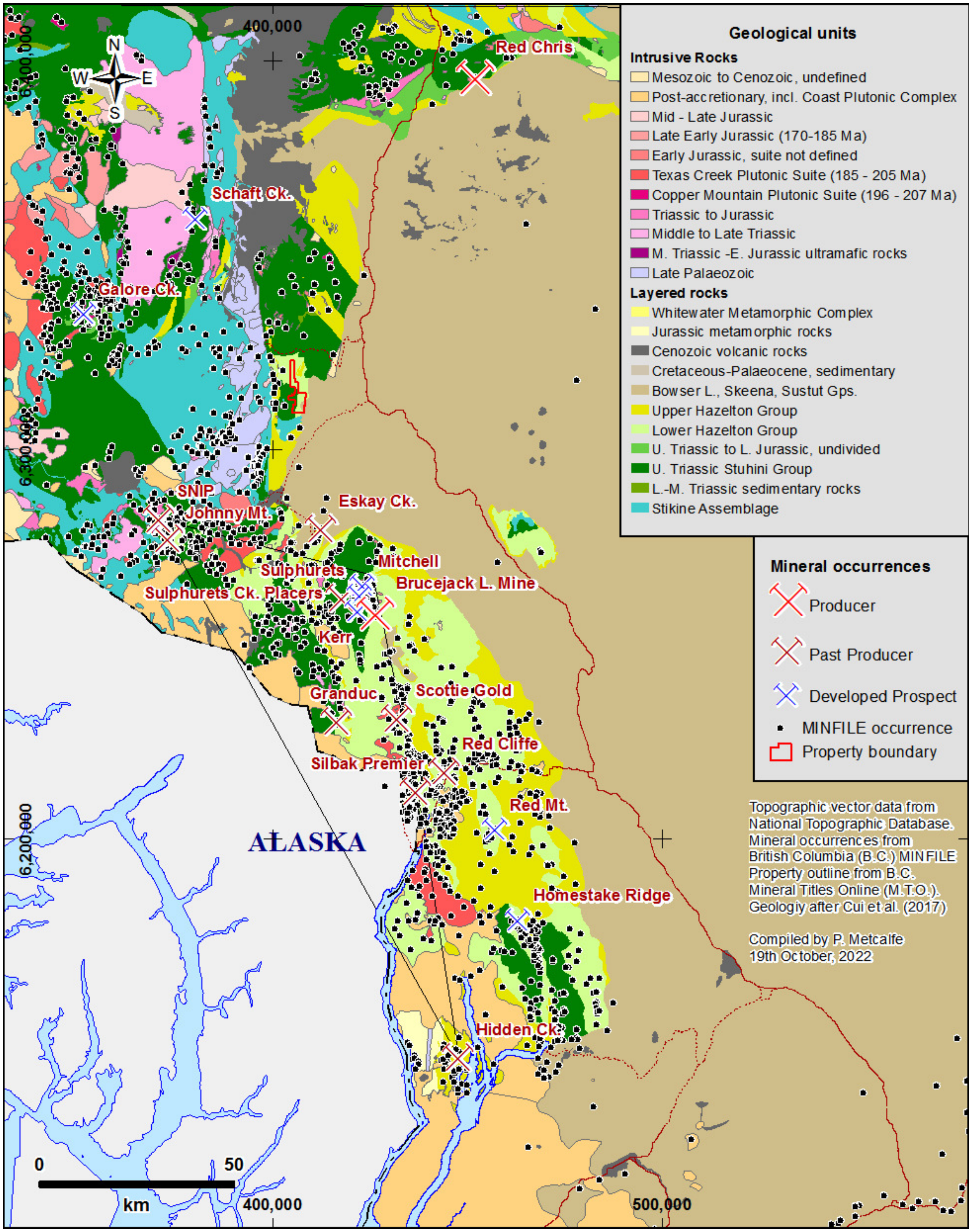

Figure 4. Generalised regional geology after Cui *et al.* (2017).

The area covered is the same as for Figure 1. The outline of the “ancestral” Golden Triangle is presented here for comparison with Figure 1.

Anderson *et al.* (*ibid.*) noted that precious metal deposits discovered to date in northwestern Stikinia are associated mainly with Mesozoic, predominantly Jurassic, arc assemblages. Formation of the island arc assemblages and their associated mineralization occurred during four magmatic episodes, each from 5-10 Ma in duration and bracketed by Triassic-Jurassic, Early Jurassic, Middle Jurassic, and Cretaceous-Eocene deformations. The broad magmatic episodes, together with examples of their derivative mineral deposits, are as follows:

1. Latest Triassic to earliest Jurassic (*ca.* 205-196 Ma) alkaline porphyry-related, deformed mesothermal Ag-Au veins (*e.g.*: Red Mountain);
2. Early Jurassic Texas Creek Plutonic Suite (*ca.* 196-187 Ma) alkaline porphyry-related epithermal, transitional and mesothermal Ag-Au veins and base and precious metal deposits (*e.g.*: Premier, Scottie Gold, Kerr, Sulphurets, Mitchell, Iron Cap, Valley of Kings, SNIP, Stonehouse, Bronson Slope);
3. Latest Early Jurassic (*ca.* 185-183 Ma) small, poorly mineralized porphyry intrusions; and:
4. Middle Jurassic (*ca.* 175-172 Ma) calc-alkaline arc and tholeiitic back-arc magmatism and syn- and epigenetic, stratabound base and precious metal deposits (*e.g.*: Eskay Creek deposit) related to the back-arc basin formation.

The reader will note that the two early magmatic episodes are each roughly 9 million years (Ma) in duration. These periods are quite lengthy and probably reflect overlapping, shorter periods of magmatism at discrete volcanoes. Alldrick (1989) was the first to remark on the necessity for locating Mesozoic volcanic centres as sources of heat and, probably, metal to form the mineral deposits. In particular, from his studies of the Silbak Premier deposit immediately north of Stewart, Alldrick coined the term "Premier Porphyry" to describe the K-feldspar megacrystic phase of the Texas Creek intrusion at Silbak Premier and identifiable as far south as Georgia River (Metcalfe 2011) and as far north as the Red Bluff Porphyry intrusion immediately adjacent to the past-producing SNIP Mine at Johnny Mountain (Atkinson *et al.* 1991, Rhys 1993). Not even the calc-alkaline Island Park Caldera at Yellowstone is large enough to cover the intervening 120 km. Alkaline volcanic centres are, in general, significantly smaller than their calc-alkaline counterparts and it is likely that the terrain between Anyox and Johnny Mountain/Sulphurets hosted several such volcanoes.

The reader will also note that Anderson *et al.*'s demarcation of time periods was designed for the previously defined area of the Stewart Complex (between Iskut and Anyox) and thereby omits an earlier time period 212-201 Ma wherein the Late Triassic Galore Creek and Copper Canyon deposits formed (Henderson 2018). The older Galore Creek intrusions overlap somewhat in time with magmatic assemblages such as the Texas Creek Suite to the south but are petrologically distinct from them. This author infers that the area north of the Iskut River is a transition zone between the influences of these two magmatic suites.

Influence of regional structures on mineralization

A set of structures was noted in the mid-1980s in the Bronson Creek area, with southeast strike and southwest dip. The first structure discovered with this orientation was that hosting the southwest-dipping shear-hosted SNIP Au deposit and, subsequently, that hosting the base metal mineralization of the Bonanza Zone immediately to the southeast from SNIP along structural strike. In 1990, the author confirmed the existence and orientation of the Sky Creek Fault, *en* *echelon* with and in the hanging wall of the SNIP structure (Atkinson *et al.* 1991). In the same year, detailed analysis of bedding surfaces in the Bonanza Zone drill core permitted the author to identify a northeast-vergent anticline with an axial plane coplanar with the SNIP structure (*ibid.*). The folded strata were overlain, with angular unconformity by a flat-lying basal dacite unit, marking the unconformity between Triassic and Jurassic (Metcalfe, unpubl. data 1992).

A distinctive feature of the structure hosting the SNIP deposit is the greater width encountered on more shallowly dipping sections of the structure. The author inferred from this geometry (Metcalfe unpubl. data 2000) that the initial movement on the structure was reverse[2], indicating compression along a southwest to northeast direction. By contrast, Rhys (1993) had shown that ore formation in the same (SNIP) structure occurred during a period of normal[3] movement, indicating extension. At the time there was no evidence that this pattern was regional. However Greig (2014) and Metcalfe (2014) noted that the compressional event was common to Stikinia, uplifting and unroofing plutons as young as Late Triassic, during an event of finite duration in the latest Triassic to earliest Jurassic.

Studies by the author south of Stewart (Metcalfe 2011) found southeast-striking, southwest-dipping major structures which had undergone, at least at one period, dextral[4] movement. Archival geophysical information, once published on the Snipgold II website, confirmed that similar dextral movement occurred on southeast-striking structures at SNIP also. The sequence of movement on these southwest-dipping structures can therefore be mimicked by pushing one's hand up a sloping surface, thence to the left, thence back down the slope of the structure: an anti-clockwise movement. The microcontinent underwent compression, transpression, transtension and then onset of extension within a comparatively short time (Metcalfe 2014, 2021; Metcalfe and Nelson 2014).

