Report to:

FAR RESOURCES

NI 43-101 TECHNICAL REPORT ON THE ZORO LITHIUM PROJECT, SNOW LAKE, MANITOBA

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July 6, 2018

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1.0 SUMMARY

This Technical Report was prepared to support ongoing exploration and development by the current property owner Far Resources Ltd. Its main objective is to provide an updated review of historic and recently discovered lithium-cesium-tantalum-bearing pegmatite dykes on the property based on exploration to date and to utilize the related geoscientific database to demonstrate sufficient technical merit to continue the assessment of known pegmatite dykes and to explore for repetitions of this style of mineralization.

The Zoro property consists of 16 claims with a total of 3603 hectares. The Zoro1 claim is wholly owned by Far Resources and 15 claims have been optioned from the property vendor Strider Resources Limited. To earn a 100% interest in the property free and clear of all liens, charges, royalties (save and except for the NSR contemplated herein) and claims of others Far Resources must satisfy the following conditions:

1. an undivided one-hundred percent (100%) interest in the Optioned Interest, subject to the NSR (the “**First Option**”); and

2. an undivided fifty percent (50%) interest in the NSR, being one-half of the NSR or a 1% Net Smelter Return, in addition to the undivided one-hundred percent (100%) interest in the Optioned Interest that has been acquired under the First Option (the “**Second Option**”).

The Optionee may exercise:

(a) the First Option by making the following cash payments and common share issuances to the Optionor:

  (1) Shares in the capital of the Optionee Valued at $25,000 and $25,000 cash within two Business Days following the Effective Date; and

  (2) Shares in the capital of the Optionee Valued at $50,000 and an additional $50,000 cash on or before the twelfth month anniversary of the Effective Date; the Optionee must spend $50,000 on Exploration Expenses by the end of the first 12 months; and

  (3) Shares in the capital of the Optionee Valued at $50,000 and an additional $50,000 cash on or before the twenty-fourth month anniversary of the Effective Date; the Optionee must have spent an accumulated total of $100,000 on Exploration Expenses by the end of the first 24 months; and

(4) Shares in the capital of the Optionee Valued at $50,000 and an additional $50,000 cash on or before the thirty-sixth month anniversary of the Effective Date; the Optionee must have spent an accumulated total of $150,000 on Exploration Expenses by the end of the first 36 months; and

(5) Shares in the capital of the Optionee Valued at $75,000 and an additional $75,000 cash on or before the forty- eighth month anniversary of the Effective Date; the Optionee must have spent an accumulated $200,000 on Exploration Expenses by the end of the first 48 months and an accumulated total of $500,000 on Exploration Expenses by the end of the first 84 months.

(b) provided it has exercised the First Option, the Second Option, by making a cash payment to the Optionor of $1,000,000, together with all accrued but unpaid NSR at the time, prior to the commencement of Commercial Production (the "**Second Option Payment**");

Spodumene mineralization on the property occurs within laterally and vertically extensive pegmatite dykes hosted within andesitic volcanic rocks and intermediate to felsic sedimentary rocks. Abundant overburden cover characterizes much of the property. Geophysically the pegmatite dykes do not have a recognizable geophysical signature.

Based on current exploration results a base case inferred resource for Dyke 1 on the property has been determined. The reporting cut-off is 0.3 percent Li_2O. Dyke 1 contains 1,074,567 tonnes at 0.91% Li_2O, 182 ppm Be, 198 ppm Cs, 51 ppm Ga, 1212 Rb, and 43 ppm Ta. The full sensitivity analysis of tonnage and grade is summarized below in Table 1, based on various Li_2O cut-off percentages, is summarized below in Table 1.1 based on Li_2O cut-off percentages. Recommendations for ongoing exploration are included.

Table 1.1: Base Case Inferred Resource estimate, using 0.3% Li_2O, and sensitivities to tonnage and grade based on increased cutoffs.

Li_2O (%) Cut-off	Tonnes	Li_2O (%)	Be (ppm)	Cs (ppm)	Ga (ppm)	Rb (ppm)	Ta (ppm)
0.3	1,074,567	0.91	182	198	51	1212	43
0.4	946,402	0.99	180	201	51	1203	43
0.5	881,815	1.03	179	203	51	1197	43
0.6	780,350	1.09	180	207	52	1196	42
0.7	721,660	1.13	179	208	52	1190	42
0.8	629,578	1.18	181	210	52	1174	42

0.9	515,652	1.26	183	211	53	1152	43
1.0	419,961	1.33	188	212	54	1135	43

The Zoro property hosts numerous historic pegmatite dykes and on the basis of its exploration program has defined additional spodumene-bearing dykes on the property. Utilizing an innovative approach to exploring beneath overburden cover based on Mobile Metal Ion Technology Far Resources has discovered an additional previously unrecognized spodumene-bearing pegmatite dyke.

It is concluded that based on results to date an ongoing program based on a combination of MMI soil geochemistry and diamond drilling have the potential to discover additional lithium-bearing pegmatite dykes on the property. A recommended program of helicopter-assisted MMI surveys and diamond drilling costing $885,500.00 has been proposed for the property.

1.1 GEOLOGICAL SETTING AND MINERALIZATION

The Zoro Lithium Project is located at the east end of the Flin Flon-Snow Lake greenstone belt. The Paleoproterozoic Flin Flon-Snow Lake Belt is approximately 200 km in strike length and has an exposed width of up to 70 km. The Belt is overlain to the south by Ordovician Red River Formation sandstone, limestone, and dolomite of the Western Canada Sedimentary Basin, and is bordered to the north by high-grade paragneiss and granitoid rocks of the Kisseynew Domain. The Flin Flon Belt is interpreted to be an accreted assemblage of oceanic to continental margin arc terrane, interspersed with oceanic basins representing back-arc, fore-arc, and oceanic settings.

At the project scale the general and detailed geology for the Zoro Lithium Project is underlain by Ocean Floor volcanic rocks of the Roberts Lake allochthon and lesser amounts of Missi Group sedimentary rocks. The Ocean Floor rocks comprise mafic volcanic rocks and related intrusions and the Missi Group consists of sandstone, siltstone, mudstone and quartzo-feldspathic gneiss and migmatite.

The major mineralizing events recognized in the Flin Flon belt took place during the three main stages of crustal development: pre-accretion, post-accretion, and continent- continent collision. The pre-accretionary stage is represented by syngenetic base metal and Au deposits. The syn- to post-accretionary stage is characterized by several examples of intrusion-hosted base and precious metal deposits, and the continental collision stage by the development of orogenic Au deposits and lithium-cesium-tantalum-enriched pegmatites.

The spodumene-bearing pegmatite dykes on the property strike northwest with steep dips and crosscut the regional foliation at a low angle. The dykes tend to be concentric in internal structure and the grain size of the constituent minerals (potassium feldspar, quartz, spodumene and black tourmaline) coarsens towards the center of the dykes. This pattern may be locally interrupted by patches of saccharoidal albite, large muscovite aggregates and coarse albite stringers with garnet and beryl. Spodumene is concentrated in the core of the majority of the dykes. Some of the dykes have been split into sub-parallel veins by post-emplacement tectonic activity.

The pegmatite dykes and the host Ocean Floor mafic volcanic rocks are transected by northwest-trending structures. The general area is also crosscut by a series of northeast and near-east-trending structures including the major Berry Creek fault that extends along Crowduck Bay, to the west of the project area.

1.2 MINERAL RESOURCE ESTIMATE

Historic lithium tonnage estimates vary. An unsubstantiated visual estimate in September 1956 suggested up to 9-11 million tonnes (10-12 million tons) of Li_2O occur within the claim group. In mid-March Dyke 1was estimated to contain 1.8 million tonnes (2 million tons) grading 1.4% Li_2O to a depth of 305 m (1000 ft.) in the main dyke (“Dyke 1”; Northern Miner, October 25, 1956; Mulligan, 1965, p. 81). A reserve estimate of 1 815 000 tonnes grading 1.4% Li_2O was reported by Bannatyne (1985). In 1957, the estimate was revised to 1.72 million tonnes averaging 1.3% Li_2O or 2.72 million tonnes (3.0 million tons) at 1.0% Li_2O in Dyke 1 (Mulligan, 1957a, 1957b). By March 1958, 12 different tonnage estimates had been made (Northern Miner, March 13, 1958). The description of mineral resources cited above is presented as a historical resource estimate and uses historical terminology for these estimates. These citations are given to provide an historical frame of reference. Although the resource estimations quoted in the text are believed to be reliable, they were calculated prior to the implementation of National Instrument 43-101. The author has not carried out work to classify these historical estimates under current mineral resource or mineral reserve terminology. The historical estimates are not meant to be interpreted as current estimates as defined in section 1.2 and 1.3 of NI43-101 and should not be relied upon. These estimates were calculated without the benefit of a quality assurance or quality control program, a lack of control over drill collar locations and the ultimate end point of the drill hole. Accordingly, the issuer is not treating the historical estimate as a current mineral resource or mineral reserve.

A current inferred mineral resource estimate is calculated to be 1,074,567 tonnes grading 0.91% Li_2O at a cut-off grade of 0.3%.

Inferred Mineral Resources are not Mineral Reserves. Mineral resources which are not mineral reserves do not have demonstrated economic viability. There has been insufficient exploration to define the inferred resources as an indicated or measured mineral resource, however, it is reasonably expected that the majority of the Inferred Mineral Resources could be upgraded to Indicated Mineral Resources with continued exploration. There is no guarantee that any part of the mineral resources discussed herein will be converted into a mineral reserve in the future.

1.3 PROJECT INFRASTRUCTURE

The Zoro Lithium Project is located approximately 20 km east of the Town of Snow Lake. Nearby infrastructure includes a power line servicing the town of Snow Lake approximately 5 km south of the property, the Snow Lake airport and an all-weather gravel road 11 km west of the property, and a rail link located at Wekusko siding, 20 km to the south of Herb Lake Landing which is 30 km south of the property. The nearest road link is a seasonal road on the east side of Wekusko Lake that accesses the village of Herb Lake Landing and Provincial Highway 392 to the south.

For purposes of exploration access to the property can be achieved using provincial highway 39 and driving north to Bartlett's Landing where a boat can be launched from the shores of Wekusko Lake. It is approximately a 20-km boat ride to the Property. From this point the property can be reached using drill roads and ATV trails. Gogal Air Services, a helicopter and float plane charter company, operates from Snow Lake.

1.4 ENVIRONMENT

The Zoro Lithium Project is an early stage exploration project despite a long history of exploration activity at the site since 1956. The status of the project precludes the undertaking of an Environment Impact Assessment (fauna, flora and social) for both federal (Canadian Environmental Assessment Agency) and provincial authorities. Regular communications are ongoing with the nearby town of Snow Lake. Work permits are routinely acquired from the Manitoba department of Sustainable Development, Snow Lake District in 2-4 weeks.

2.0 INTRODUCTION AND TERMS OF REFERENCE

2.1 INTRODUCTION

The Zoro Lithium Project is located near the east shore of Wekusko Lake in west-central Manitoba, approximately 25 km east of the mining town of Snow Lake, 249 km southeast of Thompson and 571 km north-northeast of Winnipeg.

The property has a long history of intermittent exploration commencing in 1956 and after a lengthy period of no activity exploration was re-started in 2009. This technical report is an update of 43-101 technical reports produced in 2009 (Fedikow, 2009) and in 2012 (Fedikow, 2012). These reports utilized evaluations of historic data from the Manitoba Cancelled Assessment files, Historic Corporate Files and data summarized from non-confidential assessment reports within the files of Manitoba Mining Recorder’s office (Winnipeg) in addition to results for exploration undertaken by Force Energy between 2009 and 2012. The objectives in each of the preceding technical reports was to provide an understanding of the geological setting of spodumene-bearing lithium-cesium-tantalum-bearing pegmatite dykes thereby providing guidelines for further exploration. The current report builds on these reports with recent exploration results by Far Resources Ltd. between 2016 and 2017 including diamond drilling and geological and geochemical surveys. The business of Far Resources is the acquisition, exploration and development of lithium-bearing mineral properties and to this end Far has been actively exploring for lithium in the Snow Lake area of Manitoba.

Recently, the expansion of the initial Zoro 1 claim of 52 hectares optioned from Dalton Dupasquier of Top Notch Marketing Inc. in 2016 was expanded to 3603 hectares by the acquisition through option in two separate agreements with Strider Resources Limited.

The author has visited the property on numerous occasions. The first was to undertake a mineral deposit description for the Manitoba Geological Survey as part of a program to document mineral deposits and occurrences in the Province. This field visit was made in 1986 and results published in Fedikow *et al.* (1986) and expanded upon in Fedikow *et al.* (1993). A subsequent visit was made in August of 2009 on behalf of optionee Force Energy (Colorado Springs, U.S.A.) to assess the availability of outcrop for geological mapping, to examine the property area for exposures of the pegmatite dykes, to review the immediate area for historic and new diamond drill hole collars and to assess the general area for considerations relating to further exploration and possible production decisions. The most recent property visit was undertaken on behalf of Far Resources Ltd. on May 25, 2018 to review the area for ongoing diamond drill programs and general exploration activities.

2.2 TERMS OF REFERENCE

This report entitled "Zoro Lithium Project" is prepared at the request of Far Resources Ltd. Their business address is 201-2691 Viscount Way. Richmond, British Columbia V6V 2R5, Canada. The current report is an update to the original 2009 and 2012 NI 43-101 technical report. This update report includes a description of recent survey results on the property follows the format defined by NI43-101F1. Sources of information that have been utilized to build this report include:

- The Manitoba Mining Recorders Office (Winnipeg) has provided the recorded description including the current status of the Zoro claims as well as recorded adjacent claims. Copies of archived assessment files which include historic work reports, maps, data and diamond drill logs of exploration work completed on and in the general vicinity of the property;
- Public domain geological literature from the Manitoba Geological Survey and the Geological Survey of Canada which describe the geological setting of the area of the Wekusko Gold Project property;
- Interviews with mineral exploration companies actively exploring in the general area of the Far Resources Zoro Lithium Project;
- Personal and telephone/e-mail discussions with exploration personnel that have done work on the property in the past.

This report was constructed by Mark Fedikow Ph.D. P.Eng. P.Geo. C.P.G., consulting geologist and geochemist. He has previously undertaken geological and geochemical surveys on the Zoro 1 property as a Manitoba Geological Survey geologist, as a consultant for both Force Energy Ltd. and Far Resources Ltd. Scott Zelligan, an independent resource geologist prepared sections 13 and 15.. An inferred resource of 1,074,567 tonnes grading 0.91%Li_2O at a cut-off of 0.3% Li_2O has been calculated for Dyke 1 on the property.

1,074,567	0.91

3.0 RELIANCE ON OTHER EXPERTS

Mark Fedikow Ph.D. P.Eng. P.Geo. C.P.G. (Mount Morgan Resources Ltd.) is an independent mineral exploration consultant and holds no interest in the Zoro Lithium Project. The author will be paid a fee for the preparation of this report according to normal consulting practice.

4.0 PROPERTY DESCRIPTION AND LOCATION

4.1 PROPERTY LOCATION

The Zoro Lithium Project is located near the east shore of Wekusko Lake (Table 4.1, Figure 4.1) in west-central Manitoba, approximately 25 km east of the mining community of Snow Lake, 249 km southeast of Thompson and 571 km north-northeast of Winnipeg. Provincial Road 393 occurs 23 km to the northwest. The pegmatite dykes are located northwest of the northwest corner of Johnson Lake a small lake east of the east shore of Wekusko Lake. The small historic gold mining community of Herb Lake is located about 10 km southwest of the property.

The property is located within NTS map sheet 63J/13SE (latitude: 54^{o}51.27' and longitude: 99^{o}38.46'; Township 68N; Range 15WPM). The project consists of 16 claims totaling 3,005 Ha. The ZORO 1 claim is 100% owned by Far Resources and the remaining 15 claims are owned by Strider Resources Ltd. In 2016, Far Resources and Strider Resources entered an option agreement whereby Far Resources could earn 100% interest of JAKE 3558 and an additional 350m strip of BERT 6304 and BERT 797 subject to a 2% NSR to Strider Resources. The agreement was expanded in 2017 to include JAKE 9, JAKE 1054, JAKE 2655, JAKE 3557, JAKE 54199, JAKE 10, JAKE 2412, JAKE 2413, JAKE 54745 and CRO 55734 and subsequent to that agreement, additionally BAZ 12131 and BAZ 12133 were added in the same year. There are no known additional royalties or back-in-rights to the property.No negotiations are required for access to the property and surface rights reside with the Manitoba government. Far Resources will undertake an integrated exploration program of Mobile Metal Ion soil geochemical surveys coupled with diamond drilling and a combination of cash payments and share issuances Permits for to earn its 100% interest

A work permit must be granted by Sustainable Development, a department of the Government of Manitoba, for any exploration work which may cause a disturbance to the local environment. Far Resources has adhered to the legislations put forth by the Manitoba Mines Act for any work that has been conducted on the Zoro Property. To the authors knowledge, there is no known environmental liabilities or other significant factors or risks which may affect the access, title or the ability to perform work on the property. Work permits for exploration including diamond drilling are issued 2 weeks subsequent to application being made and prior to the actual work being undertaken.