Nelson (2014, 2017) described the influence of reactivated basement structures on emplacement of the Triassic and Jurassic intrusions and on their derivative hydrothermal systems, showing in addition that the structures were sufficiently long-lived to have affected emplacement of intrusions as early as the Devono-Mississippian More Creek and Forrest Kerr plutons, close to the Property. The implications for the economic potential of the Property are considerable.

Arc activity ended with deposition of the Middle and Upper Jurassic Bowser Lake Group sedimentary rocks. As noted above, the southwestern margin of Stikinia is bounded by the Palæogene post-kinematic Coast Plutonic Complex.

Regional stratigraphy of the layered Mesozoic rocks

A very brief review of past literature on the stratigraphy of the Iskut camp will be given here. The initial construction of a stratigraphic column for the Stewart Complex was by Grove (1971, 1986). Grove identified three major chronostratigraphic units: the volcanic-rich uppermost Triassic to Lower Jurassic Unuk River formation, the Lower Jurassic, dominantly coarse-grained volcanosedimentary Betty Creek formation and the Lower to Middle Jurassic basinal fill of the Salmon River formation, this latest formation intercalated in places with pillowed andesitic or basaltic volcanic rocks and grading upsection into the Bowser Lake Group.

[2] Reverse movement on a fault or shear zone occurs when upper plate moves up the fault plane relative to the lower.
[3] Normal movement occurs when the upper plate moves down the fault plane relative to the lower.
[4] Dextral movement occurs when the opposite side of a fault moves to the right relative to an observer.

Reassessment of the stratigraphic column took place in the late 1980s and early 1990s. (The reader is reminded that, at the time the original stratigraphic column was created, the concept of plate tectonics was still in its infancy.) As before, the problem remained of accurate chronostratigraphic correlation with few constraining age measurements. As early as 1991, the recommendation was made to replace the increasingly meaningless terms “Unuk River formation” and “Betty Creek formation” with “Lower Hazelton Group” (Peter Lewis, pers. comm.to the author, 1991).

Considerable progress has been made in recent years, leading to substantial changes to the stratigraphic column (Gagnon *et al.* 2012, Nelson *et al.* 2018). The reader is referred to both these publications for a detailed review of the stratigraphy in the Iskut River area. To summarise, the Hazelton Group comprises Lower Hazelton and Upper Hazelton sections, each section with diachronous basal unconformities; basal Lower Hazelton can be as old as uppermost Norian (>205.5 Ma) or as young as basal Sinemurian (≤199.5 Ma); the latter age is indicated in the immediate vicinity of the Property in the Downpour Creek area (*ibid.*). It should be noted that diachroneity of deeply incised erosional surfaces will interfere to a major extent with the location of depth-controlled mineralization, irrespective of deposit type, whether porphyry, transitional, epithermal or sea-floor.

7.2 Local and Property geology

A geological map of the area of the Property, adapted from Alldrick *et al.* (2005a), is shown in Figure 5 (with legend in Figure 6). The Property lies across the valley of Downpour Creek, between the *massifs* of Lawrence Peak and Faisal Peak. West of the Faisal Peak and the headwaters of Downpour Creek is the north-south trending major structure of the Forrest Kerr fault which runs along Forrest Kerr Creek upstream from its confluence with the Iskut River.

The Forrest Kerr fault is a profound structural break, separating Palæozoic supracrustal and intrusive rocks to the west from a comprehensive section of the Upper Triassic to Middle Jurassic volcanosedimentary rocks east of the fault (*c.f.* Nelson 2014, 2017). Immediately east of the Property boundary, another rapid transition occurs, passing downsection to the east from Middle Jurassic siliciclastic facies of the Iskut River formation through Lower Jurassic Lower Hazelton volcaniclastic rocks to Upper Triassic Stuhini Group. Thus, the Property area is an inlier of Upper Hazelton Group sedimentary rocks bounded to east and west by older strata.

Alldrick *et al.* (2005b.) noted that: “The internal stratigraphy of the Downpour Creek facies is not completely understood.” Notwithstanding, the palæontological work in Nelson *et al.* (2018) constrains the base of the Upper Hazelton sedimentary sequence to latest Toarcian. The sedimentary section up through the Aalenian to the Lower Bathonian is preserved in the Property area. This is the stratigraphic equivalent of the host sedimentary rocks at Eskay Creek.

While Alldrick *et al.* (*ibid.*) report hornblende + plagioclase porphyritic, K-feldspar megacrystic intrusions correlated with Texas Creek Plutonic Suite well to the south of the Property, no intrusive rocks are reported by Government mapping on the Property itself.

7.3 Significant mineralized zones

No significant mineralized zones have been described on the Property to date.

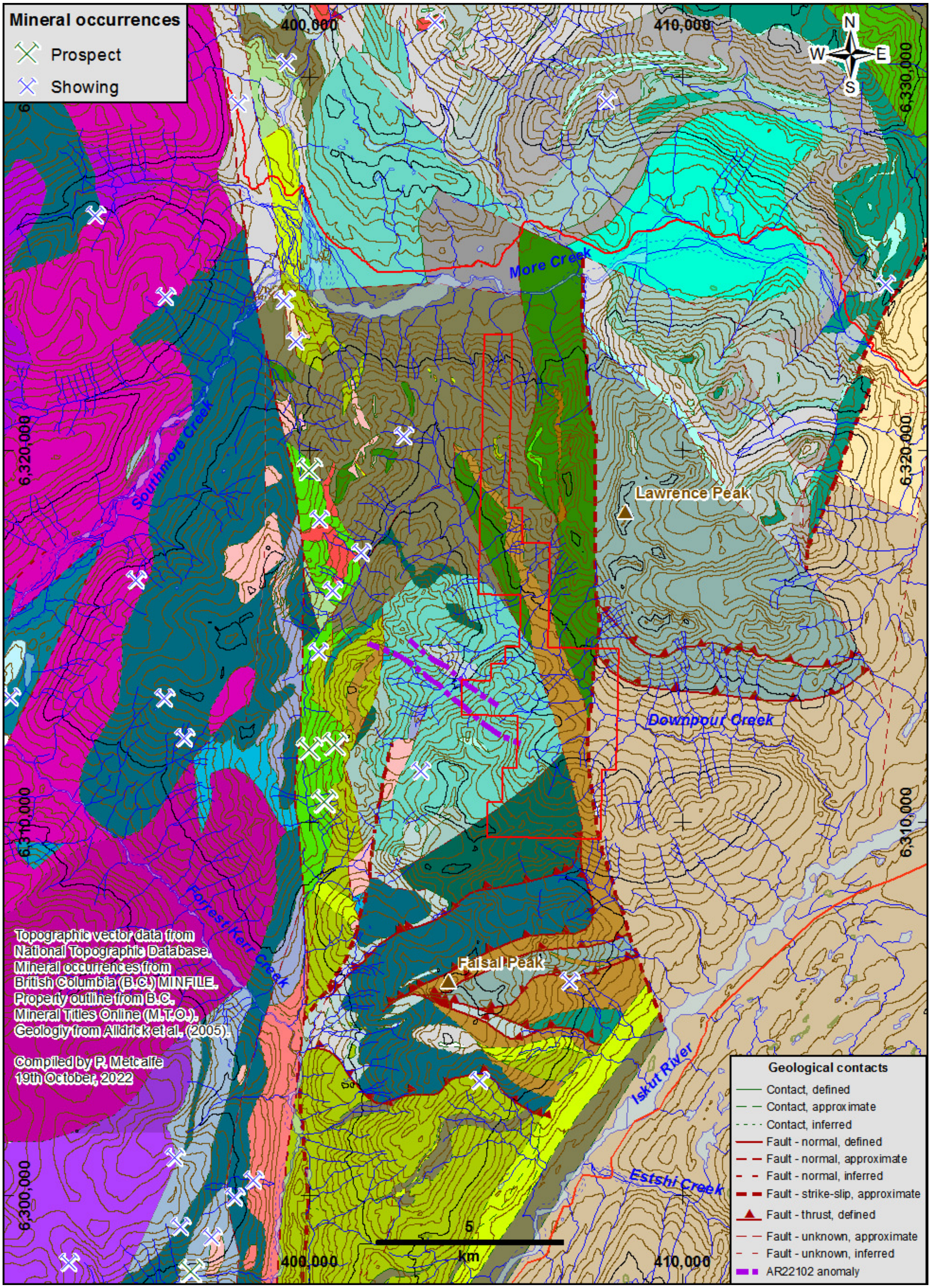

Figure 5. Geological map of the Property (after Alldrick *et al.* 2005a).

Figure 6. Legend for Figure 5.

8.0 DEPOSIT TYPES

It is instructive, when an advanced project is the subject of a technical report, to consider the data available for the project in the context of various deposit models and to comment on the various models' conformity with these available data. As new data become available, the appropriateness of a particular model may be challenged and a new model adopted for the project.

By contrast, assignment of a model or models for mineralization yet to be discovered is extremely difficult for an early stage project with little available data. A brief review of the inferred deposit types in the area will be given here, as a general guide only. The reader is cautioned not to be constrained by this list; for example: the "subaqueous hot spring" deposit type was only formulated in the wake of the 1989 Eskay Creek discovery.

The area defined by the Golden Triangle is host to more than 1000 mineral deposits identified in MINFILE (Geological Survey Branch 1991-2022). More than 700 are gold-bearing. Of these more than 10% have not been assigned a British Columbia Mineral Deposit Profile. Fully two thirds of the remainder are identified as "veins" albeit a large number of these are identified as "intrusion-related". The balance of mineral occurrences range in inferred deposit type from porphyry through vein and epithermal to near-surface types; these last, for convenience, will be grouped as "exhalative".

The MINFILE database is not invulnerable to bias based on previous reporting. Following the discovery of Eskay Creek, a noteworthy number of Assessment Reports by prospectors identified the lithologies on their properties as correlative with the stratigraphy at Eskay Creek. This was entirely understandable, but had a tendency to skew the database.

The Eskay Creek Mine remains one of the most significant and highest grade deposits discovered in the Golden Triangle. The deposit formed in the latest of the metallogenic cycles in the area (latest Toarcian to earliest Bathonian), in a rift environment and is identified as a precious metal-rich volcanogenic massive sulphide deposit. The deposit has associated epithermal characteristics with it, insofar as gold deposition is credited to boiling of the ore fluid at or near the sediment-seawater interface (Sherlock *et al.* 1999). The recent efforts by government to track the rift to the north discovered that it ran precisely through the area covered by the Property (Alldrick *et al*. 2005a, b). As noted above (Section 7.2), the Eskay Rift on the Property contains strata equivalent in age to those hosting the Eskay Creek deposit (Nelson *et al*. 2018).

Epithermal deposits are not confined to the latest metallogenic cycle. A large profile of intrusion-related deposits exists in the Golden Triangle, particularly vein-type deposits such as those at Johnny Mountain. These exhibit characteristics ranging from epithermal at higher elevations to mesothermal at lower elevations. The reader is reminded that, in this area of the world, the topography permits exposure of a section as much as 2.5 km in vertical distance through a mineralized system, without need for fault displacement. Notwithstanding the spectacular mineralization at Eskay Creek, epithermal mineralization is more typical of the lithologies above and below the Rhaetian-Hettangian sub-Hazelton unconformity in the immediate area and is intrusion-related.

The Early Jurassic intrusions in the central and southern parts of the Golden Triangle, to which many (if not most) of the veins are related, compose a suite of hornblende+plagioclase-porphyritic monzonites: the Texas Creek Plutonic Suite. Characteristically, these are polyphase intrusions which commonly include a distinctive K-feldspar megacrystic phase (*e.g.*: the “Premier Porphyry” of Alldrick 1985). These intrusions vary in age from 186 Ma to as old as 196 Ma and include the Red Bluff Porphyry, the metallogene not only for the shear-hosted veining of the SNIP deposit, but also for the adjacent porphyry mineralization of Bronson Slope. The porphyry deposits of Kerr, Sulphurets and Mitchell (KSM) are also part of the Texas Creek metallogenic event.

Porphyry mineralization in the northern Triangle is not confined to intrusions of the Texas Creek Plutonic Suite. The Late Triassic potassic undersaturated intrusions of the Galore Creek area, absent in the central and southern Triangle, entirely predate the Rhaetian unconformity. The implications for exploration in the area of the Property, where the north-trending Forrest Kerr Fault is evidently one of Nelson’s (2014, 2017) basement structures, are evident.

Of well over a thousand mineral occurrences in the Golden Triangle, less than fifty are identified in MINFILE as having skarn characteristics. Of these, no more than six have progressed in development stage beyond the status of “showing”. While this deposit type should not be discounted, it has low priority as an exploration target type for the present Property.

9.0 EXPLORATION

Fieldwork was carried out between 9th July 2022 and 5th August 2022; the author was present on the Property from the 9th of July to the 19th of July. Traverses were conducted by the author on four of the six mineral tenures extant at the time (two tenures were acquired at the south end of the Property subsequent to the author’s fieldwork). Geological mapping and geochemical sampling of accessible streams were carried out during the course of fieldwork. A total of 43 samples (16 rock, 10 heavy mineral and 17 stream sediment) were submitted by Romios to ALS labs in Vancouver and to ActLabs in Ontario (see Section 11).

The preliminary fieldwork served to confirm the presence of stratigraphy represented by Alldrick *et al.* (2005a, b) and Nelson *et al*. (2018). The northern part of the Property is underlain by a sequence of dark grey shale, siltstone and wacke, correlated (*ibides*) with Upper Hazelton Group. The sedimentary rocks are folded into a series of anticline-syncline pairs. Intercalated with the finer clastic sedimentary rocks are minor thicknesses of coarser clastic rocks interpreted as felsic volcaniclastic rocks, albeit with few signs of hot emplacement, implying that their deposition was distal from a volcanic centre (*c.f.* Alldrick 1989).

The sequence rests on a coarser clastic unit identified by Nelson *et al.* (*ibid.*) as the basal Kinaskan Conglomerate of the Upper Hazelton Group. 2022 fieldwork suggests that this conglomerate is not as extensive as mapped, but Nelson *et al.* note that the conglomerate thins to the north, grading laterally and upsection to finer clastic sedimentary rocks.

Two minor addenda are possible to Alldrick *et al.*’s map (*op.cit.*). The first is the presence of an extensive, thick fracture zone with southeast strike and moderate southwest dip crossing the southwest ridge of the Property (Figure 7). This fracture zone is cospatial with a geophysical anomaly detected during previous exploration (Fields 1992). The interpretative map from this historic geophysical survey is shown in Figure 7, with the discontinuities highlighted.

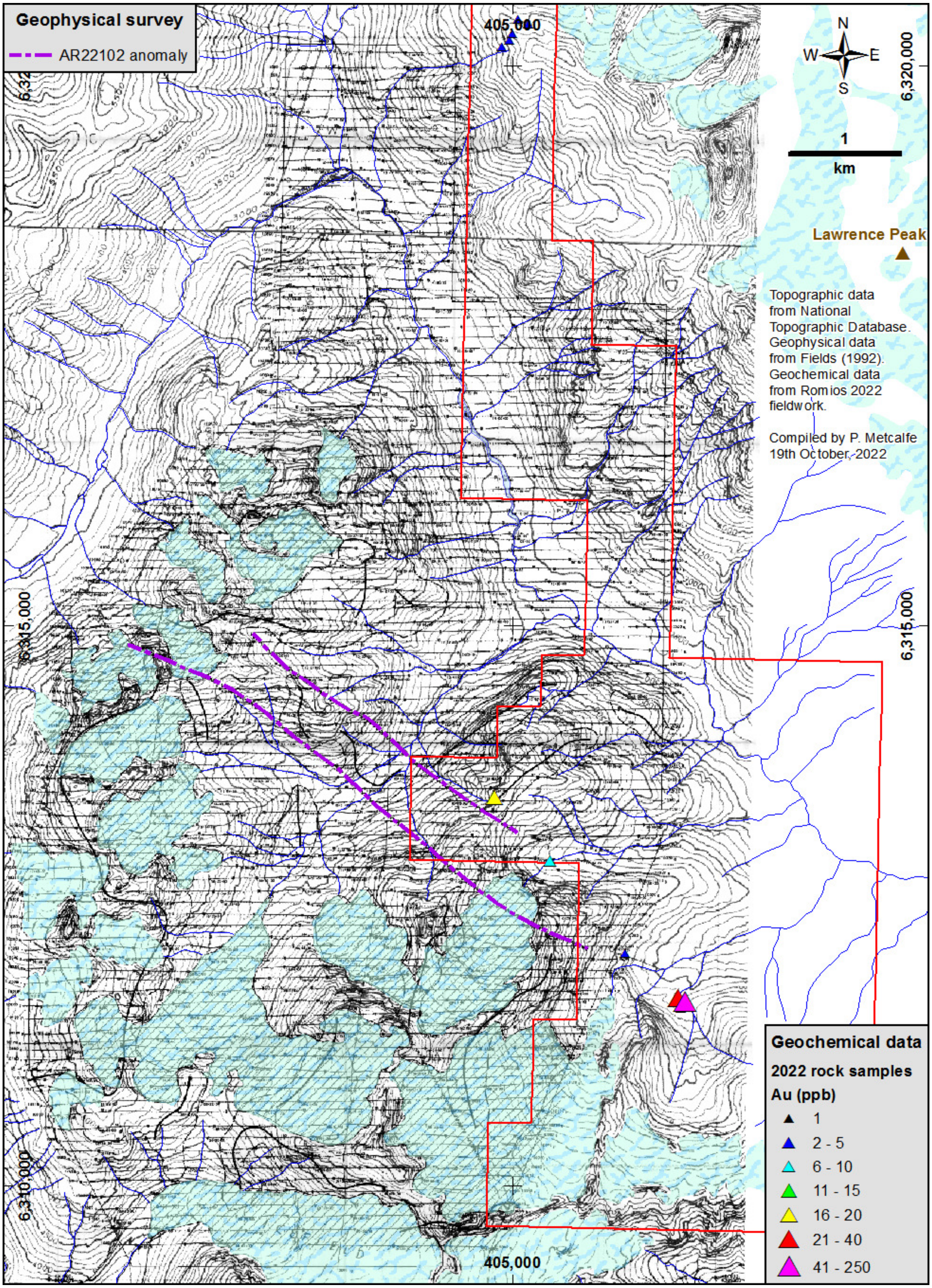

Figure 7. 2022 Rock samples overlaid on historic airborne geophysical survey (Fields 1992).
N.B.: The geophysical data are from a historic report and have not been verified.