Table 6.1: Summary of claims, Far Resources Zoro Lithium Project.

NAME	NUMBER	HOLDER	STAKED	RECORDED	EXPIRES	HECTARES
ZORO 1	P1993F	FAR RESOURCES LTD.	2/13/1994	3/14/1994	5/13/2067	52
JAKE	P3558F	STRIDER RESOURCES LTD.	6/7/1996	7/3/1996	9/1/2030	250
BERT	MB6304	STRIDER RESOURCES LTD.	2/26/2008	3/17/2008	5/16/2030	28
BERT	MB797	STRIDER RESOURCES LTD.	5/28/1999	6/16/1999	8/15/2030	27
JAKE 9	P3031F	STRIDER RESOURCES LTD.	3/20/1995	3/27/1995	5/26/2030	256
JAKE 1054	MB1054	STRIDER RESOURCES LTD.	4/27/2002	5/17/2002	7/16/2030	240
JAKE 2655	MB2655	STRIDER RESOURCES LTD.	4/28/2002	5/17/2002	7/16/2030	255
JAKE 3557	MB3557	STRIDER RESOURCES LTD.	6/6/1996	7/3/1996	9/1/2030	256
JAKE 54199	W53199	STRIDER RESOURCES LTD.	10/16/1996	11/8/1996	1/7/2030	131
JAKE 10	P3032F	STRIDER RESOURCES LTD.	1/30/1995	2/27/1995	4/28/2030	173
JAKE 2412	MB2412	STRIDER RESOURCES LTD.	4/29/2002	5/17/2002	7/16/2030	256
JAKE 2413	MB2413	STRIDER RESOURCES LTD.	4/30/2002	5/17/2002	7/16/2030	196
JAKE 54745	W54745	STRIDER RESOURCES LTD.	4/23/1997	5/9/1997	7/8/2030	245
CRO 5734	MB5734	STRIDER RESOURCES LTD.	1/20/2010	2/11/2010	4/12/2030	192
BAZ 12131	MB12131	STRIDER RESOURCES LTD.	12/20/2017	1/10/2018	3/10/2020	192
BAZ12133	MB12133	STRIDER RESOURCES LTD.	12/21/2017	1/10/2018	3/10/2020	256

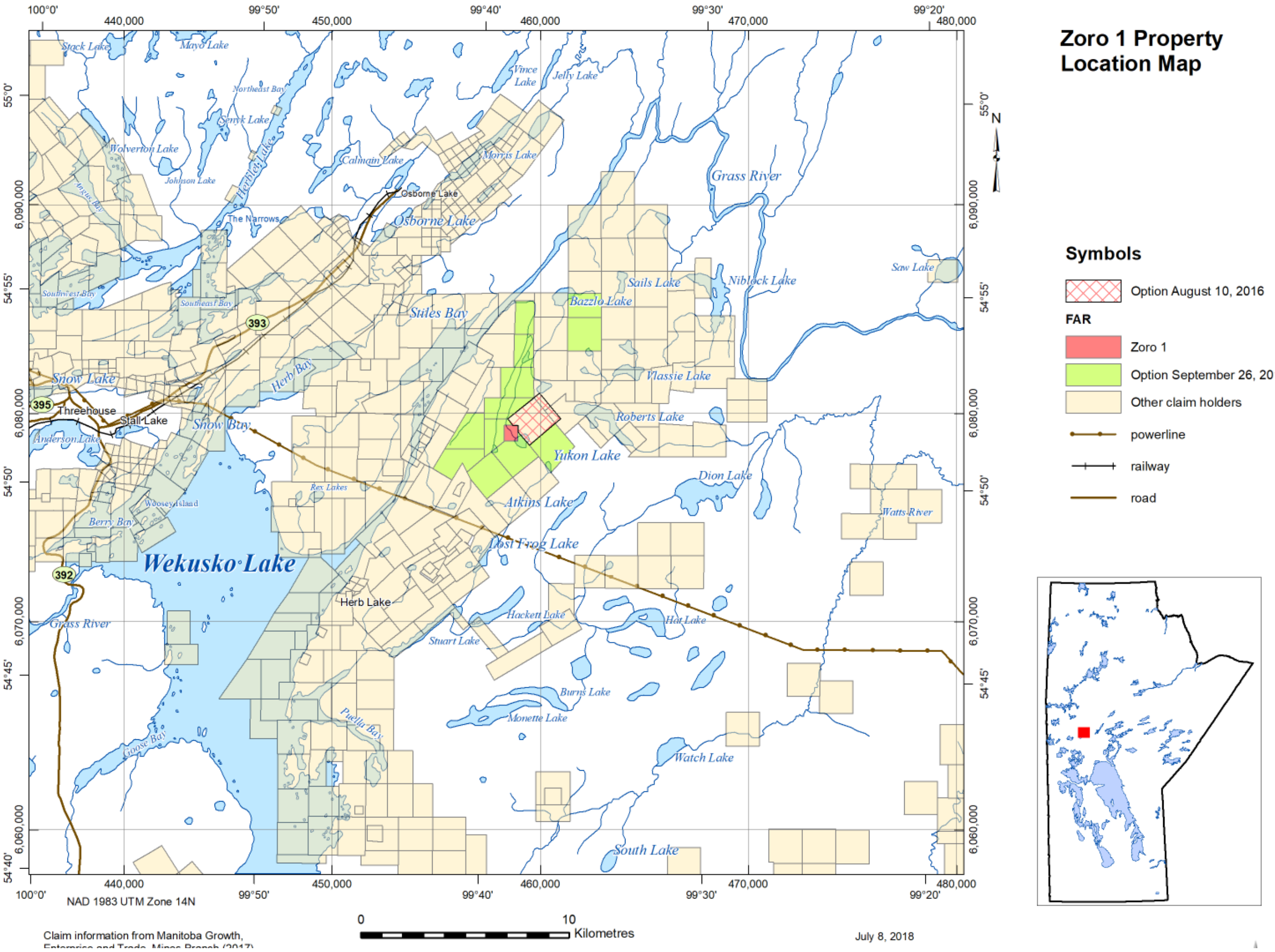


Figure 4.1: Location map for the Zoro Lithium Project, Snow Lake, Manitoba.

4.2 OTHER FACTORS AFFECTING ACCESS

Work and drill permits for exploration are received within two weeks from the local Snow Lake office of Sustainable Development. The Zoro Lithium Project occurs within the immediate area of the historic mining town of Snow Lake and as such very little negative community and social impact is evident. Exploration and mine development has been a part of the local community for 80 years. Currently the project is an early stage exploration project. Accordingly, environmental studies have not been undertaken.

5.0 ACCESSIBILITY, CLIMATE, LOCAL RESOURCES, INFRASTRUCTURE AND PHYSIOGRAPHY

5.1 ACCESS

The Property is located approximately 20 km east of the Town of Snow Lake. Access can be achieved using the main highway #39 and driving north on highway #392 to Bartlett's Landing where a boat can be launched from the shores of Wekusko Lake. It is approximately a 20-km boat ride to the Property. From this point the property can be reached using drill roads and ATV trails. The nearest rail link is at Wekusko siding, some 20 km to the south of Herb Lake Landing. The nearest road link is a seasonal road on the east side of Wekusko Lake that accesses the village of Herb Lake Landing and Provincial Highway 392 to the south. Gogal Air Services, a helicopter and float plane charter company, operates from Snow Lake and provides easy access to the property.

The Zoro Lithium Project is located within 10 km of paved Highway 392 that connects to the historic mining town of Snow Lake where an experienced mining labour force is available and to the local airport (Figure 15.1). Highway 392 also connects to paved provincial road 39 providing access to Flin Flon and Thompson. The property is 1.2 km north of the power line that services Snow Lake. Abundant drill roads and ATV trails crosscut the property.

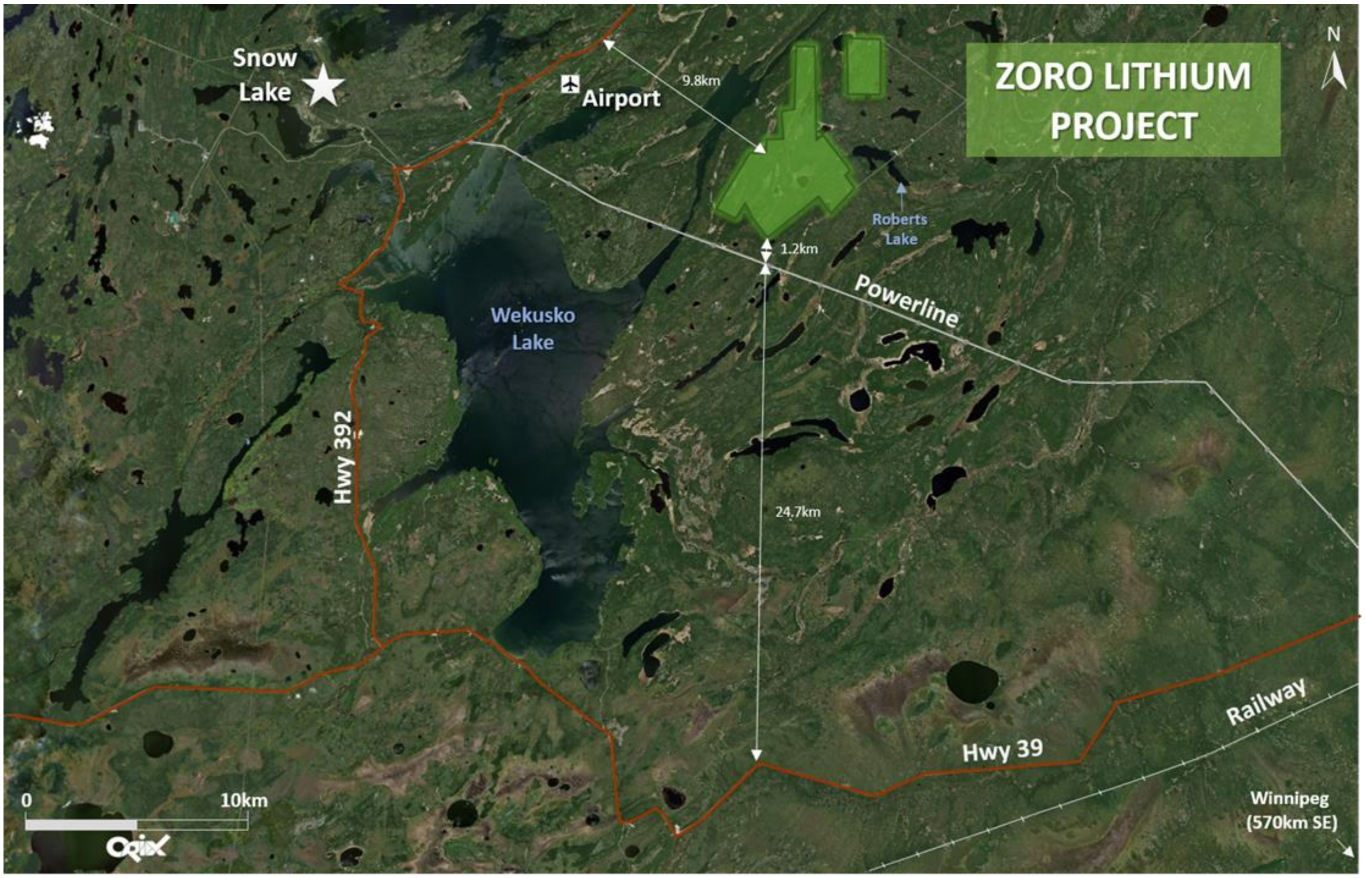


Figure 15.1: Available infrastructure in the area of the Zoro Lithium Project.

5.2 CLIMATE AND PHYSIOGRAPHY

Temperature averages for the community of Flin Flon, Manitoba are given in Table 5.2.1 for general information only. No detailed data were available from Environment Canada for Snow Lake. The Snow Lake climate is continental and characterized by cold winters (January mean temperature -21.1^{o}C and relatively warm summers (July mean temperature +18.3^{o}C). Total average precipitation is 477.9 mm per annum with 342.6 mm falling as rain and 137.2 mm falling as snow (for the years 1927-1990). Wind directions and velocities are well distributed, but predominate to the southeast and southwest, with strong components to the northwest, north and south. Electrical storms are common and forest fires can be problematic. Summer exploration work is best conducted between the months of May to September. Winter exploration work can be conducted from November to March.

The Property is located near the eastern shore of Crowduck Bay at the northeast end of Wekusko Lake. The shoreline of the bay is marked by approximately 10 metre slopes surrounded by heavily wooded flat areas, interspersed with low outcrop ridges. Lake elevation is around 260 m A.S.L. with the highest topographical ridge having an elevation of 290 m. Wekusko Lake is a large body of water 25 km long by 3 to 10 km wide. Crowduck Bay is located at the end of a 12 km long narrow channel that is the head of the Grass River.

The ecoregion is classified as sub-humid high boreal eco-climate. It forms part of the continuous coniferous boreal forest that extends from northwestern Ontario to Great Slave Lake in the southern Northwest Territories. The predominant vegetation comprises black spruce and jack pine with ericaceous shrubs and ground cover of mosses and lichens. Black spruce is the predominant species. Depending on the nature of local drainage and surficial sediments, trembling aspen, white birch, white spruce and, to a lesser extent, balsam fir predominate.

Bedrock exposures are covered with lichen. Poorly drained peat-filled depressions are marked by stunted black spruce with ericaceous shrubs and a ground cover of sphagnum moss. Permafrost is distributed throughout the ecoregion but is only widespread in organic surficial materials. In the western part of the ecoregion, uplands are blanketed by discontinuous sandy acidic tills, whereas extensive thin clay-rich lacustrine deposits and locally prominent, sandy fluvio-glacial uplands are common in the eastern section. Exposed bedrock occurs throughout the ecoregion and is locally prominent. Dystric and Eutric Brunisols are associated with sandy uplands, whereas Gray Luvisols occur on clayey lacustrine uplands and loamy to silty fluvioglacial deposits. On level and in depressions, Gleysolic soils are associated with clayey sediments, whereas Mesisols and Organic Cryosols are associated with shallow to deep peatlands.

A pulpwood and dimension lumber industry operates to a limited extent in the southern part of the ecoregion. Wildlife includes barren-ground caribou, moose, black bear, lynx, wolf, beaver, muskrat, snowshoe hare and red-backed vole. Bird species include raven, common loon, spruce grouse, bald eagle, grey jay, hawk owl and waterfowl, including ducks and geese. Trapping, hunting, fishing and tourism are the dominant uses of land in this region.

Table 1(5.2.1): Monthly climate averages for Flin Flon, including temperature highs, lows and precipitation values. Flin Flon climate data from the airport (54º41'N, 101º41'W at elevation 303.9 m) (source: Environment Canada website).

Temperature	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec	Year
Daily Average (°C)	-21.4	-16.7	-9.3	0.7	8.8	14.9	17.8	16.6	9.8	2.7	-8.4	-18.4	-0.2
Daily Maximum (°C)	-16.6	-11	-2.9	6.9	15	20.4	23.1	21.8	14.2	6.2	-5.1	-14	4.8
Daily Minimum (°C)	-26.2	-22.3	-15.8	-5.5	2.6	9.3	12.6	11.4	5.4	-0.8	-11.7	-22.6	-5.3
Extreme Maximum (°C)	9.5	10	15	27	32.5	35	35	33.9	30	24	17.5	8.3	
Date (yyyy/dd)	1993/30	1984/21	1993/24	1980/30	1986/28	1988/05	1989/21	1970/08	1991/01	1987/03	1978/02	1969/01	
Extreme Minimum (°C)	-44.5	-45.6	-41	-31	-13	-2	4.4	-1.5	-6.7	-16.5	-35	-44	
Date (yyyy/dd)	1996/30	1974/01	1995/04	1979/06	1990/02	1987/03	1969/05	1982/27	1974/28	1996/30	1985/28	1989/19	
Precipitation													
Rainfall (mm)	0.1	0.3	0.9	8.6	36.9	66.6	76.5	66.6	55.3	25.6	1.4	0.4	339.2
Snowfall (cm)	19.6	14.6	19.1	20	3.7	0	0	0	2	13	25.4	23.9	141.3
Precipitation (mm)	17.6	13.4	19	28.3	40.6	66.6	76.5	66.6	57.3	38.3	24.8	21.8	470.8
Average Snow Depth (cm)	33	39	32	8	0	0	0	0	0	1	11	25	
Extreme Daily Rainfall (mm)	2.2	3.8	12.2	25.4	62.6	54	78.2	53.8	55.6	24.9	9.1	10	
Date (yyyy/dd)	1984/02	1986/25	1987/20	1971/16	1985/04	1993/24	1981/23	1988/21	1984/07	1969/03	1974/08	1987/07	
Extreme Daily Snowfall (cm)	13	14.2	24	39.4	18	0.4	0	0	14.2	29.6	25.4	18.6	
Date (yyyy/dd)	1973/02	1987/12	1982/12	1973/20	1975/20	1987/03	1969/01	1969/01	1984/22	1991/27	1973/28	1981/06	

5.3 LOCAL RESOURCES AND INFRASTRUCTURE

The property is adequate in terms of area to permit the commercial exploitation of the lithium-bearing pegmatite dykes discovered to date. Exploration is assisted by helicopter support to mobilize drill and exploration crews in and out of the property. Helicopter flights from the Snow Lake airstrip to the property take approximately 15 minutes. Personnel and equipment can also be transported overland via Wekusko Lake and winter drill roads during freeze-up.

Nearby infrastructure (Figure 5.3.1) includes a power line servicing the town of Snow Lake approximately 5 km south of the property, the Snow Lake airport and an all-weather gravel road 11 km west of the property, and a rail link is located at Wekusko siding, 20 km to the south of

Herb Lake Landing which is 30 km south of the property. The nearest road link is a seasonal road on the east side of Wekusko Lake that accesses the village of Herb Lake Landing and Provincial Highway 392 to the south (Figure 5.3.1).

The closest community to the property is Snow Lake (about 20 km west), which has a population of approximately 800 people. The town of Flin Flon is about 150 km west of Snow Lake, has a population of 7,000 and is a provincial regional government centre. The town is a major support centre for smaller communities in northwestern Manitoba. There is a regional hospital in Flin Flon, a newly built shopping centre, and all necessary infrastructures to support the local town. Hudson Bay Mining and Smelting Ltd. (HBMS) operates the Lalor mine and a concentrator to the south of Snow Lake.

The Snow Lake area has the necessary infrastructure in place to provide the basis for supporting other mining operations in the district. The HBMS operations represent the largest employer in the area. The presence of HBMS also provides the basis for secondary support and supply local businesses. The municipal, regional and provincial government activities in the area also provide significant employment. There are tourist camps and lumber operations in the district. Consequently, there is a stable and experienced work force possessing the necessary skills in exploration and mining in the area.

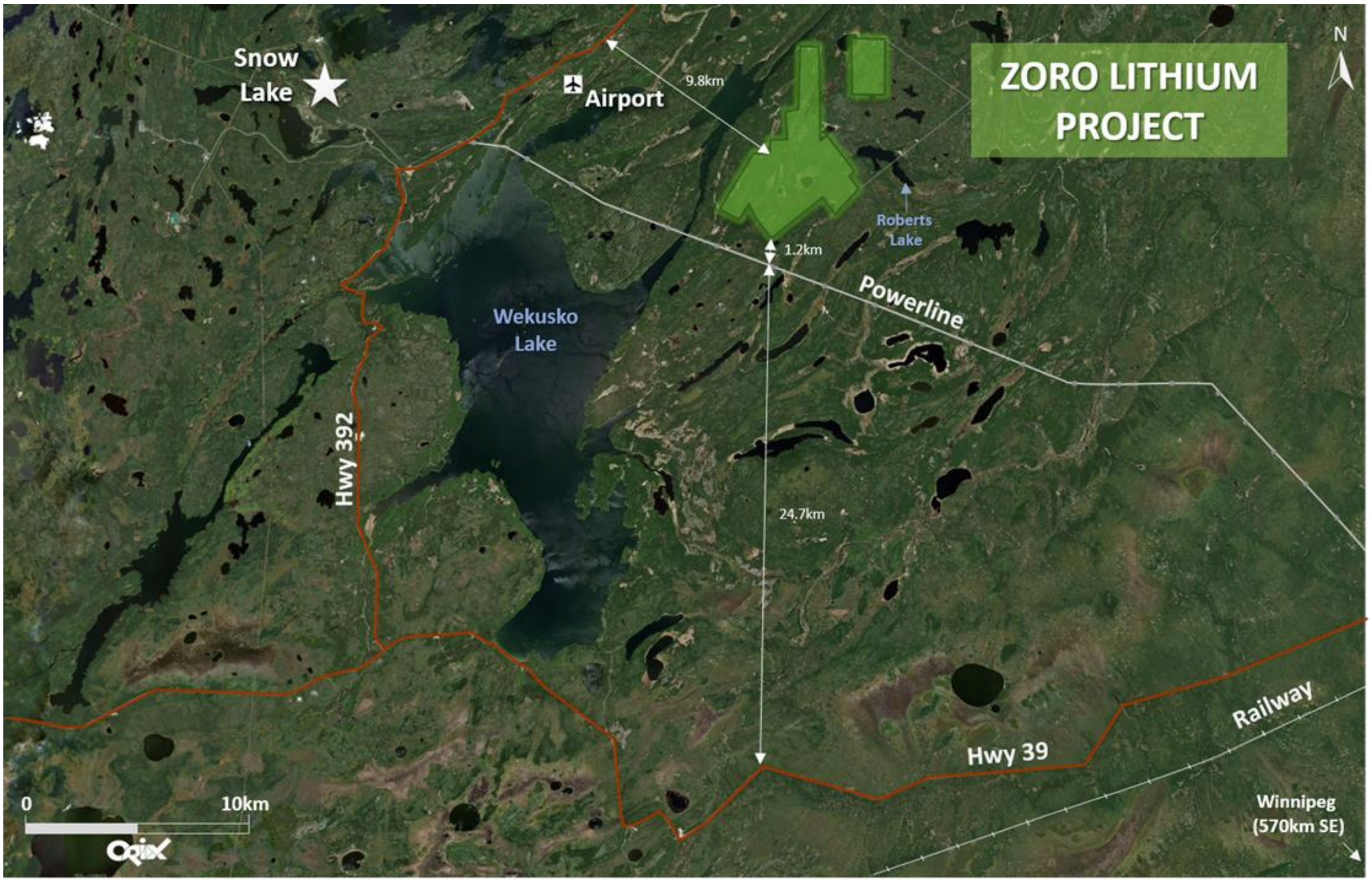


Figure 5.3.1: Infrastructure in the vicinity of the Zoro Lithium Project

6.0 HISTORY

The Zoro1 pegmatite dykes are located on the north side of a small lake between Roberts Lake and the south end of Crowduck Bay. Early in 1953, Cs No. 3-10, 12 (P 26973-80, 82), S.R. No. 1-6 (P 7877-82) and Linda 1 (P 26983) were staked by Mrs. Johanna Stoltz, Eric Stoltz, Carl Stoltz and Edwin Stoltz, and Key No. 1-4, 8-14 (P 27159-62, 27226-27, 27164-68) were staked by John Tikkanen, Hjalmar Peterson, and Loren Fredeen. These were cancelled the following year.

Lit Nos. 11-5 (P 31758-62) was staked by J.J. Johnson in 1954. In 1955 **Lit** Nos. 6-1l8 (P 35014-26) were added by J.A. Syme. All the **Lit** claims were assigned to Green Bay Uranium Limited in 1956 which changed its name to Green Bay Mining & Exploration Ltd.

Early in 1956, before drilling commenced, samples containing more than 2% Li_2O were reported (Northern Miner, January 12, 1956). A shipment of 136 kg (300 lbs.) of spodumene was sent to Ottawa for testing in 1956. This sample assayed 1.19% Li_2O, with minor NbO_5. Historic ore dressing tests concluded that good liberation and separation could not be effected (Mineral Dressing and Process Metallurgy Report in Green Bay Mining & Exploration Ltd., Corporation File).

Over 6096 m (20 000 ft.) of diamond drilling was done on **Lit** No. 1-4, with at least 3048 m (10 000 ft.) of this on the main dyke. Results of the drilling on dykes 1, 3, 5 and 7 were reported to be "promising". Assays of 2.42% to 7.28% Li_2O were reported from Dyke 5 (Green Bay; Corporation File). Dyke 5 was apparently 305 m long x 12 m wide (1000 x 40 ft.); Dyke No. 7, over 457 m x 24 m (1500 x 80 ft.). Several of the holes went deeper than 305 m (1000 ft.). Drilling on **Lit** 10, 16 and 17 amounted to 1950 m (6399 ft.). Gold was also found on the property, with a 3.3 kg (7.25 lb.) sample across 3.4 m (11 ft.) yielding 0.17 ounces per ton gold (Green Bay; Corporation File).

Historic lithium tonnage estimates vary. An unsubstantiated visual estimate in September 1956 suggested up to 9-11 million tonnes (10-12 million tons) of Li_2O occur on the entire group. In mid-March the main dyke was estimated to contain 1.8 million tonnes (2 million tons) grading 1.4% Li_2O to a depth of 305 m (1000 ft.) in the main dyke ("Dyke 1"; Northern Miner, October 25, 1956; Mulligan, 1965, p. 81). A reserve estimate of 1 815 000 tonnes grading 1.4% Li_2O was reported by Bannatyne (1985). In 1957, the estimate was revised to 1.72 million tonnes averaging 1.3% Li_2O or 2.72 million tonnes (3.0 million tons) at 1.0% Li_2O in Dyke 1 (Mulligan, 1957a, 1957b). By

March 1958, 12 different tonnage estimates had been made (Northern Miner, March 13, 1958). Also by that time, a permanent camp and a 4-mile road into the property had been built. Plans for a heavy media separation plant on the property were being prepared by the Lummus Co. of New York together with Knowles Associates and the Colorado School of Mines (Green Bay Mining & Exploration Ltd., Corporation File). The description of mineral resources cited above is presented as historical resource estimates and use historical terminology for these estimates. These citations are given to provide an historical frame of reference. Although the resource estimations quoted in the text are believed to be reliable, they were calculated prior to the implementation of National Instrument 43-101. The author has not carried out work to classify these historical estimates under current mineral resource or mineral reserve terminology. The historical estimates are not meant to be interpreted as current estimates as defined in section 1.2 and 1.3 of NI43-101 and should not be relied upon. FAR does not consider these as current resources.

The claims were assigned to J.A. Syme in 1963.

Several airborne surveys were done in the area between 1948 and1973:

1. Inco/1948: Aeromagnetic Survey; Non-confidential assessment file 91614.
2. Canadian Nickel/1957: Airborne Electromagnetic Survey; Non-confidential assessment file 91624.
3. Hudson Bay Exploration and Development/1965: Helicopter-borne Electromagnetic and Radiometric survey; Non-confidential assessment file 91650.
4. Falconbridge Nickel Limited/1973: Airborne Electromagnetic and Magnetic Surveys; Non-confidential assessment file 91564.
   Databases from these surveys are unavailable.

In 1980, J.A. Syme cancelled the **Lit** No. 6-18 claims and obtained Explored Area Lease No. 40 for the **Lit** No. 1-4 claims.

Sampling and detailed geological mapping (Scale 1:1200) of the deposit was done by Cerny *et al.* (1981, p. 155). The analysis of four samples of "core muscovite" had an average content (in weight %) of 0.171% Li, 0.792% Rb, 0.0702% Cs, 0.0021% Be; nine samples of beryl averaged 0.331% Li, 0.903% Na, 0.939% Cs; three samples of spodumene averaged 0.23% Na_2O, 0.943% Fe as Fe_2O_3 (Cerny *et al*., 1981, p. 192).

The **Lit** Nos. 6-18 claims were re-staked under Nor 5 and 6 (W 49000, 49001) by Ross Colon and Moses Crane, respectively, for Noranda Exploration Company Limited in 1983. Fedikow *et al.* (1986) examined quartz veins and outcrop (*c.f.* mineral occurrence RL-95) in the general area. The Nor 6 claim was cancelled in 1987; the Nor 5 in 1988. In 1989, this area was staked as Kelly 3 (P 8412E) by Strider Resources Limited.

The property owner Dalton Bruce Dupasquier optioned the Zoro 1 claim to Force Energy Ltd. of Denver, Colorado (U.S.A.) in 2011. Force Energy defaulted in 2012 and the claim was optioned to Far Resources Ltd. (Vancouver) in 2016. Far Resources completed the acquisition of the Zoro 1 mineral claim as announced on May 9, 2017 in consideration for common shares of the Company and a non-interest bearing promissory note for $100,000 payable in 12 months. Subsequently two option agreements were struck with Strider Resources Limited (Snow Lake, Manitoba) to enlarge the property. In the first option agreement (August 10, 2016) Far Resources increased the size of the property by 600% acquiring an undivided 100% interest in all lithium-bearing pegmatite dykes on Claim Jake 3558 (P3558F) and a 350-metre wide strip along the northeast edge of claim Jake 3558 and a portion of adjacent claims Bert 6304 (MB6304) and Bert 797 (MB797). The claims are contiguous with its Zoro 1 claim. The second option agreement (September 28, 2017) with Strider Resources expanded the property by an additional 2200 hectares. Claims Jake 9 (P3031F), Jake 1054 (MB 1054), Jake 2655 (MB 2655), Jake 3557 (MB 3557), Jake 54199 (W54199), Jake 10 (P3032F), Jake 2412 (MB 2412), Jake 2413 (MB 2413), Jake 54745 (W54745), CRO 5734 (MB 5734) were included in this second option agreement. Recently claims BAZ 12131 (MB12131) and BAZ 12133 (MB12133) have been acquired by Far Resources. The current total area of the property is 3005 hectares.

6.1 HISTORIC RESOURCE WORK

A reserve of "1,727,550 undiluted tons grading 0.945% Li_2O" based on a length of 444' and a vertical extent of 800'" was calculated for the westernmost or "Principal Dyke" on the property (Huston, 1956). This figure was strongly influenced by "an average grade of 2.0% Li_2O over a length of 600' across an average width of 40' in surface exposure."

The parameters of the calculations for the above stated tonnage and grade are presented here with the original units intact (Tables 14.1.1 and 14.1.2).

Table 14.1.1: Parameters for the calculation of grade and tonnage at Dyke 1, lithium-bearing pegmatite dyke.

Horizon/Level	Length (feet)	True Width (feet)	Grade Li_2O (%)
Surface	600	40	2
100'	720	42.5	0.73
330'	165	74.8	0.752
550'	290	87.2	1.052

Table 14.1.2: Combined calculations used to derive grade and tonnage, Dyke 1.

	Length x Width x Depth = Factor	xGrade%	Grade Factor
Surface	600'x40'x50'=1,200,000	2.0	2,400,000
100'	720'x42.8'x140'=4,314,240	0.73	3,149,395
330'	165'x74.8x250'=3,085,500	0.752	2,314,125
550'	290'x87.2'x220'=5,563,360	1.052	5,841,528
Average	444'x52'x660'=14,163,100	0.967	13,705,048

Based on an 11-cubic foot per ton factor (in place): 14,163,100/11=1,287,550 tons to 660'. Diamond drill hole 17 (zl-56-017intersected a zone grading 0.882% over 204.4' at 780 vertical feet below surface. This intersection is interpreted to be 0.882% across 110 feet true width and is assumed for the purposes of this calculation to extend 100' laterally and 220' vertically.

This gives:

- *200' long x 110' wide x 220' depth/11 = 440,000 tons grading 0.882%.*

The final historic resource estimate produced for the westernmost dyke was:

- *1,287,550 tons grading 0.967% Li_2O to a depth of 660’ vertically below surface outcrop plus a possible 440,000 tons grading 0.882% to a depth of between 660’ and 880’ below outcrop.*

The total undiluted tonnage was given as 1,727,550 at 0.945% Li_2O. An appropriate dilution factor was given as 5%.