More significantly, the author located a single small area of proximal float, possibly subcrop, of K-feldspar megacrystic porphyry on the same southwestern ridge, near where weak anomalous arsenic (As) was returned from a talus fine sample. The occurrence lies between the traces of Fields' geophysical discontinuities. Its significance was only realized after float of similar lithology was discovered in the cirque at the eastern end of the more southerly discontinuity. This lithology is designated herein by the author as the Black Cat Porphyry.

The Black Cat Porphyry (Figure 8) is a crowded K-feldspar porphyry, containing as much as 45% euhedral, thin tabular K-feldspar megacrysts as large as 50 mm, seriate with a microcrystalline groundmass, commonly dark grey or black in colour. The lithology may also contain as much as 10% relic phenocrysts with outlines distinctive of prismatic amphibole, pervasively altered. Megacrysts are peripherally corroded and, particularly in the more crowded areas are aligned which can give the whimsical impression of cats' eyes, hence the name.

This lithology, despite its superficial resemblance to K-feldspar-megacrystic phases of the Texas Creek Plutonic Suite, is not a plagioclase porphyry. It most closely resembles intrusive lithologies exposed at Burgundy Ridge (Enduro Metals) and at Galore Creek (Henderson 2018) and which historically have been identified as hosts for copper (and gold) mineralization (Souther 1971). Significantly, these intrusions, at Galore Creek, are of Late Triassic age (212-200.7 Ma; *ibid.*). If that were the case on the Property, which the author considers likely, such an intrusion (and associated mineralization) would predate the entire Hazelton Group section.

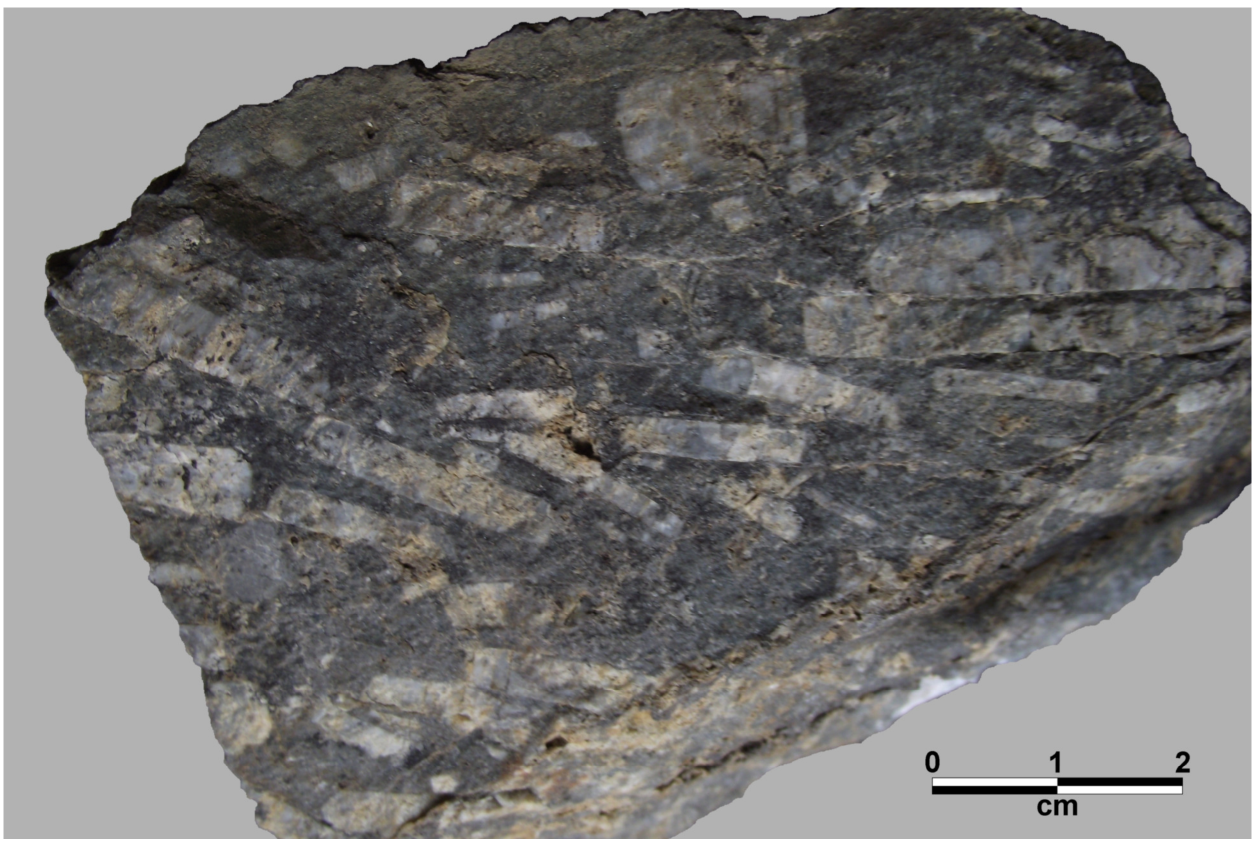

Figure 8. Black Cat Porphyry.

The rock is a true K-feldspar porphyry, unlike the plagioclase porphyries of the Texas Creek Plutonic Suite

No significant mineralization was encountered during examination of the northern and central parts. Float discovered on its southwestern part was mineralized with pyrite and returned weakly anomalous Au (Figure 7).

Geochemical rock samples taken are shown in Figure 7. Samples were generally low in precious metals and, indeed, in base metals. However, elevated values of arsenic (As) were detected in the sample of talus fines (PM22-055) taken on the ridge crest at the centre of the southeasterly striking fracture zone. Elevated As was also associated with the two anomalous Au values taken beyond the eastern limit of the 1992 geophysical survey (Figure 7).

Stream sediment sampling in the valley bottom was tested by the author in a single traverse. It was discovered that the swath peripheral to Downpour Creek, mapped as Qal (Quaternary alluvium) by Alldrick *et al.* (2005a) represents only a part of the Quaternary deposits in the valley. A deposit of glacial material, not shown in Figure 5, extends at least 500 m to either side of Downpour Creek up the lower slopes of the valley side. This lack of outcrop in the valley bottom is a considerable hindrance to mapping and, in the author's opinion, precludes the use of direct (endogenic) geochemical sampling, owing to the transported nature of the material.

Stream sediment and heavy mineral concentrate samples taken in the valley bottom and in areas of thick moraine in adjacent cirques were low in concentrations of both Au and the pathfinder element As. The author has confidence in these 2022 data, but notes that the time constraints precluded comprehensive coverage of the Property, particularly on the valley sides.

Owing to these sparse 2022 data, the author compiled data from historic studies (Nicholson 1990a, b, Gal and Walker 1990, Walker and Moore 1990). These data cannot be verified, owing to their age, the lack of published blanks or standards and the different method of analysis (dissolution in perchloric and nitric acid, followed by inductively coupled plasma emission spectroscopy (ICP-ES). Nevertheless, the results are informative and are presented in Figure 9.

The historic data are of sufficient density to illustrate a pattern. An elevation in As values occurs proximal to the southeast-striking anomalies noted by Fields (*op. cit.*) and to the fracture zone discovered during 2022 exploration. The reader is invited to compare this pattern with the regional pattern visible in the digital elevation model shown in Figure 1. It is apparent that north-vergent, southeast-striking structures described by the author in 2014 (*opera citata*) are also present north of the Iskut River. These are powerful vectors for mineralization.

To summarise, 2022 fieldwork returned generally low values from geochemical sampling, consistent with previous studies (in 1990) on ground enclosed by the present Property. Moreover, the areas selected for geological investigation confirmed, with two exceptions, previous geological mapping. The two exceptions were a southeast-striking, southwest-dipping fracture zone, probably a fault, spatially associated with a subcrop or proximal float of a K-feldspar megacrystic porphyry. The close lithological similarity of this intrusive lithology with an intrusive phase explicitly associated with the Galore Creek deposit is remarkable.

Significantly, the locations of weak geochemical anomalies in Au and As encountered in 1990 and 2022 are cospatial both with the location of the geophysical discontinuities identified in 1992 and with the southeast-striking fracture zone containing subcrop or proximal float of the K-feldspar megacrystic porphyry. This coincidence has never been remarked before and certainly warrants further investigation.