The description of mineral resources cited above is presented as an historical resource estimate and uses historical terminology for these estimates. These citations are given to provide an historical frame of reference. Although the resource estimations quoted in the text are believed to be reliable, they were calculated prior to the implementation of National Instrument 43-101. The author has not carried out work to classify these historical estimates under current mineral resource or mineral reserve terminology. The historical estimates are not meant to be interpreted as current estimates as defined in section 1.2 and 1.3 of NI43-101 and should not be relied upon. FAR is not treating these as current resources.

6.2 GREEN BAY MINING & EXPLORATION LTD. (PREVIOUSLY GREEN BAY URANIUM LIMITED)

Green Bay Mining and Exploration Ltd. (Green Bay) undertook basic prospecting, rock chip sampling from blasted trenches and assay work prior to a major drill campaign between 1956 and 1958. Most work was undertaken on the seven known dykes on this property (Figure 9.1.1) with emphasis on Dyke 1 where assays from trench samples documented >2% lithium (Northern Miner, January 12, 1956). A geological map was produced for Dyke 1 at a scale of 1”:50 feet by C.C. Huston and Associates acting for Green Bay (Figure 9.1.2) showing trench locations. Dykes 2 through 7 were also assessed by rock chip sampling and assay with Dyke 5 returning assays of 2.42% to 7.28% Li_2O. Numerous additional assays are available for this work by Green Bay however assay certificates do not contain details of when and where the assay sample was collected from drill core. Assays were performed by Correlation Laboratories Ltd. of Cobden, Ontario.

Over 6096 m (20 000 ft.) of diamond drilling was done on the property with a minimum of 3048 m (10 000 ft.) of this on Dyke 1. Results of the drilling on dykes 1, 3, 5 and 7 were reported to be "promising". Gold was also found on the property, with a 3.3 kg (7.25 lb.) sample across 3.4 m (11 ft.) yielding 0.17 ounces per ton gold (Green Bay Exploration Corporation File).

Review of historic drill and assay information was significantly hampered by the lack of control over historic drill collar locations and information documenting the location of the drill hole upon completion. No down hole information is available. Assay results are incomplete with location of the sample collected for assay not detailed. For the construction of three-dimensional imagery for Dykes 1 through 7 the approach has been to use available historic information supplemented by the identification of recognizable collar locations in the field.

6.2.1 GEOLOGY OF DYKE 1 (FIGURES 9.1.1 AND 9.1.2)

The historic geologic map in Figure 9.1.2 is based on available outcrop at the time of exploration by Green Bay. Dyke 1 is a 280 m long sinuous pegmatite dyke with coarse grained spodumene that intrudes variably textured mafic volcanic rocks and a variety of sedimentary rocks. The north end of the dyke is marked by a faulted extension that forms a northeast-trending arm of spodumene-bearing pegmatite. The contact between the pegmatite dyke and adjacent sediments is schistose. Outcrop is scarce in the surrounding area.

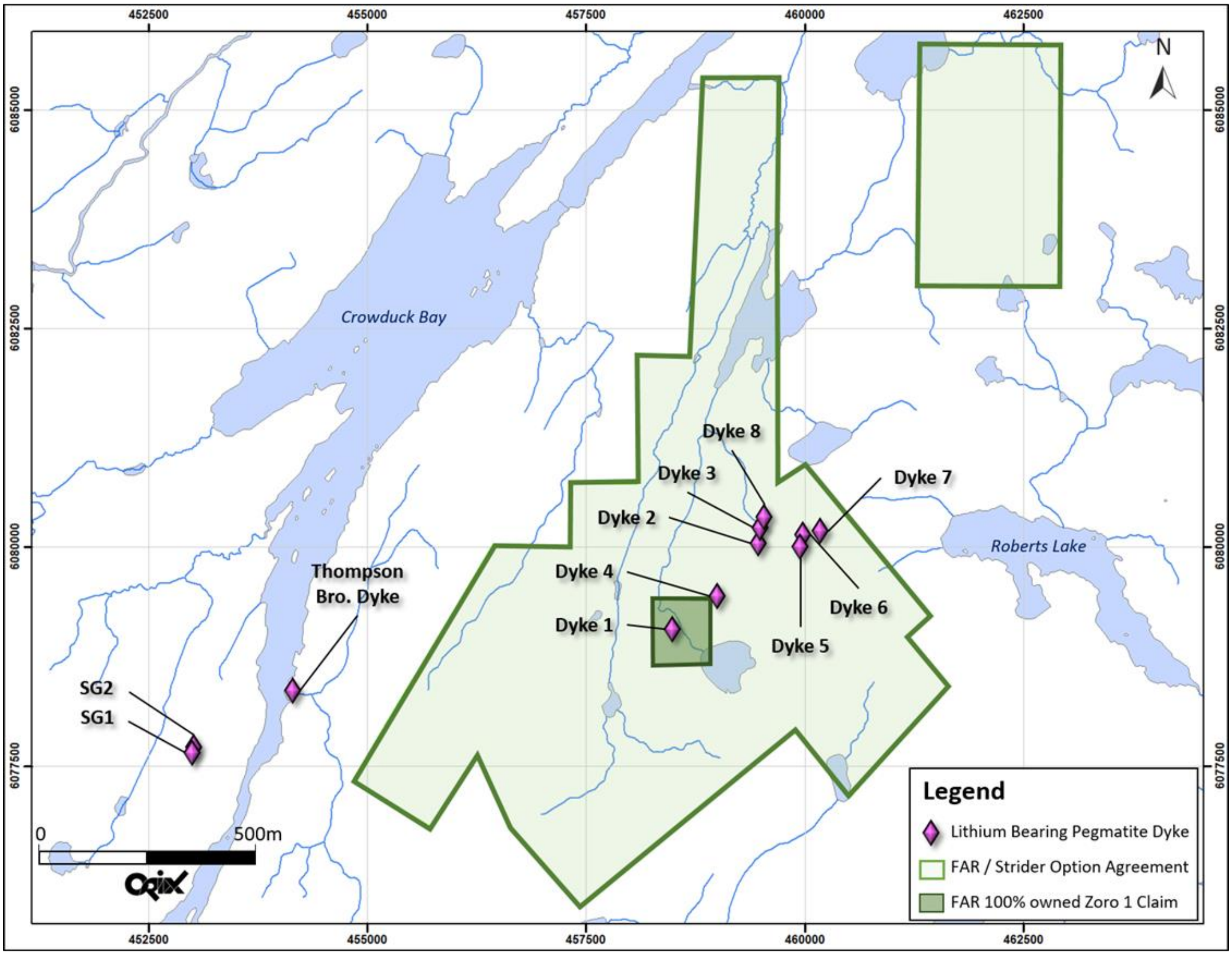


Figure 9.1.1: Location map for the seven known pegmatite dykes comprising the Zoro Lithium Project.

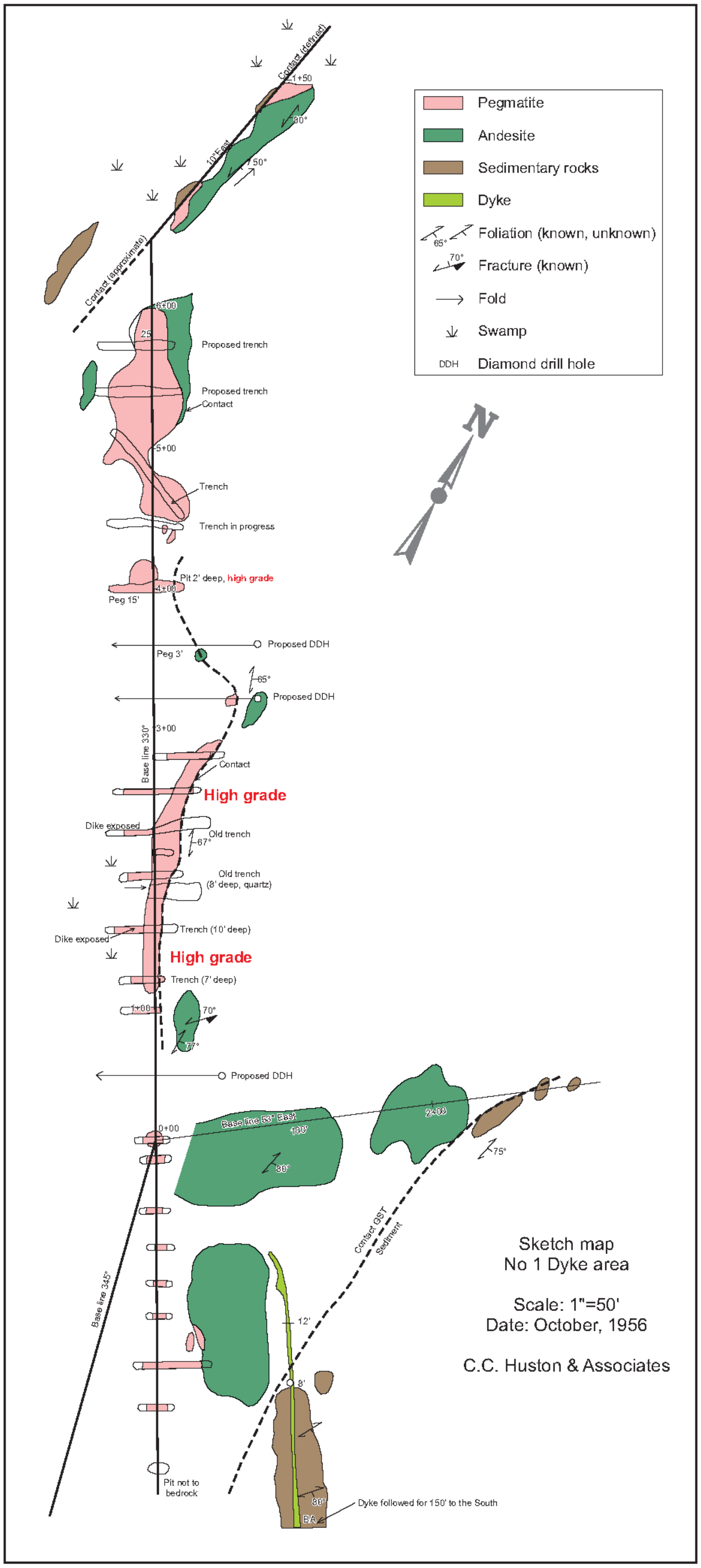


Figure 9.1.2: Outcrop, geology and trench location map, Dyke 1, Zoro Lithium Project. Data from historic information sourced in Manitoba Mining Recorder's assessment files.

6.3 FORCE ENERGY LTD.

Exploration commencing in 2012 by Force Energy was designed to assess the lithium potential of Dyke 1 on the Zoro 1 claim. The work consisted of cleaning out, washing, channel and chip sampling of pegmatite exposed in historic trenches exposed on the property. Samples were analyzed for lithium and a multielement suite at Activation Laboratories (Ancaster, Ontario). The locations of individual trenches that were sampled are given in Figure 9.2.1. Figures 9.2.2 through 9.2.8 give sample numbers and sampling intervals for each trench. Table 9.2.1 summarizes the channel and chip samples collected from the Zoro 1 claim. A total of 165 channel samples were cut with a rock saw and 5 representative chip samples were collected using a sledge hammer and chisel. Weight averaged Li_2O assay data from each trench are summarized in Table 9.2.2. Analytical results for channel samples documented elevated lithium concentrations in the spodumene-bearing pegmatite on the Zoro 1 claim.

Table 9.2.1: Summary of channel and grab samples collected from the Zoro 1 claim. UTM coordinates (datum NAD 83, Zone 14).

Trench	Samples	UTM North	UTM East
Trench-01	ZR-1 to ZR-7	6078940	458534
Trench-02	ZR-8 to ZR-12	6078941	458524
Trench-03	ZR-13 to ZR-16	6078949	458520
Trench-04	ZR-17 to ZR-22	6078964	458518
Trench-05	ZR-23 to ZR-28	6078970	458513
Trench-06	ZR-29 to ZR-36	6079015	458498
Trench-07	ZR-37 to ZR-46	6079025	458495
Trench-08	ZR-47 to ZR-55	6079035	458496
Trench-09	ZR-56 to ZR-64	6079040	458494
Trench-10	ZR-65 to ZR-81	6079045	458494
Trench-11	ZR-82 to ZR-90	6079061	458496
Trench-12	ZR-91 to ZR-105	6079102	458474
Trench-13	ZR-106 to ZR-120	6079108	458473
Trench-14	ZR-121 to ZR-142	6079121	458470
Trench-15	ZR-143 to ZR-154	6079138	458463
Trench-16	ZR-155 to ZR-165	6079146	458458
Representative Chip Sample	ZR-166	6079295	458869
Representative Chip Sample	ZR-167	6079305	458870
Representative Chip Sample	ZR-168	6079313	458875
Representative Chip Sample	ZR-169	6079315	458876
Representative Chip Sample	ZR-170	6079317	458879

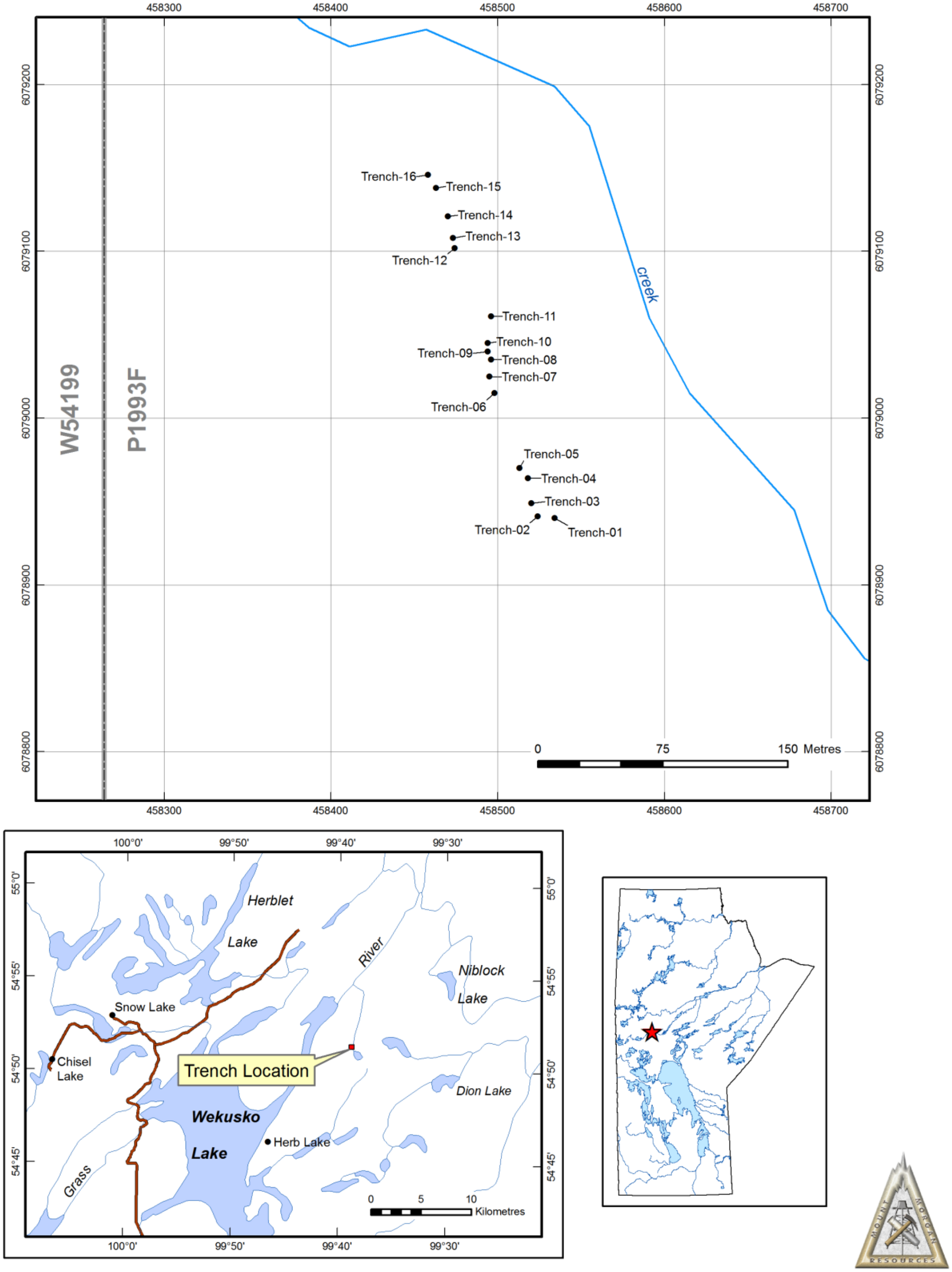


Figure 9.2.1: Location of historic trenches on the Zoro 1 claim, Dyke 1 that have been mucked out, washed and channel sampled.

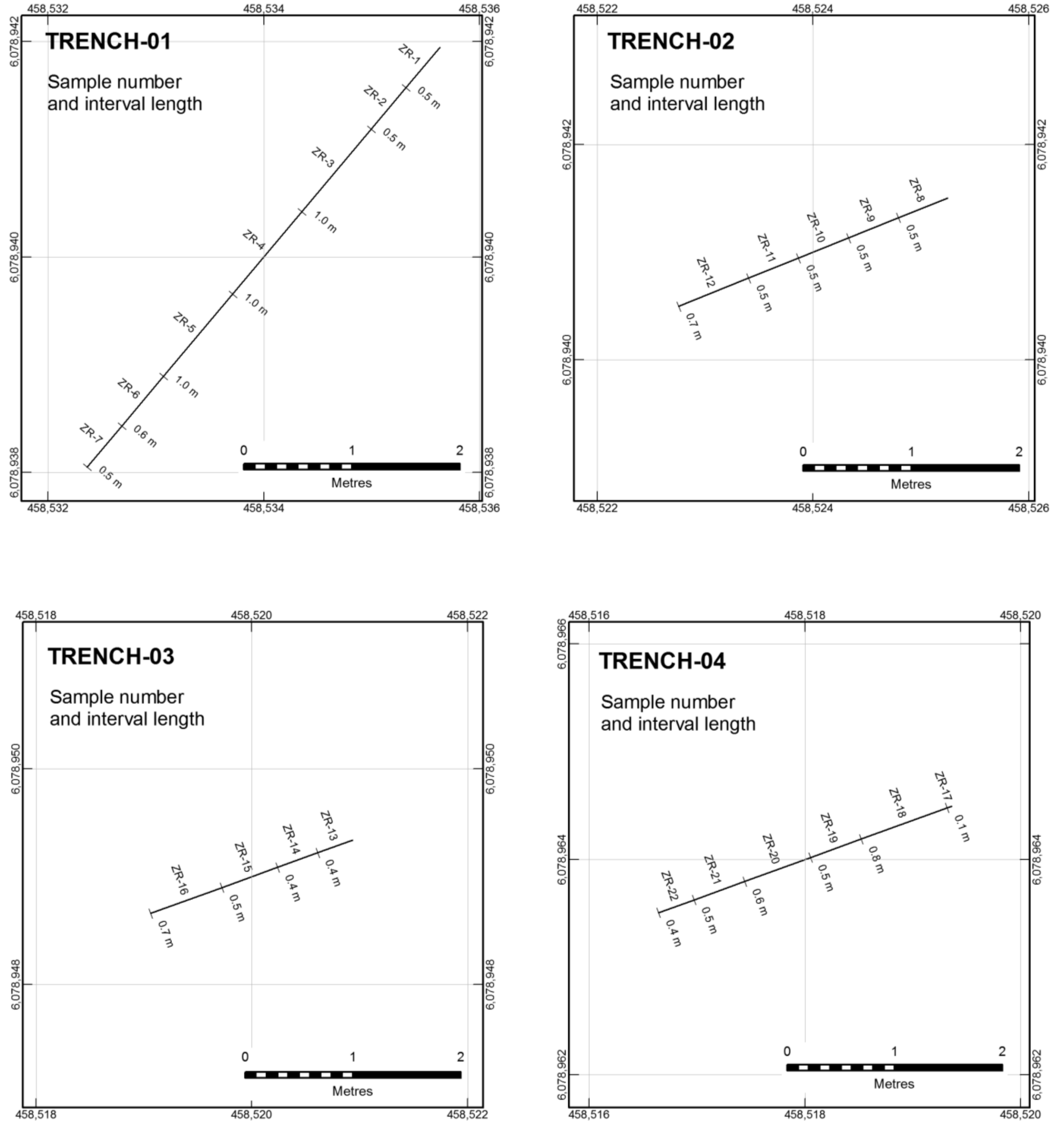


Figure 9.2.2: Channel samples (number and sampling interval) for trenches 01, 02, 03 and 04, Zoro 1 claim.

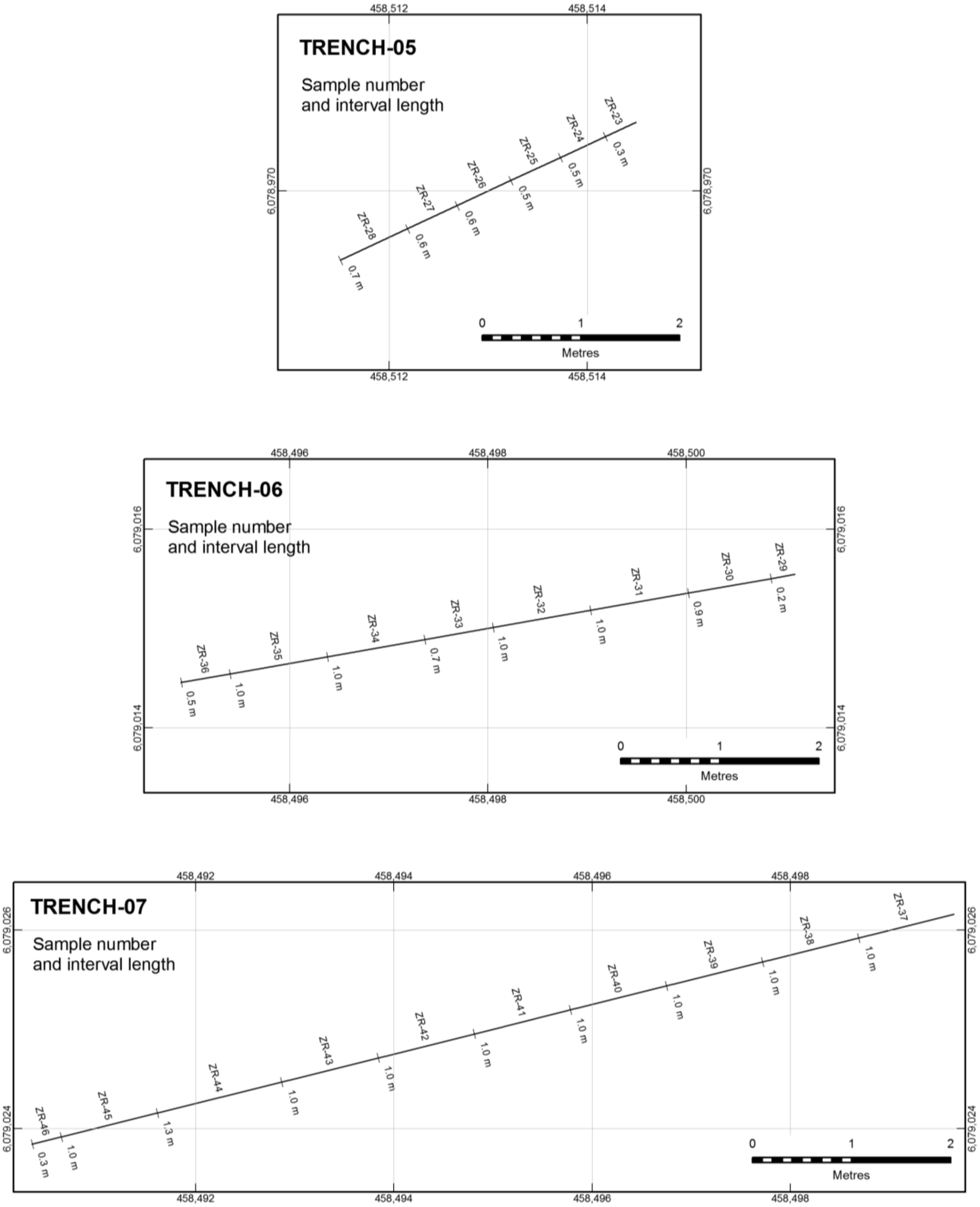


Figure 9.2.3: Channel samples (number and sampling interval) for trenches 05, 06 and 07, Zoro 1 claim.

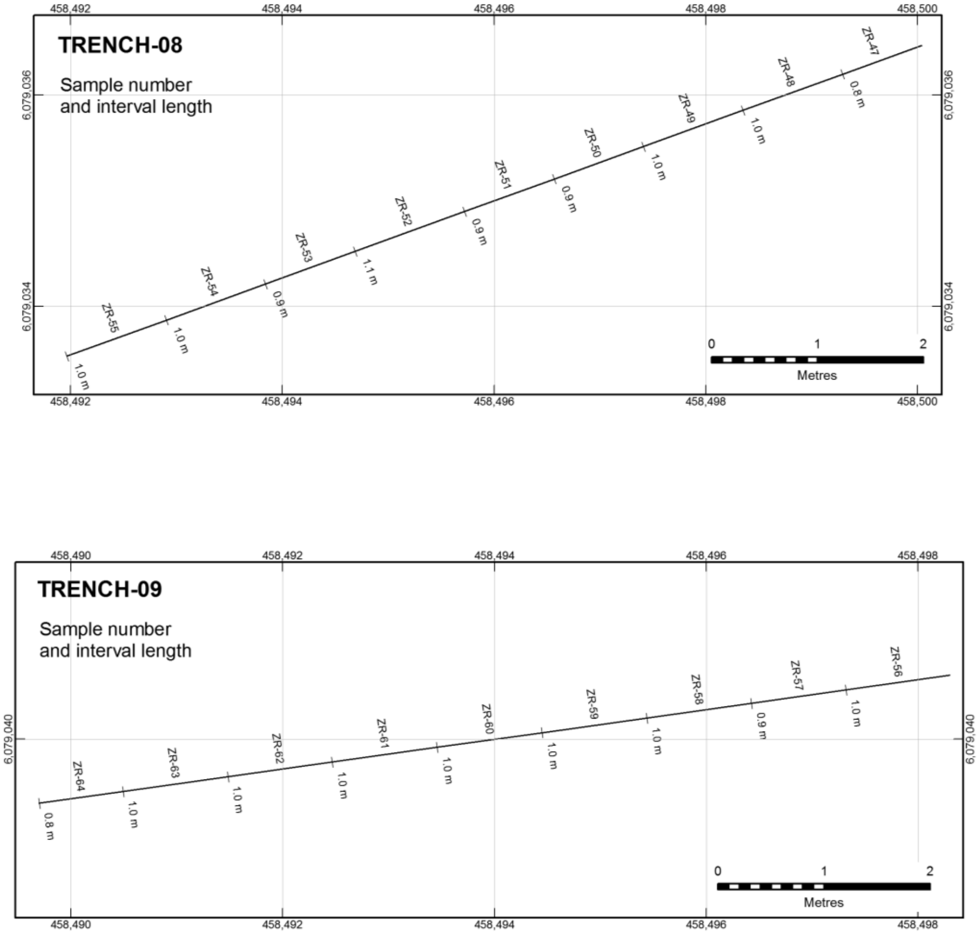


Figure 9.2.4: Channel samples (number and sampling interval) for trenches 08 and 09, Zoro 1 claim.

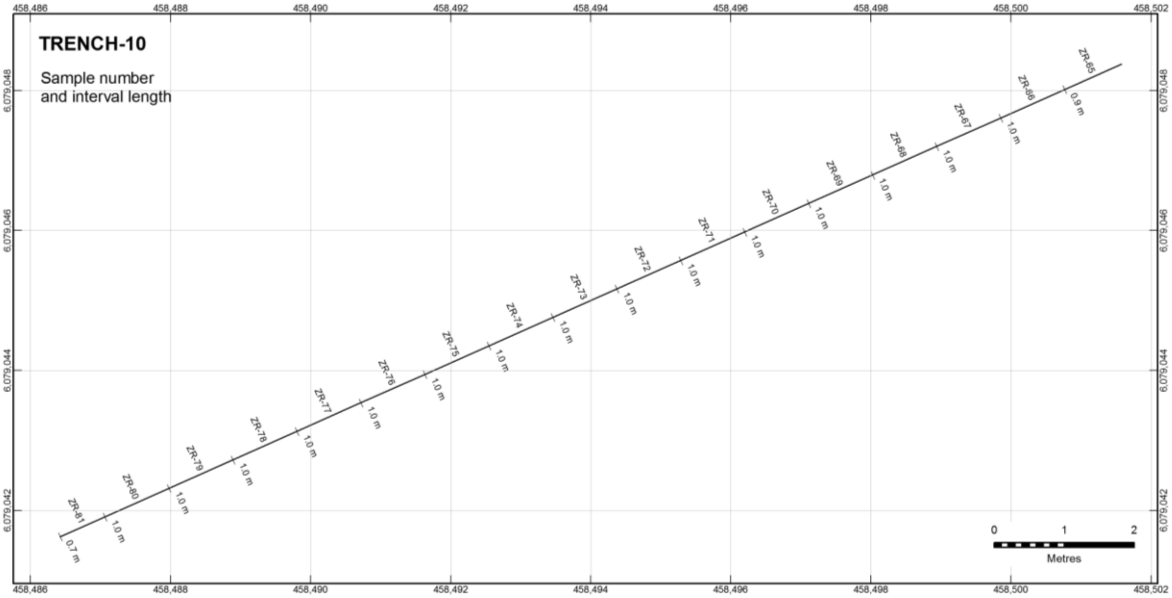


Figure 9.2.5: Channel samples (number and sampling interval) for trench 10, Zoro 1 claim.

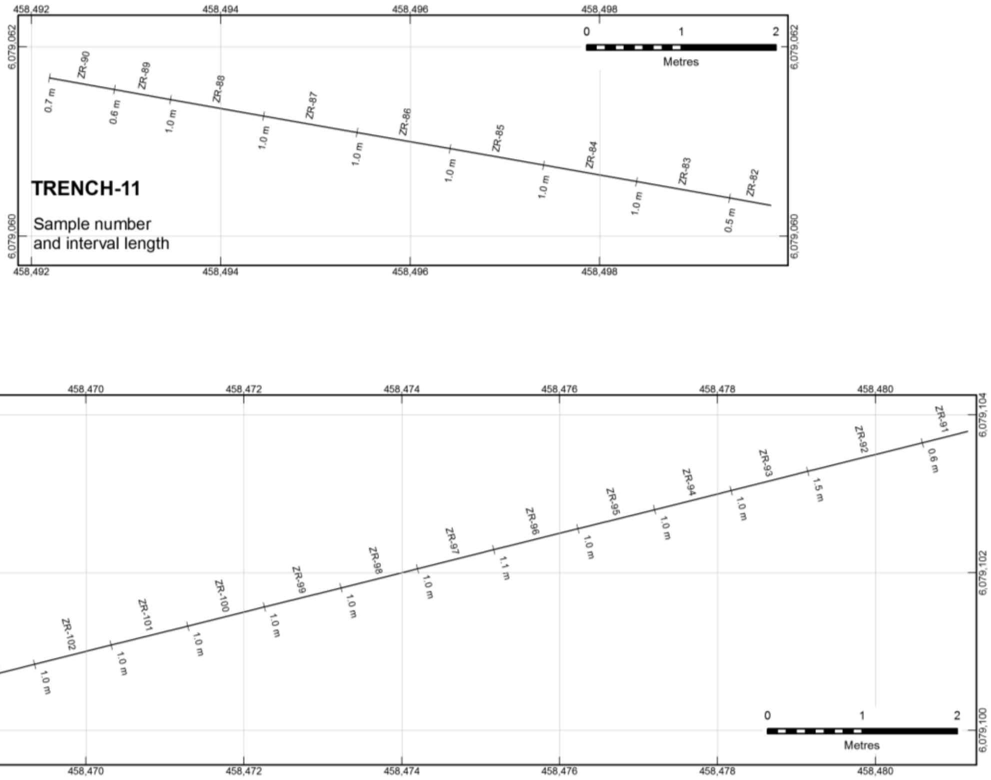


Figure 9.2.6: Channel samples (number and sampling interval) for trenches 11 and 12, Zoro 1 claim.