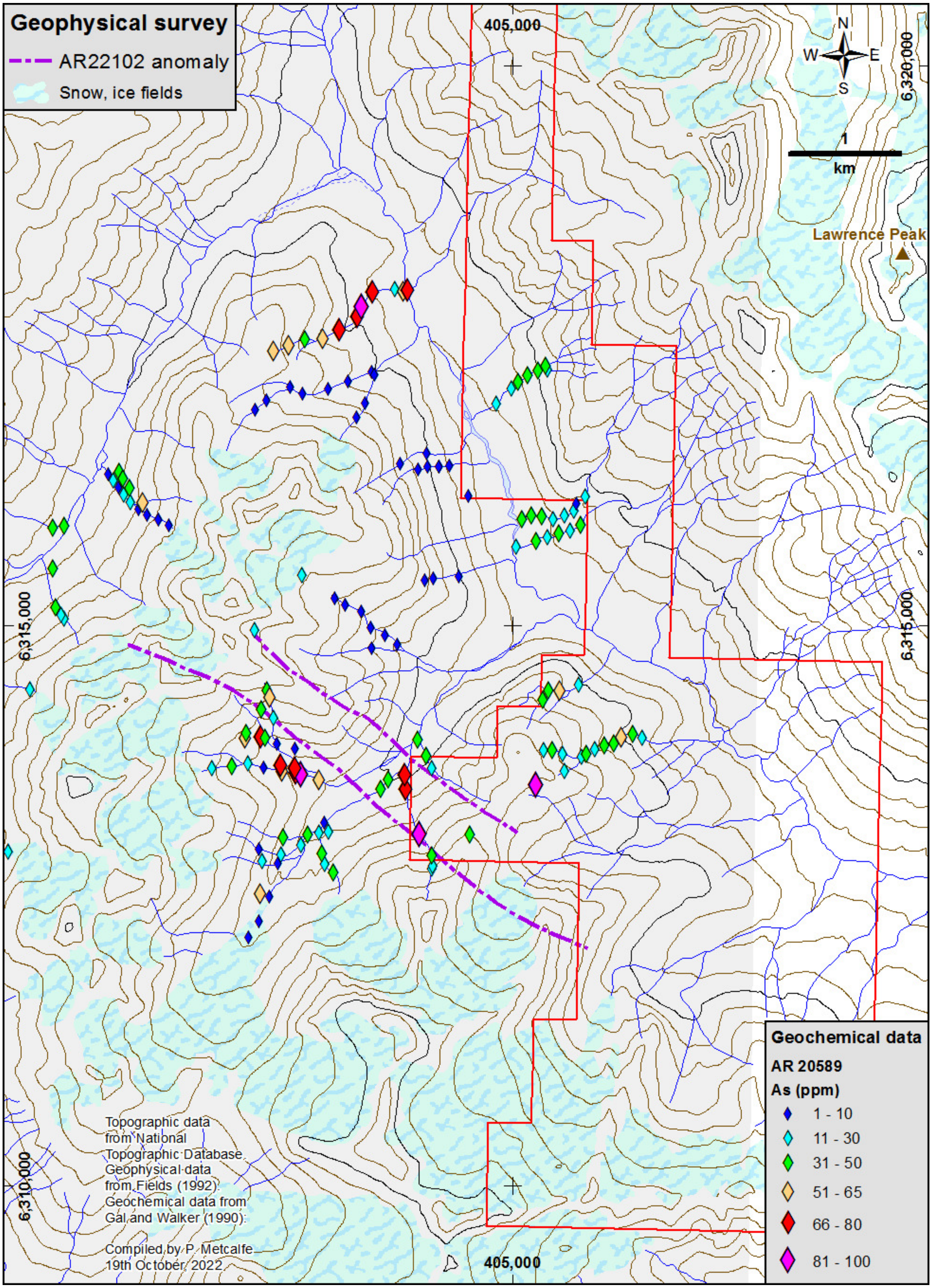

Figure 9. Historic As concentrations in stream sediments, from previous (1990) reports. Data from Nicholson (1990a, 1990b), Gal and Walker (1990), Walker and Moore (1990).

10.0 DRILLING

There has never been a diamond drill project conducted on any part of the Property.

11.0 SAMPLE PREPARATION, ANALYSES, AND SECURITY

Stream sediment samples and heavy mineral concentrates were dried insofar as conditions in camp would allow. No other field procedures were used prior to sample shipping. Samples were shipped, one batch to ALS Analytical Laboratories in North Vancouver. The remaining rock samples and all stream sediment and heavy mineral concentrate samples were shipped to an ActLabs analytical facility in Ancaster, Ontario. The accreditations of the laboratories are, respectively:

ALS: Accredited Testing Laboratory ISO/IEC 17025:2017 and ISO 9001:2015

ActLabs: Accredited Testing Laboratory (ISO/IEC 17025:2017): No.266

Neither laboratory has any relationship to Copperhead Resources Inc., to Romios Gold Resources Inc. or to Palatine Geological Ltd.

Stream sediment samples and heavy mineral concentrates were dried, then sieved to -80 mesh. Analysis was by induced neutron activation analysis (INAA). Rocks were crushed and a split of the crushed rock pulverized to -180 mesh. Analysis at ALS was by *aqua regia* digestion followed by inductively coupled plasma-mass spectrometry (ICP-MS). Analysis at ActLabs was by INAA.

“Natural” blanks (unmineralized rock), unmarked replicates and unmarked standards (OREAS 505 for rocks, OREAS 47 for stream sediments) were inserted with the samples to be analysed.

The sample preparation and analytical procedures for samples from 2022 exploration were carried out by qualified personnel at internationally certified laboratories. The author has examined their relevant certifications and is satisfied with the certifications regarding sample preparation and analytical procedures. Concentrations of pathfinder elements returned from analysis of standards and lab blanks are consistent with, respectively, published concentrations for the standards and below-detection limits for the blanks. While “natural” blanks returned above detection limits in pathfinder elements, there were no anomalously high values.

Regarding sample security, the author has drawn from the historic work (Nicholson 1990a, 1990b, Gal and Walker 1990, Walker and Moore 1990) carried out in the area prior to the institution of National instrument 43-101. The close similarity between the historic data and results of the present sampling indicate that either both sample sets are representative or both are in error by identical amounts. The former hypothesis minimises entities and is preferred.

12.0 DATA VERIFICATION

The author is not a geophysicist and took no steps to verify the data used in the preparation of Fields (1992), beyond a preliminary attempt to retrieve the original data file, for later review by a qualified geophysicist. The author will remark that the anomalies reported (*ibid.*) are consistent with the pattern of geochemical anomalies reported in historic surveys and with present geological mapping.

As noted above, the author has examined, where available, the historic data in the assessment reports and has noted that these data bear a very strong resemblance to those obtained during 2022 exploration. In brief, the historic geochemical data validate the security measures taken for the 2022 data and the 2022 data at least support the general integrity of the historic data.

The author is of the opinion that the data available both from previous exploration and from 2022 exploration provide adequate grounds for the opinions expressed in this technical report.

13.0 MINERAL PROCESSING AND METALLURGICAL TESTING

Neither mineral processing nor metallurgical testing analyses have been carried out, nor are any contemplated at the time of writing.

14.0 MINERAL RESOURCE ESTIMATES

There is neither presently, nor in the past, a mineral resource estimate of any kind on any part of the Property, nor is any such estimate contemplated in the near future.

15.0 MINERAL RESERVE ESTIMATES

There is neither presently, nor in the past, a mineral reserve estimate of any kind on any part of the Property, nor is any such estimate contemplated in the near future.

16.0 MINING METHODS

Absent a significant discovery, mining methods are not contemplated at this stage.

17.0 RECOVERY METHODS

Recovery methods are not contemplated for an early stage property such as Red Line.

18.0 PROJECT INFRASTRUCTURE

For an early stage property such as Red Line, infrastructure and logistic requirements for the project (roads, port facilities, location of power lines) have already been discussed in Section 5.

19.0 MARKET STUDIES AND CONTRACTS

Neither market studies nor production contracts are presently contemplated (no production).

20.0 ENVIRONMENTAL STUDIES, PERMITTING AND SOCIAL OR COMMUNITY IMPACT

Despite being an early stage exploration project, a brief discussion of the impact of permitting and environmental requirements is appropriate here. As noted above, non-intrusive exploration activities do not require a permit. There are presently no pending permit applications, nor are any required for the recommended work (see Section 26: "Recommendations"). However, as noted in Section 4, mineral tenure does not confer surface rights, therefore any form of camp construction, trenching or drilling on the Property or even performance of an induced polarization survey requires a permit. Permit application is through the government of British Columbia and such applications are routinely referred to interested parties, including local communities. The Government of British Columbia recommends that permit applicants engage with interested parties early in exploration, rather than relying on government.

Notwithstanding regulatory requirements, a baseline environmental assessment is often carried out prior to intrusive work. The cost can easily be absorbed by making collection of samples part of an exploration stream sediment sampling program.

21.0 CAPITAL AND OPERATING COSTS

There are no planned activities on the Property that require capital and operating cost estimates.

22.0 ECONOMIC ANALYSIS

The Red Line property is in the earliest stages of operation and no economic analysis is possible.

23.0 ADJACENT PROPERTIES

The Property is bordered to west, north and east by three claim blocks owned, respectively, by Aben Resources Ltd., Galore Creek Mining and Tower Resources Ltd. The ground covered by Galore Creek Mining comprises a near-continuous swath of tenures which extend from Highway 37 to the developed prospect of Galore Creek 50-60 kilometres northwest of the Property. Most of the work carried out there is geotechnical and not immediately relevant to the Property.

To the west of the Property, the mineral tenures held by Aben Resources Ltd. cover ground along the north-south trending Forrest Kerr Fault. The reader will note that the majority of the mineral occurrences in the area, including all those elevated to the status of “prospects” are proximal to that structure; the activities of Aben Resources have, naturally, been in that area.

To the east of the Property, there is a rapid transition downsection, in part stratigraphic, in part owing to north-south faulting; the rocks exposed around Lawrence Peak are Triassic in age. Information from Tower Resources’ property would be of restricted use except with regard to regional structure.

The author is aware of roughly 70 assessment reports describing work carried out for assessment purposes on mineral tenures in the area. Of the tenures extant in the area and to the best of the author’s knowledge, no assessment work has been published on adjacent properties that would add information material to this Property. Five historic studies carried out in the immediate area of the Property have already been disclosed in Section 9 (Exploration).

24.0 OTHER RELEVANT DATA AND INFORMATION

The author is not aware of any other information relevant to the technical report.

25.0 INTERPRETATION AND CONCLUSIONS

The Red Line property is located in steep, glaciated topography 18 km west of Highway 37 in the eastern Boundary Ranges of British Columbia's Coast Mountains. Owing to the presence of major drainages on three sides of its immediate physiographic area, access is, perforce, by helicopter only. The Property is within 15 km of a major powerline and, were a significant discovery made, access would be trivial by contrast with Brucejack Lake or Galore Creek.

The area is part of British Columbia's Golden Triangle, home to more than a thousand mineral occurrences, including past-producing precious metal mines and two extant producing gold mines. The Triangle is part of the ancient microcontinent of Stikinia, whose geological history includes at least three major Mesozoic metallogenic events. Lithologies from each of these time periods were mapped on the Property by government (Alldrick *et al.* 2005a, b).

The immediate area of the Property was shown (*ibides*) to lie within the northerly extension of the Eskay Rift which, south of the Iskut River, hosts the world-class Eskay Creek deposit. Geological fieldwork by Romios (Biczok 2021 and the present study) confirmed the presence of this Upper Hazelton sedimentary section in the northern part of the Property. However, four conjugate circumstances have hindered exploration in this valley and, in the author's opinion, have contributed to the lack of MINFILE occurrences recorded to date in the area enclosed by the Property.

The first of these circumstances is topographic. The steepness of the valley sides constrained the regional mapping of Alldrick *et al*. and the (to date brief) exploration of Romios to the ridge crests. As a consequence, comprehensive examination of the entire Upper Hazelton section, crucial in exploration for any stratabound deposit, has never been carried out in this area.

The second circumstance is the presence of northerly trending faults along the valley, juxtaposing the prospective (for Eskay Creek-type deposits) Upper Hazelton section against Lower Hazelton and older strata. Vertical movement on such faults has the potential to create a "blind" exploration target, *i.e.*: one not exposed at surface. This is readily apparent in the southeast corner of the Property, where younger, Middle to Upper Jurassic sedimentary rocks are juxtaposed against basal Upper Hazelton rocks (Figure 5).

The third circumstance is the presence of significant deposits of glacial material, of unknown thickness, extending at least 500 m to either side of Downpour Creek. Not only do these deposits have the potential to cover prospective bedrock, they comprise transported material which can occlude any endogenic geochemical anomaly derived from an underlying deposit.

The fourth circumstance is conceptual. The 1989 discovery of the world-class Eskay Creek deposit created a substantial shift in exploration philosophy towards exploration of rift-type environments in Upper Hazelton stratigraphy both north and south of the Iskut River. In particular, attention was drawn to exploration activities in the area of the Forrest Kerr Fault (Figures 3 and 5), around which the majority of historic, local exploration activity took place. Exploration at remove from the fault was, perforce, of lower intensity and, at least in its later stages, focused exclusively on an Eskay Creek model (*c.f.* Fields 1992). Owing to this, the southwestern part of the Property itself has received disproportionately less historic attention.

Previous geophysical work (Fields 1992) identified two southeast-trending geophysical discontinuities, shown in Figures 7 and 9. Significantly, these discontinuities are not represented on geological maps of the area. The discontinuities are coincident with a wide structural break encountered during 2022 fieldwork, with a zone of elevated As values in 1990 stream sediment samples and with topographic fabric visible in regional digital elevation data. This zone and the attendant As anomaly are interpreted as the uppermost parts of a mineralized conduit.

2022 fieldwork confirmed the general findings of regional mapping (Alldrick *et al, opera citata*) in the north of the Property, albeit with considerable reduction in observed thickness of the basal Kinaskan Conglomerate. In addition, little evidence of hot emplacement was observed in the volcanogenic rocks assigned (*ibides*) to the Lower Hazelton Group. With the exception of weakly anomalous As, precious and base metal values were low in both rock and stream sediment samples. Similarly low values, with the exception of arsenic, had been returned from historic (1990) regional sampling. As noted above, this may owe in part to the presence of extensive glacial deposits on the valley floor and braes of Downpour Creek. However, float of intrusive rock discovered on the southwestern part of the Property is mineralized with pyrite and is weakly anomalous in Au and lies along the structural break identified by 2022 mapping.

Fieldwork located a single subcrop or proximal float of K-feldspar megacrystic porphyry on its southwestern ridge, near the trace of the southeast-trending structural zone. This lithology most closely resembles intrusive lithologies exposed at Burgundy Ridge and at Galore Creek and which historically have been identified as hosts for copper (and gold) mineralization. Significantly, these intrusions are of Late Triassic age (212-200.7 Ma) and therefore predate the entire Hazelton Group section. They are also mineralized:

> *"The alkaline intrusions (unit 12), including the Galore Creek body, that occupy a broad north-south belt through the west-central part of the map-area exhibit many features in common with porphyry copper stocks. Traces of copper are present in all the mapped bodies, particularly where the rock is highly fractured. The possibility that other alkaline stocks are present in this zone makes it particularly interesting for further exploration."*
>
> J. G. Souther (1971)

If the author's inferences regarding the K-feldspar porphyry are correct, then any mineralized system deriving from a Triassic intrusion will be covered, in the northern half of the Property, by Jurassic supracrustal rocks, deposited after the Rhaetian-Hettangian unconformity described by Nelson *et al.* (2018). In such a circumstance, the overlying Jurassic rocks would not be mineralized.

Notwithstanding the previous paragraph, the Property lies within the Eskay Rift and stratified rocks equivalent in age to those hosting the high-grade, past-producing Eskay Creek deposit are inferred to be present in its northern part. At the time of writing, the author is not aware of any protracted discussion regarding the source of the gold so effectively concentrated by the Eskay Creek mineralizing system. It is therefore proposed here that likely sources are the metallogenic intrusions of the Texas Creek Plutonic Suite or the K-feldspar-megacrystic intrusions correlated with those at Galore Creek. If so, the Eskay Rift at this location incorporates sufficient faults as to act as efficient plumbing for any remobilisation of Au.

Strata as young as Bathonian occur on the Property, well upsection from Eskay Creek-equivalent stratigraphy. In addition a thick blanket of transported syn-and post-glacial Quaternary deposits covers the valley bottom and sides. Future exploration should proceed on the presumption that mineralization formed during any of the three major Mesozoic metallogenic events may occur on the Property as a blind target, to be located by geophysical and exogenic geochemical methods.

26.0 RECOMMENDATIONS

The following are recommended:

1. Consultation with independent experts in geophysics and in exogenic geochemistry, for the purposes of recommending optimal techniques for the location of blind, precious metal-bearing exhalative deposits, porphyry-style deposits and intrusion-related precious metal vein deposits under glacial and bedrock cover in mountainous terrain;
2. Planning and execution of the recommended geophysical survey.
3. Isotopic age measurement and geochemical analysis of a sample of Black Cat Porphyry where located in outcrop, using U-Pb laser ablation mass spectrometry on zircon to allow for determination of the primary oxidation state of the magma (and thereby fertility);
4. Detailed mapping of the Property, with particular attention to intrusive rocks and any associated alteration, accompanied by intensive prospecting;
5. Creation of a test line for evaluation of various exogenic geochemical methods for locating blind porphyry, vein and Eskay Creek-style exhalative targets and:
6. Planning and execution of the recommended exogenic geochemical survey.

A budget for the recommendations is presented in Table 2. The second phase would be contingent on positive results from the first, particularly results from the geophysical survey.

Table 2 Proposed budget for exploration.