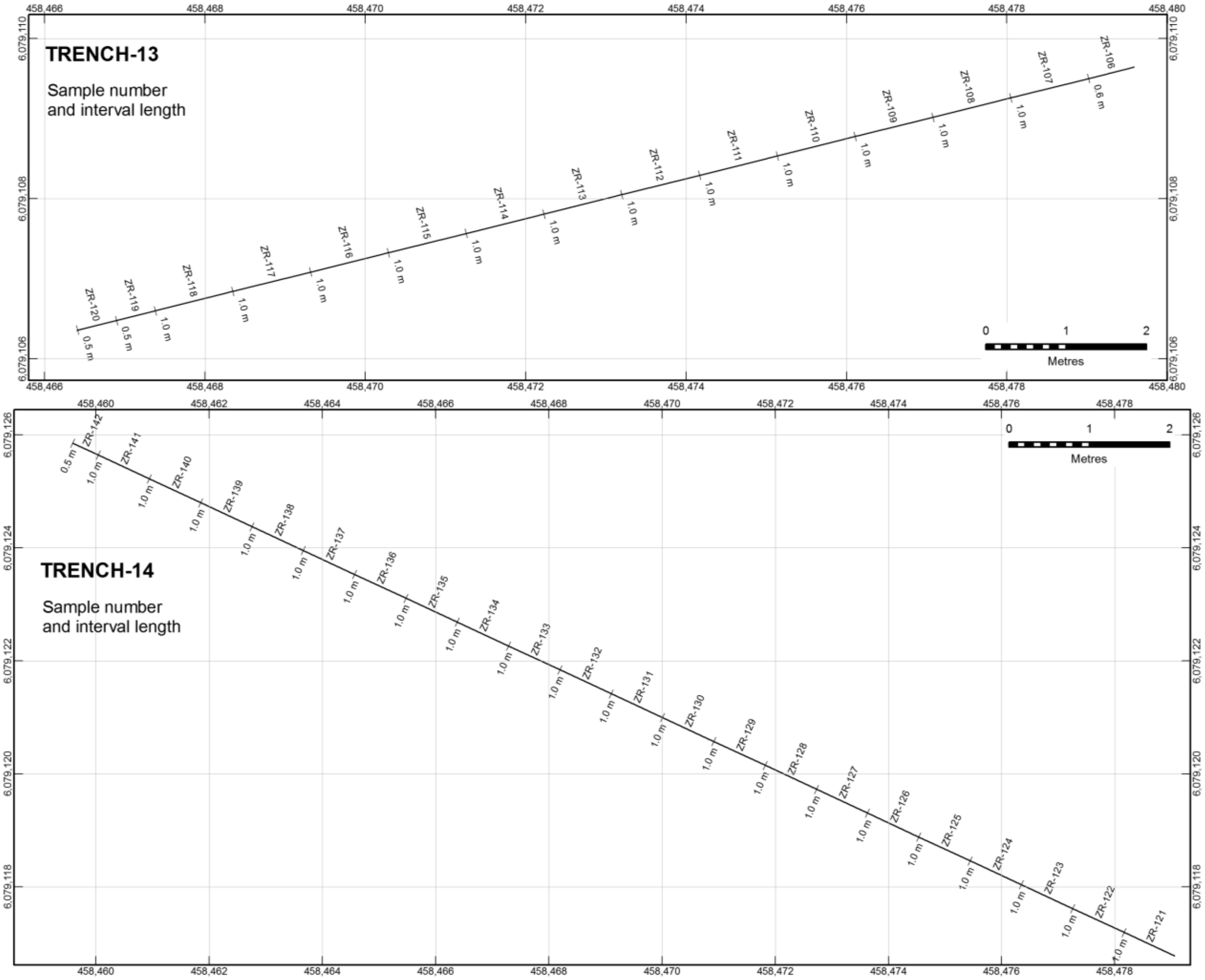


Figure 9.2.7: Channel samples (number and sampling interval) for trenches 13 and 14, Zoro 1 claim.

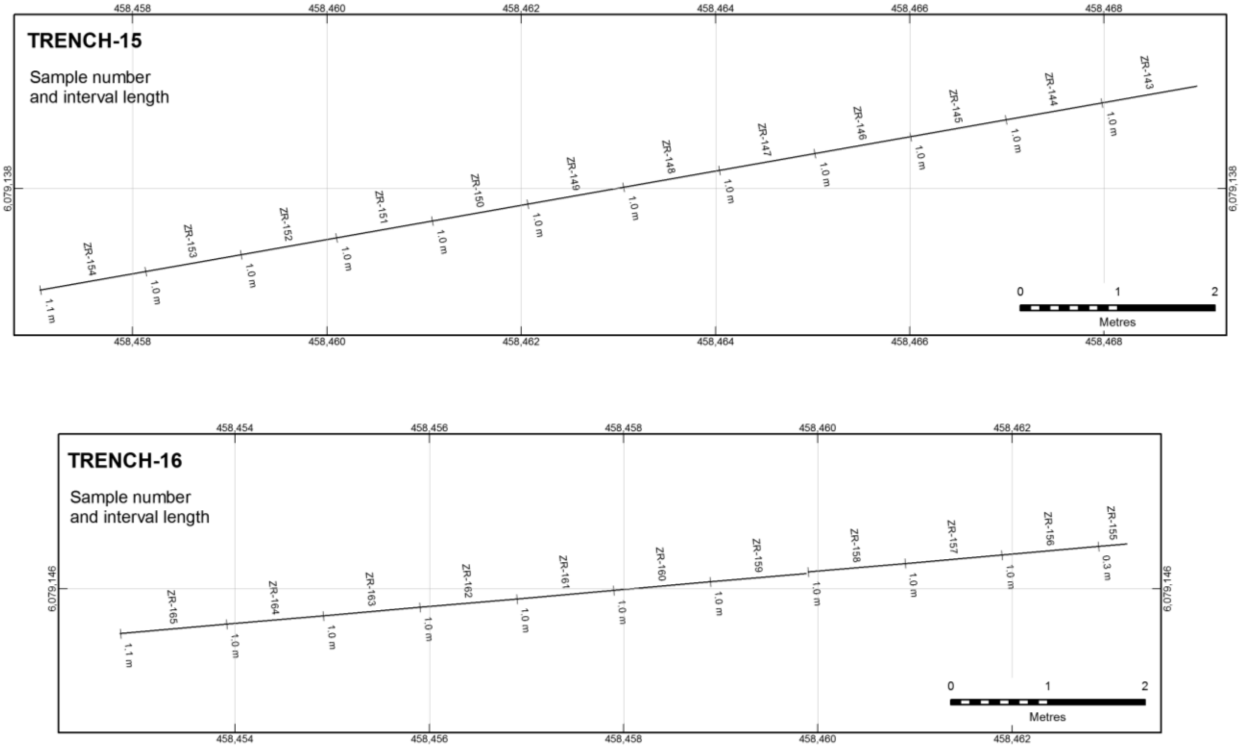


Figure 9.2.8: Channel samples (number and sampling interval) for trenches 15 and 16, Zoro 1 claim.

Table 9.2.2: Summary of lithium analyses presented as weighted averages per trench and conversion to Li_2O and Li_2CO_3.

Trench (n of samples)	Weighted Average Li (ppm) Per Trench	% Li_2O Per Trench	% Li_2CO_3 Per Trench
1 (n=7)	1225	0.26	0.65
2 (n=5)	1531	0.33	0.82
3 (n=4)	1966	0.42	1.05
4 (n=5)	2610	0.56	1.39
5 (n=6)	884	0.19	0.47
6 (n=8)	3028	0.65	1.61
7 (n=10)	2069	0.45	1.1
8 (n=9)	1930	0.42	1.02
9 (n=9)	3350	0.72	1.78
10 (n=17)	2408	0.52	1.28
11 (n=9)	2852	0.61	1.51
12 (n=15)	2332	0.5	1.24
13 (n=15)	1392	0.3	0.74
14 (n=22)	2310	0.5	1.23
15 (n=12)	1521	0.33	0.8
16 (n=11)	1639	0.35	0.87

6.4 HISTORICAL DRILLING

Initial drilling on the Zoro dykes (historically referred to as the LIT Group) occurred between 1956 and 1957 and was completed by Green Bay Uranium Group, based in Edmonton, Alberta. Seven major dykes were initially identified due to their exposure in outcrop and were subsequently uncovered through trenching and blasting. Six of these dykes were targeted with drilling to test their depth extent. Historical drilling is reported in cancelled assessment report #93562. Historical drilling comprises 78 diamond drillholes totaling 8,469.3 m (27,786.6 ft). Table 10.1.1 summarizes the breakdown of drilling on each of the 7 major dykes.

Table 10.1.1: Summary of historical drilling by pegmatite dyke.

Dyke	Number of DDH	Total Feet Drilled	Total Meters Drilled
Dyke 1	48	20,015.6 ft	6,100.8 m
Dyke 2	11	2,689.0 ft	819.6 m
Dyke 3	4	864.0 ft	263.3 m
Dyke 4	3	612.0 ft	186.5 m
Dyke 5	8	2,840.0 ft	865.6 m
Dyke 6	*****Not Drilled*****		
Dyke 7	4	766.0 ft	233.5 m
TOTAL	**78**	**27,786.6 ft**	**8,469.3 m**

The collar locations for these historical drillholes were referenced to a local grid with no modern reference control points, therefore their exact location on the ground remains uncertain. Planview maps for Dykes 1, 2, 3, 4, and 7 were included in assessment report #93562 showing the location of the drillhole collars relative to each other. The historical drillhole collar locations can also be found in the GIS Map Gallery drillhole database maintained by the Mineral Resources Division of the Government of Manitoba Department of Growth, Enterprise, and Trade. These collar locations approximated the relative spatial arrangement of the report plan view maps, however appeared to be displaced approximately 150m northeast from where the dykes are located on the ground.

Far Resources mobilized a reconnaissance field crew to ground-truth the locations of the collars. This effort permitted successful verification of 19 historical collar locations around Dyke 1 using the historical plan view map as a guideline. UTM coordinates for these 19 historical collar locations were recorded by a handheld Garmin GPS and plotted in ArcGIS (Table 10.1.2, Figure 10.1.1). Utilizing these 18 collars, the guideline plan view map was rubber sheeted into the correct geographic space, and the remaining 29 drillholes around Dyke 1 were digitized into place.

The reconnaissance crew also identified surface trenches. However, the crew were unable to identify drill collars for the remaining 6 dykes; therefore, the true locations for the remaining 48 collar locations cannot be verified. The surficially exposed trenches were used as control points for the historical plan view maps to approximate collar placement for the 48 unverified collar locations targeted under trenches. The collars were placed as close to their true location as could be interpreted based on surface exposure of the dykes and location of trenches and pits. Drillhole information for all 78 historical drillholes are listed in Table 10.1.3. Figure 10.1.2 illustrates the nature of much of the historic infrastructure on the Zoro Lithium Project.

Table 10.1.2: Historical diamond drillhole collar locations identified by Far Resources field crew. (*UTM coordinates displayed in NAD 83, Zone 14N.)

BHID	Easting*	Northing*	Waypoint
zl-56-011	458555	6079091	WP0001
zl-56-029 and zl-56-032	458491	6079131	WP0002
zl-56-069	458511	6079076	WP0003
zl-56-068	458491	6078976	WP0004
zl-56-063	458495	6078992	WP0005
zl-56-048	458521	6078897	WP0006
zl-56-049	458528	6078881	WP0007
zl-56-075	458530	6078856	WP0008
zl-56-022	458369	6078943	WP0009
zl-56-021	458353	6079017	WP0010
zl-56-020	458385	6079023	WP0011
zl-56-071	458500	6079103	WP0012
zl-56-001 or zl-56-073	458504	6079092	WP0013
zl-56-067	458472	6079002	WP0014
zl-56-064	458486	6079005	WP0015
zl-56-018	458334	6079049	WP0016
zl-56-074	458508	6079040	WP0017
zl-56-078	458620	6079043	WP0018

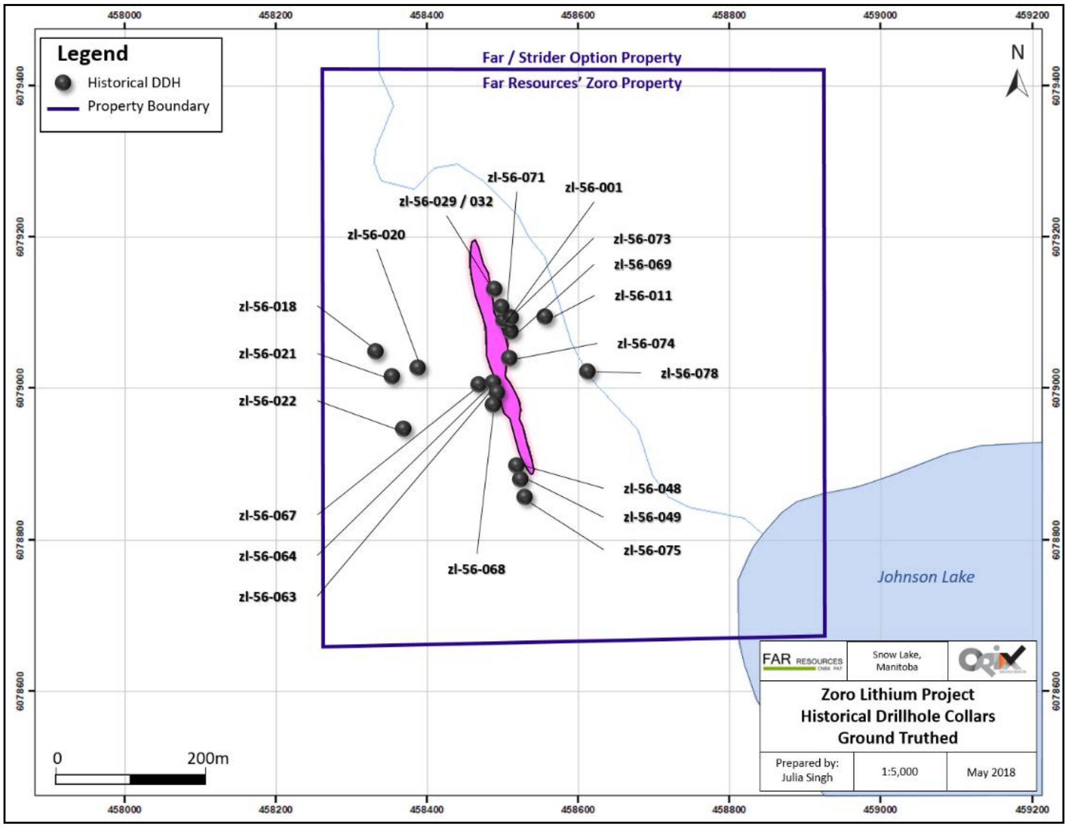


Figure 10.1.1: Historical drillhole collars located by Far Resources.

Figure 10.1.2: Drill core in collapsed core racks from the Green Bay Exploration and Development historic diamond drill programs, Zoro Lithium Project.

Table 10.1.3: List of historical drillhole information.

BHID	Easting*	Northing*	Azimuth	Dip	EOH (m)	Target	Pegmatite Intervals (m)
zl-56-001	458504	6079092	253°	-45°	66.45	Dyke 1	29.50-58.46
zl-56-002	458503	6079117	253°	-45°	56.39	Dyke 1	36.27-50.13
zl-56-003	458492	6079142	253°	-45°	75.29	Dyke 1	58.73-69.18
zl-56-004	458519	6079055	253°	-45°	71.93	Dyke 1	24.50-58.67
zl-56-005	458527	6079024	253°	-45°	76.20	Dyke 1	44.59-68.76
zl-56-006	458534	6078997	253°	-45°	67.97	Dyke 1	48.46-61.87
zl-56-007	458542	6078969	253°	-45°	67.67	Dyke 1	43.58-56.38
zl-56-008	458548	6078943	253°	-45°	68.88	Dyke 1	11.73-12.13 41.97-45.87
zl-56-008A	458559	6078915	253°	-45°	58.83	Dyke 1	41.75-42.51
zl-56-011	458555	6079091	253°	-55°	172.82	Dyke 1	131.36-132.16 139.75-140.90 146.21-148.83
zl-56-012	458551	6079057	253°	-55°	163.98	Dyke 1	115.51-117.10 122.68-127.55 130.42-155.41
zl-56-013	458571	6079032	253°	-55°	142.95	Dyke 1	108.47-137.98
zl-56-014	458580	6079005	253°	-55°	145.39	Dyke 1	125.27-134.75
zl-56-015	458586	6078975	253°	-55°	136.86	Dyke 1	122.19-132.25
zl-56-016	458595	6078948	253°	-55°	122.53	Dyke 1	109.81-116.64
zl-56-017	458615	6079039	253°	-60°	333.45	Dyke 1	207.87-212.65 230.73-232.25 263.34-325.64
zl-56-018	458334	6079049	73°	-65°	289.56	Dyke 1	189.15-195.95 201.28-262.82
zl-56-019	458441	6078997	73°	-45°	101.80	Dyke 1	54.92-92.38
zl-56-020	458385	6079023	73°	-50°	211.84	Dyke 1	125.15-126.55 136.73-142.55
zl-56-021	458353	6079017	73°	-60°	238.05	Dyke 1	165.81-226.58
zl-56-022	458369	6078943	73°	-67°	329.18	Dyke 1	None
zl-56-022A	458369	6078943	77°	-62°	460.55	Dyke 1	300.31-302.30
zl-56-023	458342	6079071	75°	-60°	282.85	Dyke 1	163.98-164.01 171.90-175.44 230.42-231.22
zl-56-024	458332	6079101	75°	-60°	351.01	Dyke 1	160.62-161.11
zl-56-025	458453	6079231	113°	-45°	112.47	Dyke 1	None
zl-56-026	458439	6079250	113°	-45°	134.42	Dyke 1	40.53-41.94
zl-56-027	458262	6078964	75°	-60°	410.26	Dyke 1	78.51-78.94 358.74-360.70 364.84-368.89 377.83-378.13 378.43-379.29
zl-56-028	458430	6079022	147°	-60°	103.33	Dyke 1	21.03-21.09 22.55-22.61 29.87-30.60

BHID	Easting*	Northing*	Azimuth	Dip	EOH (m)	Target	Pegmatite Intervals (m)
zl-56-029	458491	6079131	253°	-30°	73.76	Dyke 1	25.29-32.88 44.80-45.44 46.51-48.09 53.27-54.86 57.05-64.80
zl-56-030	459480	6080047	218°	-40°	84.43	Dyke 2**	55.04-58.52
zl-56-031	459517	6079993	268°	-35°	64.01	Dyke 2**	48.06-49.04
zl-56-032	458491	6079131	253°	-45°	76.50	Dyke 1	26.76-30.48 44.80-47.54 48.40-49.19 53.15-54.25 60.35-60.74 63.97-67.90
zl-56-033	460149	6080195	11°	-45°	35.36	Dyke 7**	7.62-12.98
zl-56-034	459411	6079998	63°	-50°	83.82	Dyke 2**	29.74-30.02 63.24-65.95
zl-56-035	460150	6080176	23°	-50°	52.43	Dyke 7**	4.45-4.87 9.75-10.24
zl-56-036	459502	6080022	243°	-45°	77.42	Dyke 2**	24.07-24.65
zl-56-037	459435	6080115	288°	-45°	64.01	Dyke 2**	44.80-46.93
zl-56-038	459439	6080116	218°	-45°	64.01	Dyke 2**	45.04-54.28
zl-56-039	460021	6079971	67°	-40°	20.42	Dyke 5**	11.88-12.98 13.41-14.56
zl-56-040	459379	6080080	63°	-55°	80.16	Dyke 2**	61.75-67.51
zl-56-041	460024	6079967	353°	-30°	43.59	Dyke 5**	14.11-20.42 24.56-24.78
zl-56-042	458449	6079134	38°	-30°	67.06	Dyke 1	2.25-2.43 36.14-36.97 46.45-47.97
zl-56-043	458534	6078950	253°	-45°	34.14	Dyke 1	1.82-3.04 25.17-28.62
zl-56-044	459479	6080199	93°	-40°	39.62	Dyke 3**	15.63-16.70 23.16-25.14
zl-56-045	459479	6080200	31°	-30°	51.82	Dyke 3**	12.03-12.86 26.21-31.79
zl-56-046	458539	6078934	253°	-45°	36.58	Dyke 1	12.83-13.5 20.17-21.70 24.50-28.59
zl-56-047	458516	6078910	73°	-45°	30.18	Dyke 1	19.59-25.20
zl-56-048	458521	6078897	73°	-45°	41.45	Dyke 1	4.26-5.30
zl-56-049	458528	6078881	73°	-45°	47.55	Dyke 1	8.32-9.14
zl-56-050	460032	6080048	223°	-45°	183.49	Dyke 5**	16.45-16.61 24.56-25.17 72.66-72.93 73.39-73.97 75.62-76.10
zl-56-051	459970	6079962	43°	-45°	125.58	Dyke 5**	34.19-34.86 47.64-49.59 76.41-76.50 94.33-94.79
zl-56-052	459943	6079962	43°	-43°	135.94	Dyke 5**	23.22-23.95 52.91-53.94 64.49-64.80 87.99-90.15 90.70-91.89 100-103.54

BHID	Easting*	Northing*	Azimuth	Dip	EOH (m)	Target	Pegmatite Intervals (m)
zl-56-053	459916	6079991	43°	-45°	117.35	Dyke 5**	34.47-34.71 35.14-35.84 48.24-48.88 85.19-86.95 102.96-103.93
zl-56-054	459911	6080030	43°	-50°	118.57	Dyke 5**	11.06-13.99 27.12-28.25 64.80-65.83 88.97-91.13
zl-56-055	459882	6080043	43°	-50°	120.70	Dyke 5**	15.24-15.81 16.88-17.03 17.83-19.05 29.13-29.80 32.61-33.00 67.97-68.12 94.85-96.98
zl-56-056	460116	6080181	23°	-30°	69.19	Dyke 7**	6.40-6.88 20.6-25.78
zl-56-057	460150	6080194	77°	-45°	76.50	Dyke 7**	6.70-9.75 28.34-29.56 32.00-32.18 55.47-55.77
zl-56-058	458604	6078912	253°	-35°	99.36	Dyke 1**	58.46-58.85
zl-56-059	459438	6080200	48°	-35°	97.54	Dyke 3**	23.71-25.23 51.17-51.57 52.12-52.48 60.62-61.53
zl-56-060	459433	6079957	63°	-50°	74.98	Dyke 2**	33.95-34.35 58.97-64.00
zl-56-061	459399	6080050	63°	-45°	88.39	Dyke 2**	57.42-59.55
zl-56-062	459372	6080122	63°	-45°	69.80	Dyke 2**	35.35-35.93
zl-56-063	458495	6078992	73°	-35°	20.42	Dyke 1	9.66-14.66
zl-56-064	458486	6079005	73°	-40°	12.80	Dyke 1	4.26-12.80
zl-56-065	458448	6079002	73°	-35°	76.81	Dyke 1	34.59-64.98
zl-56-066	459437	6080201	93°	-35°	74.37	Dyke 3**	26.24-32.27 37.21-37.61 52.54-53.00 58.55-58.85 63.88-64.25
zl-56-067	458472	6079002	73°	-35°	53.34	Dyke 1	17.03-36.33 38.25-38.95
zl-56-068	458491	6078976	73°	-40°	26.82	Dyke 1	19.81-22.55
zl-56-069	458511	6079076	253°	-35°	60.96	Dyke 1	18.68-51.48
zl-56-070	459473	6079947	63°	-45°	68.58	Dyke 2**	3.96-5.18
zl-56-071	458500	6079103	253°	-35°	52.12	Dyke 1	19.81-42.27
zl-56-072	459013	6079372	62°	-45°	46.02	Dyke 4**	26.94-27.79 34.89-36.33
zl-56-073	458509	6079091	253°	-35°	48.16	Dyke 1	18.19-38.80
zl-56-074	458508	6079040	253°	-40°	57.61	Dyke 1	20.51-49.92
zl-56-075	458530	6078856	73°	-45°	73.15	Dyke 1	67.69-67.84
zl-56-076	458939	6079458	62°	-40°	62.48	Dyke 4**	47.09-54.10
zl-56-077	459015	6079483	242°	-40°	78.03	Dyke 4**	35.69-35.84
zl-56-078	458612	6079022	248°	-55°	259.08	Dyke 1	169.89-181.50 182.08-182.33

BHID	Easting*	Northing*	Azimuth	Dip	EOH (m)	Target	Pegmatite Intervals (m)
							183.18-183.73 184.49-191.20 199.12-205.61 207.75-219.97 220.73-222.80 223.63-225.94 228.60-246.15

***UTM coordinates in NAD 83, Zone 14N**
****Historical drillhole collar locations for Dykes 2-7 cannot be verified.**

Included in assessment report #93562 were several assay results reported in Li_2O% for the major dykes. However, it was not possible for Far Resources to match sample ID to the corresponding drillhole nor to a specified interval because the naming convention was inconsistent and the drill logs themselves rarely identified sample intervals. Four historical drillholes, however, were re-sampled in 1957 for use in a resource estimate. That assay documentation could be easily related back to the four historical drillholes. Far Resources calculated composites to be used for drillhole targeting. The four historical composites are listed in Table 10.1.4.

Table 10.1.4: Historical drillhole weighted averages.

BHID	Composite
zl-56-005	1.2% Li_2O over 22.8m
zl-56-013	1.0% Li_2O over 18.2m
zl-56-017 and	1.2% Li_2O over 7.9m 1.0% Li_2O over 15.2m
zl-56-021	1.4% Li_2O over 49.8m

7.0 GEOLOGICAL SETTING AND MINERALIZATION

7.1 BEDROCK GEOLOGY - INTRODUCTION

The Paleoproterozoic Flin Flon-Snow Lake Belt is approximately 200 km in strike length and has an exposed width of up to 70 km. The Belt is overlain to the south by Ordovician Red River Formation sandstone, limestone, and dolomite of the Western Canada Sedimentary Basin, and is bordered to the north by high-grade paragneiss and granitoid rocks of the Kisseynew Domain (Figure 7.1.1).

The Flin Flon Belt is interpreted to be an accreted assemblage of oceanic to continental margin arc terrane, interspersed with oceanic basins representing back-arc, fore-arc, and oceanic settings (Lucas *et al*., 1996; Syme *et al*., 1996). It is part of the Reindeer zone, a largely juvenile portion of the Trans-Hudson Orogen separating the Archean Superior and Hearne provinces (Figure 7.1.1, 7.1.2 and 7.1.3). Recent tracer isotope studies have confirmed the presence of >3.0 Ga Archean crust (the Sask Craton) below parts of the Trans-Hudson (Lucas *et al*., 1996). The Shield Margin National Mapping Program (NATMAP; Lucas *et al*., 1996) traced the Flin Flon Belt assemblages below the Phanerozoic to the south and recognized highly metamorphosed and deformed Flin Flon volcanic and sedimentary formations (Zwanzig, 1990, 1999) within the Kisseynew Domain to the north. To the east, the Flin Flon Belt is separated from the Paleoproterozoic Thompson Nickel Belt by Kisseynew Domain rocks. To the west, the Flin Flon Belt is terminated against the Tabernor Fault Zone (Figure 7.1.1).

The Geological Survey of Canada (GSC) - Manitoba-Saskatchewan NATMAP Shield Margin Project and LITHOPROBE Trans-Hudson Orogen transect built on an extensive existing geological database that led to a much-improved understanding of the components and evolution of the southeastern Reindeer Zone, including the Flin Flon Belt (e.g., Lucas *et al*., 1996). These investigations have shown that, at a crustal scale, the Flin Flon "greenstone" belt is only one of three components in a northeast-dipping stack, juxtaposed during 1.84 to 1.80 Ga collisional deformation:

1. At the lowest structural level (exposed in the Pelican Window, Figure 7.1.2, within the Hanson Lake Block): metaplutonic rocks and paragneisses (3.20-2.40 Ga) of the "Sask Craton" (Corrigan *et al*., 2007).

2. At intermediate structural levels: Flin Flon Belt (now defined to include the Attitti Block and Paleoproterozoic rocks in the Hanson Lake Block) and Glennie Domain, shown in Figures 7.1.1 and 7.1.2, (together comprising the "Flin Flon-Glennie Complex" [FFGC]; Lucas *et al*., 1996).
3. At the highest structural levels: marine turbidites (Burntwood Group; 1.85-1.84 Ga) and distal facies of alluvial-fluvial sandstones (Missi Group) in the Kisseynew Domain (Figure 7.1.1).

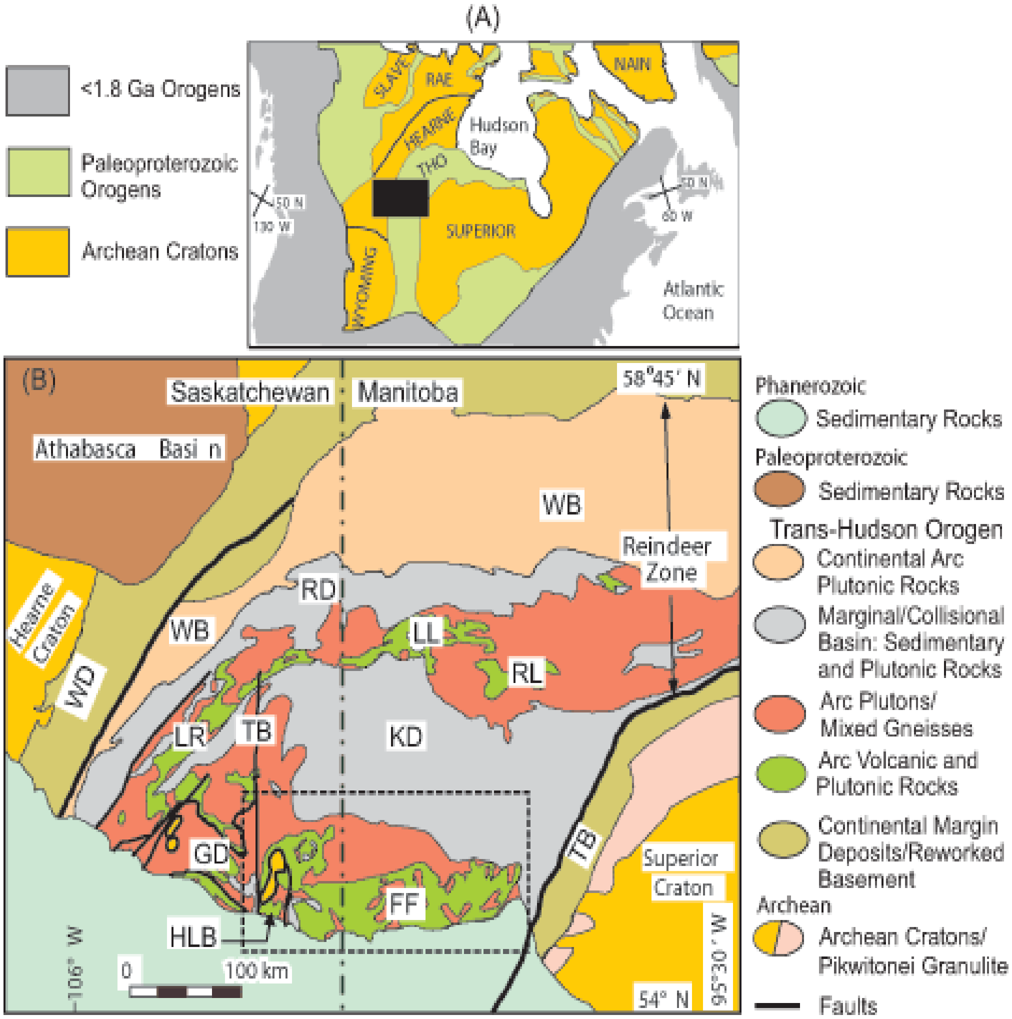


Figure 7.1.1: Location map after Hoffman (1989) illustrating the position of the Flin Flon-Snow Lake Belt in relation to the Precambrian geology of North America (A) and central Canada (B). The latter illustrates the components of the dominantly juvenile core to the Trans-Hudson Orogen in relation to the bordering Archean terranes, middle Proterozoic Athabasca Basin and overlying Phanerozoic strata of the Western Canada Sedimentary Basin. FFB: Flin Flon Belt; GD: Glennie Domain; HLB: Hanson Lake Block: LL-LT: Lynn Lake-LaRonge Belt; KD: Kisseynew Domain; RD: Rottenstone Domain; RL: Rusty Lake Belt; TB: Thompson Belt; TB: Tabernor Fault Zone; WB: Wathaman-Chipewayan Batholith; WD: Wollaston Domain. Modified from Lucas *et al.* (1996) from the original by Hoffman (1989). Dashed box represents area shown in more detail in Figure 7.1.2.