Phase 1					
Work Type	**Item**	**Cost/unit (CD$)**	**# of units**	**Units**	**Cost (CD$)**
Isotopic analysis	Isotopic age measurement	1500	1	Analysis	1,500
Consultants	Geophysical and geochemical consultation	1500	4	Days	6,000
Airborne geophysics	Deep- penetration airborne electromagnetic and magnetic survey[5]	250	250	Line km	62,500
Geochemical (exogenic)	Test geochemical line	300	100	Analyses	30,000
	Total				**100,000**
Phase 2					
Work Type	**Item**	**Cost/unit (CD$)**	**# of units**	**Units**	**Cost (CD$)**
Geological mapping and prospecting	Mapping crew	2000	56	Days	112,000
Crew maintenance	Camp costs	300	56	Days	16,800
Exogenic geochemistry	Sampling and analysis	100	1200	Analyses	120,000
Mobilization	Crew travel	1700	4	People	6,800
	Total				**255,600**

[5] The reader should note that the author is not a Qualified Person in the field of airborne geophysical techniques, hence the recommendation for consultation with an independent Qualified geophysicist

ACKNOWLEDGEMENTS

The author is indebted to the Romios crew, Hanna Tiitto, Jacob Mohr-Wise, Kieran Kristoffersen and Daniel Thomson, for their tolerance of bad weather and (in Jacob's case) bad bush; also for their help with samples and with demobilization in fraught conditions. Thanks are also owing to Chris Blount and Shane Menzel of Aberdeen Helicopters, who got us there and brought us home. The author is also grateful for the hospitality of the Matrix team at Truffle Camp.

Ms. Marie Brannstrom is thanked for her rapid review of the preliminary document. In particular, the author wishes to acknowledge, with thanks, the scientific review and suggestions from Romios' Vice-President of Exploration, John Biczok; these suggestions greatly benefited this document.

Lastly, this report is based on the findings of more than two generations of researchers from academe and geological surveys, both federal and provincial. Notwithstanding all these contributions, this author takes full responsibility for any misinterpretation of their work herein.

27.0 REFERENCES

References cited in the text are emboldened; the remainder are presented for the reader's convenience.

Alldrick, D.J., 1985: Stratigraphy and Petrology of the Stewart Mining Camp; *in*: Geological Fieldwork 1984, B.C. Ministry of Energy, Mines and Petroleum Resources Paper 1985-1, pp.316-342.

Alldrick, D.J., 1987: Geology and Mineral Occurrences of the Salmon River Valley, Stewart Area (104A, B); B. C. Ministry of Energy, Mines and Petroleum Resources Open File 1987-22.

Alldrick, D.J., 1989: Volcanic centres in the Stewart Complex; *in*: Geological Fieldwork 1988; British Columbia Ministry of Energy. Mines and Petroleum Resources Paper 1989-1, pp.233-240.

Alldrick, D.J., 1993: Geology and Metallogeny of the Stewart mining Camp, Northwestern British Columbia; British Columbia Ministry of Energy, Mines and Petroleum Resources, Geological Survey Branch Bulletin 85, 105p., 2 maps.

Alldrick, D.J. and Britton, J.M., 1988: Geology and Mineral Deposits of the Sulphurets Area (NTS 104A/5, 12; 104B/8, 9); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1988-04.

Alldrick, D.J. and Britton, J.M., 1992: Unuk River Area Geology (NTS 104B/7E, 8 & 9W, 10E); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1992-22.

Alldrick, D.J., Britton, J.M., MacLean, M.E., Hancock, K.D., Fletcher, B.A., Hiebert, S.D., 1990: Geology and Mineral Deposits of the Snippaker Area (NTS 104B/6E, 7W, 10W, 11E); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1990-16.

Alldrick, D.J., Britton, J.M., Webster, I.C.L. and Russell, C.W.P., 1989: Geology and Mineral Deposits of the Unuk Area (NTS 104B/7E, 8W, 9W, 10E); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1989-10.

Alldrick, D.J., Mawani, Z.M.S., Mortensen, J.K. and Childe, F., 1996: Mineral deposit studies in the Stewart District (NTS 103O/P and 104A/B); *in*: Exploration in British Columbia 1995, Part B - Geological descriptions of properties. British Columbia Ministry of Energy, Mines and Petroleum Resources, pp. 89-109.

Alldrick, D.J., Nelson, J.L and Barresi, T., 2005a: Geology of the Volcano Creek - More Creek Area, British Columbia (NTS 104B/9,10,15,16; 104G/1,2); British Columbia Geological Survey Open File 2005-5, 1 map.

Alldrick, D.J., Nelson, J.L and Barresi, T., 2005b: Geology and Mineral Occurrences of the Upper Iskut River Area: Tracking the Eskay Rift through Northern British Columbia (Telegraph Creek NTS 104G/1, 2; Iskut River NTS 104B/9, 10, 15, 16); *in*: Geological Fieldwork 2004, British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Paper 2005-1, pp. 1-30.

Alldrick, D.J., Nelson, J.L., Barresi, T., Stewart, M.L., and Simpson, K.A., 2006. Geology of the upper Iskut River area, British Columbia British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Open File 2006-2, scale 1:100,000.

Anderson, R.G., 1989: A stratigraphic, plutonic and structural framework for the Iskut River Map Area, Northwestern British Columbia; *in*: Current Research, Part E. Geological Survey of Canada Paper 89-1E, pp. 145-154.

Anderson, RG., 1993: A Mesozoic stratigraphic and plutonic framework for northwestern Stikinia (Iskut River area), northwestern British Columbia, Canada; *in*: Dunne, G. and McDougall, K. (eds.): Mesozoic palæogeography of the Western United States-II; Society of Economic Palæontologists and Mineralogists, Pacific Section, vol. 71, pp. 477-494.

Anderson, R.G., Simpson, K., Alldrick, D., Nelson, J. and Stewart, M., 2003: Evolving ideas on the Jurassic tectonic history of northwestern Stikinia, Canadian Cordillera; Geological Society of America Abstracts with Programs, Vol. 35, No. 6, September 2003, p.89.

Atkinson, J.R., Metcalfe, P. and Moore, M.J. 1991: Summary report, 1990 program: Skyline Gold Corporation / Placer Dome Inc. joint venture: Bronson Project; Skyline Gold Corporation Report (unpubl. report), 72p.

Biczok, J., 2021: Report on the 2021 summer exploration program on the Red Line claims; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report, 54 p. (in review)

British Columbia Geological Survey Branch, 1991-2022: MINFILE; http://minfile.gov.bc.ca

Cui, Y., Miller, D., Schiarizza, P., and Diakow, L.J., 2017. British Columbia digital geology. British Columbia Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey Open File 2017-8, 9p. (Data version 2019-12-19).

Evenchick, C.A., 1991a: Geometry, evolution, and tectonic framework of the Skeena Fold Belt, north-central British Columbia; Tectonics, v. 10, no. 3, pp. 527-546.

Evenchick, C.A., 1991b: Structural relationships of the Skeena Fold Belt west of the Bowser Basin, northwest British Columbia; Canadian Journal of Earth Sciences, v. 28, p. 973-983.

Fields, M., 1992: Assessment Report for the For property; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report 22102, 54 p., 7 maps.

Fletcher, B.A. and Hiebert, S.D., 1990: Geology of the Johnny Mountain Area (NTS 104B/11); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1990-19.

Gagnon, J.-F., Barresi, T., Waldron, J.W.F., Nelson, J.L., Poulton, T.P., Cordey, F. and Colpron, M., 2012: Stratigraphy of the upper Hazelton Group and the Jurassic evolution of the Stikine terrane, British Columbia; Canadian Journal of Earth Sciences, 49, pp.1027-1052.

Gal, L. and Walker, R., 1990: Geological and geochemical summary report on the For claim group; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report 20598, 98 p., 5 maps.

Galloway, J.D., 1932: Lode-gold deposits of British Columbia; British Columbia Department of Mines Bulletin No. 1932-1, 144p.

Geological Survey of Canada, 1957: Stikine River area, British Columbia; Geological Survey of Canada Map 9-1957, 2 sheets.

George, S.W.M., Nelson, J.L., Alberts, D., Greig, C.J. and Gehrels, G.E., 2021: Triassic-Jurassic Accretionary History and Tectonic Origin of Stikinia From U-Pb Geochronology and Lu-Hf Isotope Analysis, British Columbia; Tectonics 40, 27p.

Government of Canada, 2022: Canadian climate normals, 1971-2000 station data, Bob Quinn AGS; https://climate.weather.gc.ca/climate_normals/results_e.html?searchType=stnProv&lstProvince=BC&txtCentralLatMin=0&txtCentralLatSec=0&txtCentralLongMin=0&txtCentralLongSec=0&stnID=1484&dispBack=0

Greig, C.J., 2014: Latest Triassic-Earliest Jurassic contractional deformation, uplift and erosion in Stikinia, NW B.C.; Abstracts, 2014 Geological Society of America Annual Meeting, Vancouver, British Columbia (19–22 October 2014)

Grove, E.W., 1971: Geology and Mineral Deposits of the Stewart Area, British Columbia; British Columbia Ministry of Energy, Mines and Petroleum Resources Bulletin 63, 219 p., 40 illus., 19 Plates, 3 maps, 1 section.

Grove, E.W., 1986: Geology and mineral deposits of the Unuk River-Salmon River-Anyox area; British Columbia Ministry of Energy, Mines and Petroleum Resources Bulletin 63, 434 p.

Gunning, M.H., 1993: Character of Upper Paleozoic Strata and Plutons, Lower Forrest Kerr Creek area, northwestern British Columbia; *in* Current Research, Part A, Cordillera and Pacific Margin / Recherches En Cours, Partie A, Cordillere Et Marge Du Pacifique; Geological Survey of Canada, Paper 93-1A, pp. 27-36.

Henderson, S., 2018: Galore Creek: A Complex High Grade Cu-Au Alkalic Porphyry System; Galore Creek Mining Corporation presentation; Prospectors and Developers Association of Canada, Annual Meeting, Toronto, Ontario.

Hunter, R.C. and van Straaten, B.I., 2020: Preliminary stratigraphy and geochronology of the Hazelton Group, Kitsault River area, Stikine terrane, northwest British Columbia; *in*: Geological Fieldwork 2019, British Columbia Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey Paper 2020-01, pp. 101-118.

Lefebure, D.V. and Gunning, M.H., 1989: Geology, Lithogeochemistry and Mineral Occurrences of the Bronson Creek Area (NTS 104B); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1989-28.

Logan, J.M., Koyanagi, V.M. and Drobe, J.R., 1990: Geology, Geochemistry and Mineral Occurrences of the Forrest Kerr - Iskut River Area, Northwestern B.C. (NTS 104B/10, 15); British Columbia Ministry of Energy, Mines and Petroleum Resources Open File 1990-02

Logan, J.M., Drobe, J.R., Koyanagi, V.M. and Elsby, D.C., 1997: Geology of the Forrest Kerr-Mess Creek area, northwestern British Columbia (105B/10, 15 & 105G/2 and 7W): Ministry of Employment and Investment Geoscience Map 1997-3, scale 1:100 000.

McConnell, R.G., 1913: Portions of Portland Canal and Skeena Mining Divisions, Skeena District, British Columbia; Geological Survey of Canada, Memoir no. 32, 101p.

Macdonald, A.J., Lewis, P.D., Ettlinger, A.D., Bartsch, R.D., Miller, B.D. and Logan, J.M., 1993: Basaltic rocks of the Middle Jurassic Salmon River Formation, northwestern British Columbia (104A, B, G); *in*: Geological Fieldwork 1992; British Columbia Ministry Of Energy, Mines and Petroleum Resources Paper 1993-1, pp.307-314.

Macdonald, A.J., Lewis, P.D., Thompson, J.F.H., Nadaraju, G., Bartsch, R., Bridge, D.J., Rhys, D.A., Roth, T., Kaip, A., Godwin, C.I., and Sinclair, A.J., 1996: Metallogeny of an Early to Middle Jurassic arc, Iskut River area, northwestern British Columbia; Economic Geology, 91, pp.1098-1114.

Marsden, H., and Thorkelson, D.J., 1992. Geology of the Hazelton volcanic belt in British Columbia: Implications for the Early to Middle Jurassic evolution of Stikinia. Tectonics, 11, pp. 1266-1287.

Metcalfe, P., 2011: Geological map of part of the Georgie River property, Skeena Mining Division, B.C.; British Columbia Ministry of Energy and Mines Assessment Report 32000, 139p, 2 maps.

Metcalfe. P., 2014: The Georgia River deposit, Stewart, B.C.: a template for Early Jurassic mineralization in the Golden Triangle? Kamloops Exploration Group Annual Conference, Abstracts, April 2014

Metcalfe. P., 2021: Reconnaissance geochemical sampling and reassessment of historic diamond drilling at the Georgia River Mine, Stewart, British Columbia; British Columbia Ministry of Energy and Mines Assessment Report 39326, 114p.

Metcalfe. P. and Nelson, J., 2014: Changes in stress field during Texas Creek mineralization in northwestern British Columbia: an exploration tool and possible record of microcontinent collision. Geological Society of America, Annual Meeting, Vancouver, B.C., Abstracts.

Miller, E.A., Kennedy, L.A. and van Straaten, B.I., 2020: Geology of the Kinskuch Lake area and Big Bulk porphyry prospect, northwestern British Columbia: Syndepositional faulting and basin formation during the Rhaetian (latest Triassic) transition from the Stuhini to the Hazelton Group; *in*: Geological Fieldwork 2019, British Columbia Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey Paper 2020-01, pp. 77-99.

Nelson, J.L., 2014: Reactivated basement structures controlled emplacement of Triassic-Jurassic Stikine Cu-Au porphyry deposits; Abstracts, 2014 Geological Society of America Annual Meeting, Vancouver, British Columbia (19–22 October 2014)

Nelson, J.L., 2017: Composite pericratonic basement of west-central Stikinia and its influence on Jurassic magma conduits: Examples from the Terrace-Ecstall and Anyox areas; *in*: Geological Fieldwork 2016, British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Paper 2017-1, pp. 61-82.

Nelson, J.L., 2019: Iskut region geological compilation: Supporting data and working files; Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey GeoFile 2019-02, 2 p.

Nelson, J., Waldron, J., van Straaten, B., Zagorevski, A. and Rees, C., 2018: Revised stratigraphy of the Hazelton Group in the Iskut River region, northwestern British Columbia; *in*: Geological Fieldwork 2017, British Columbia Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey Paper 2018-1, pp.15-38.

Nicholson, J.A., 1990a: Geological and Geochemical Report on the For Claim; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report 19634, 44p., 1 map.

Nicholson, J.A., 1990b: Geological and Geochemical Report on the Ning Claim Group; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report 19635, 38p., 1 map.

Rescan Environmental Services, 2013: Brucejack Gold Mine Project: 2012 Meteorology Baseline Report. Rescan Environmental Services Ltd. Report for Pretium Resources Inc., 174p.

Rhys, D.A., 1993: Geology of the Snip Mine, and its relationship to the magmatic and deformational history of the Johnny Mountain area, northwestern British Columbia; University of British Columbia M.Sc. thesis, 289p.

Sherlock, R.L., Roth, T., Spooner, E.T.C. and Bray, C.J., 1999: Origin of the Eskay Creek precious metal-rich volcanogenic massive sulfide deposit; fluid inclusion and stable isotope evidence; Economic Geology 94, pp.803–824.

Souther, J.G., 1971: Telegraph Creek map-area, British Columbia; Geological Survey of Canada Paper 71-44, 40p.

Walker, R.T. and Moore, M.P., 1990: Geological and geochemical summary report on Bing Structure and Ning groups; British Columbia Ministry of Energy, Mines and Petroleum Resources Assessment Report 20885, 121 p., 5 maps.

Wang, T., Hamann, A., Spittlehouse, D.L. and Aitken S.N., 2006: Development of scale-free climate data for western Canada for use in resource management; International Journal of Climatology 26, pp.383–397

CERTIFICATE OF AUTHOR

I, **Paul Metcalfe**, a consulting geologist, do hereby certify that:

1. I am a resident of British Columbia and the principal of Palatine Geological Ltd. (Permit to Practice 1003060), with a business address at P.O. Box 289, Gabriola, B.C. V0R 1X0
2. I am the author of the document entitled "Technical report on the Red Line property", effective on 4th November, 2022;
3. I am a graduate of the University of Durham (B.Sc. Hons. *Dunelm.* 1977), a graduate of the University of Manitoba (M.Sc. 1981) and a graduate of the University of Alberta (Ph.D. 1987), with a specialty in volcanology;
4. I am a member, in good standing, of Engineers and Geoscientists BC (registration number 23944) a Fellow of the Geological Society of London and a member of the Society of Economic Geologists;
5. I have worked as a geologist for more than 45 years since my graduation from the University of Durham, including employment as a postdoctoral research fellow by the Mineral Deposits Research Unit (Iskut Project) at the University of British Columbia and as a postdoctoral research fellow (volcanology) at the Geological Survey of Canada; in particular, I have worked in the Golden Triangle area of northwestern Stikinia on numerous properties since 1988;
6. I have read the definition of qualified person set forth in National Instrument 43-101 (NI 43-101) and certify that by reason of education, affiliation with a professional association as defined in NI 43-101, and past relevant work experience, I fulfil the requirements to be a qualified person for the purposes of NI 43-101, and that this Technical Report has been prepared in compliance with the Instrument;
7. I inspected the Red Line property for the first and only time between 9th July and 19th July 2022;
8. I am the sole author and solely responsible for the contents of "Technical report on the Red Line property";
9. I am independent of the Issuer, Copperhead Resources Inc., Romios Gold Resources Inc. and any subsidiaries, applying all the tests in section 1.5 of NI 43-101 and:
10. As of the date of this Certificate, to the best of my knowledge, information, and belief, this Technical Report contains all scientific and technical information required to be disclosed to make the Technical Report accurate and not misleading.

DATED at Gabriola Island, British Columbia, Canada this 4th day of November, 2022

"Paul Metcalfe"

Dr. Paul Metcalfe P.Geo. FGS