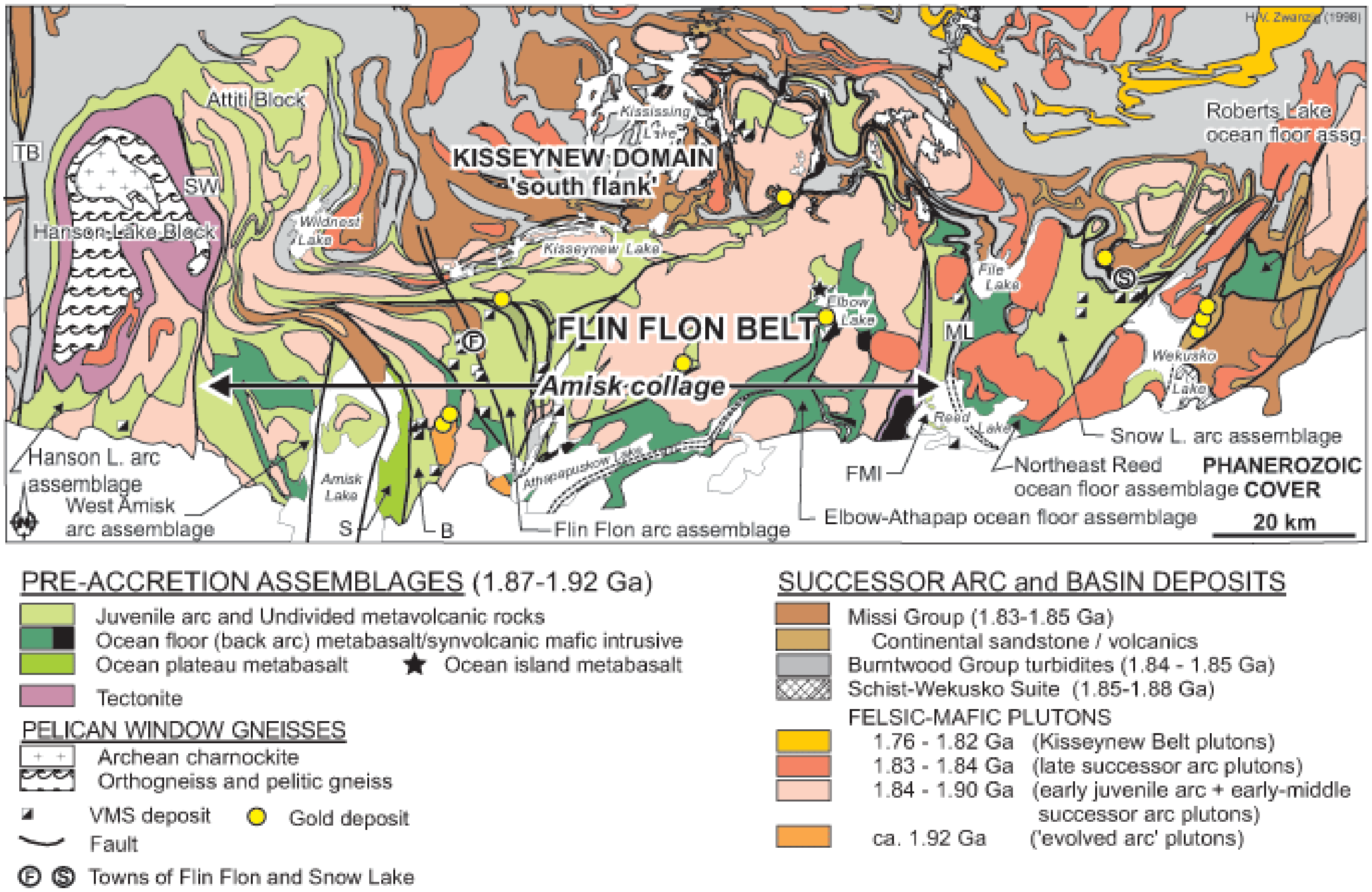


Figure 7.1.2: Map of the Flin Flon-Snow Lake Belt, illustrating the tectonio-stratigraphic assemblages, the location of the various accretionary assemblages, and major mineral deposits. B: Birch Lake assemblage; FMI: Fourmile Island assemblage; ML: Morton Lake fault zone; S: Sandy Bay assemblage; TB: Tabernor Fault Zone; SW: Sturgeon-Weir fault zone. Modified from Zwanzig (1999) and Lucas *et al.* (1996). Zoro Lithium Project is marked on Figure 7.1.3.

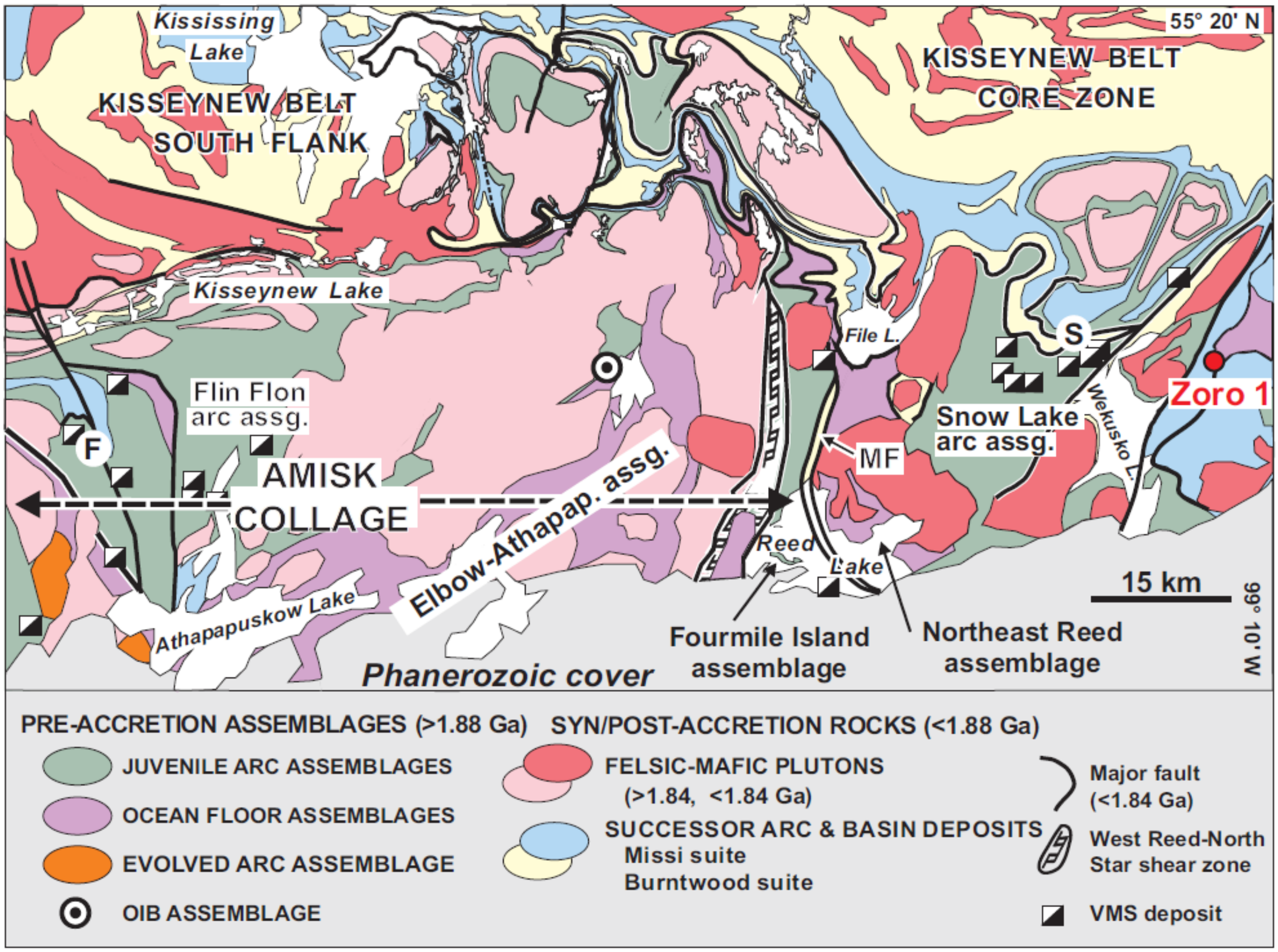


Figure 7.1.3: Regional geology map of the Flin Flon-Snow lake greenstone belt emphasizing pre- and syn/post-accretion rocks.

7.2 GEOLOGICAL EVOLUTION AND COMPONENT OF THE FLIN FLON SNOW LAKE GREENSTONE BELT

The stratigraphy of the Flin Flon-Snow Lake Belt has been previously subdivided into two major groups, the Amisk Group metavolcanic rocks and Missi Group continental metasedimentary rocks (Bruce, 1918; Harrison, 1951). The Flin Flon Belt is now recognized as consisting of several 1.9 to 1.88 Ga terranes comprised of four main tectono-stratigraphic assemblages that represent both juvenile and continentally underlying oceanic segments of a Paleoproterozoic ocean basin that were accreted during formation of the Trans- Hudson orogen (Syme, 1990;

Syme and Bailes, 1993; Stern *et al*., 1995a, b; Lucas *et al*., 1999). The orogen was formed by oblique collision between the Superior and Hearne Archean terranes, and the resulting collage is separated into assemblages distinguished by unique tectono-stratigraphy and dismembered by fault systems that were originally thrust surfaces (Syme, 1995, Lucas *et al*., 1996). Each tectonostratigraphic assemblage is a distinct package of rocks in terms of its stratigraphy, geochemistry, isotopic signature, age, and inferred plate tectonic setting (see below; Syme and Bailes, 1993; Lucas *et al*., 1996; Corrigan *et al*., 2007).

The tectonostratigraphic assemblages were juxtaposed in an accretionary complex at ca. 1.88 to 1.87 Ga, probably as a result of arc-arc collision (D1; Lucas *et al*., 1996; Stern *et al*., 1999; Figure 7.1.3). Accretionary collage-bounding structures were largely obliterated by subsequent deformation and metamorphic events (D2-D5) but are inferred where juxtaposed terranes are "stitched" together by calc-alkaline plutons related to a 1.866 to 1.838 Ga successor arc formation (Whalen *et al*., 1999). Coeval subaerial volcanism is recorded in ca. 1.87 to 1.85 Ga calc-alkaline to shoshonitic volcaniclastic sequences (Syme, 1988; Bailes and Syme, 1989; Lucas *et al*., 1996; Stern *et al*., 1996). Unroofing of the accretionary collage, development of a paleosol, and deposition of alluvial-fluvial sedimentary rocks (Missi suite; Bailes and Syme, 1989; Holland *et al*., 1989) occurred ca. 1.85 to 1.84 Ga (Ansdell, 1993). These events were coeval with the waning stages of post-accretion arc magmatism (Stern and Lucas, 1994; Whalen and Hunt, 1994; Lucas *et al*., 1996). Development of the Kisseynew turbidite basin (now part of the Kisseynew Domain) was synchronous with continental sedimentation in the Flin Flon Belt (Ansdell, 1993; David *et al*., 1993, 1996; Machado and Zwanzig, 1995; Connors, 1996; Connors *et al*., 1999).

The transition from Kisseynew basin extension to collisional collapse occurred rapidly at about 1.840 Ga, although sedimentation and magmatism continued through to ~1.830 Ga (Ansdell and Norman, 1995; Machado and Zwanzig, 1995; David *et al*., 1996; Connors *et al*., 1999). The Kisseynew Domain was thrust over the Amisk collage along the southern flank of the Kisseynew Domain (Harrison, 1951; Zwanzig, 1990; Lucas *et al*., 1994; Connors, 1996; Connors *et al*., 1999; Zwanzig, 1999). Following collisional thickening and peak metamorphism at 1.83 to 1.80 Ga, the Flin Flon Belt experienced protracted intracontinental deformation to ca. 1.69 Ga (Lucas *et al*., 1996; Stern *et al*., 1999).

The Flin Flon-Snow Lake Belt consists of two principal segments (Amisk collage and Snow Lake arc assemblage) that were juxtaposed during southwest-verging continent-continent collision between 1.84 and 1.82 Ga. To the west of the Amisk collage is a volcano-sedimentary domain of similar age (Hanson Lake arc assemblage; Figure 7.1.2). Although not traditionally designated as part of the Flin Flon Belt, the Hanson Lake arc assemblage will be included here due to the presence of volcano-sedimentary rock units and several VMS occurrences and deposits of similar age (Maxeiner *et al*., 1993, 1999).

Although each of these three tectonic segments has a distinct character, their metallogenic tenor is a product of the same three tectonically controlled evolutionary stages of the region. The first is represented by syngenetic polymetallic base metal and precious metal deposits and occurrences that formed during a pre-accretionary stage within distinct oceanic, supra-subduction environments. The second consists of post-accretion intrusion-related mineralization associated with successor arc formation and extensional magmatism. The third is comprised of shear zone-related orogenic Au deposits that formed during periods of collision, oblique compression, and crustal thickening.

The Amisk collage is comprised of a series of fault-bounded tectonostratigraphic assemblages (Syme, 1995; Lucas *et al*., 1996; Figure 7.1.2). These are intruded by post-accretionary plutons and are overlain by fluvial-alluvial sedimentary rocks of the Missi Group. The collage is bounded to the west by the Sturgeon-Weir fault system, to the east by the Morton Lake fault zone, and to the north by the southern flank of the Kisseynew domain. It extends to the south below the Phanerozoic cover. The Amisk collage contains the West Amisk, Birch Lake, Flin Flon, and Fourmile Island oceanic arc assemblages, and the Sandy Bay and Elbow-Athapapuskow back-arc basin basalt assemblages (Stern *et al*., 1999; Syme *et al*., 1999; Figure 7.1.2).

The arc and ocean floor assemblages in the eastern part of the Flin Flon Belt are collectively sufficiently distinct from the Amisk collage arc assemblages to suggest that they represent remnants of unrelated arc terranes (Lucas *et al*., 1996; Syme *et al*., 1996). This eastern part of the belt is characterized by a number of allochthons in a thrust stack that is bordered to the west by the Morton Lake fault zone and to the east and north by the overthrust Kisseynew Domain (Bailes *et al*., 1994; Syme, 1995). These allochthons are comprised of the Snow Lake arc and the Northeast Reed and Roberts Lake ocean floor assemblages (Figure 7.1.2) that are separated by major bounding fault systems. The 1.89 Ga Snow Lake arc assemblage (David *et al*., 1996;

Bailes and Galley, 1999) is the only one that contains significant VMS mineralization. It is exposed in a thrust stack that includes several structurally imbricated slivers of 1.84 to 1.83 Ga post-accretion sedimentary strata of the Burntwood suite (Stern *et al*., 1995a; Connors *et al*., 1999). The >6 km-thick dominantly juvenile oceanic to crustally contaminated arc succession of the Snow Lake arc assemblage consists of three conformable volcanic successions that record the evolution from nascent or primitive arc through mature arc, to rifting and opening of a back-arc basin.

The Hanson Lake arc assemblage structurally overlies Archean crust and contains coeval volcanic and sedimentary assemblages that are found within the neighbouring Flin Flon Belt. It is also host to numerous VMS deposits and occurrences. The fault-bounded area containing the Hanson Lake arc assemblage and underlying Archean crust is known as the Hanson Lake Block (HLB) (Figures 7.1.1 and 7.1.2).

The HLB is composed of a highly deformed and metamorphosed assemblage of 1.91 to 1.85 Ga volcanic and sedimentary rock, and 1.86 to 1.81 Ga syntectonic intrusions and migmatitic gneisses that have been thrust over the ca. 2.5 Ga Neoarchean charnokitic and enderbitic intrusive rocks known as the Pelican Window (surface expression of the Sask Craton; Ashton *et al*., 1987; Maxeiner *et al*., 1993, 1999; Ashton and Lewry, 1994; Figure 7.1.2). The HLB is terminated to the west by the Tabernor fault zone and to the east by the Sturgeon-Weir fault zone, which separates the HLB from those assemblages traditionally included within the Flin Flon Belt and extends southward below the Phanerozoic cover. The principal reason for the original exclusion of the arc assemblages of the HLB from those of the Flin Flon Belt was that they were originally believed to have formed upon Archean crust and were, therefore, not considered part of the Flin Flon oceanic supra-subduction suite.

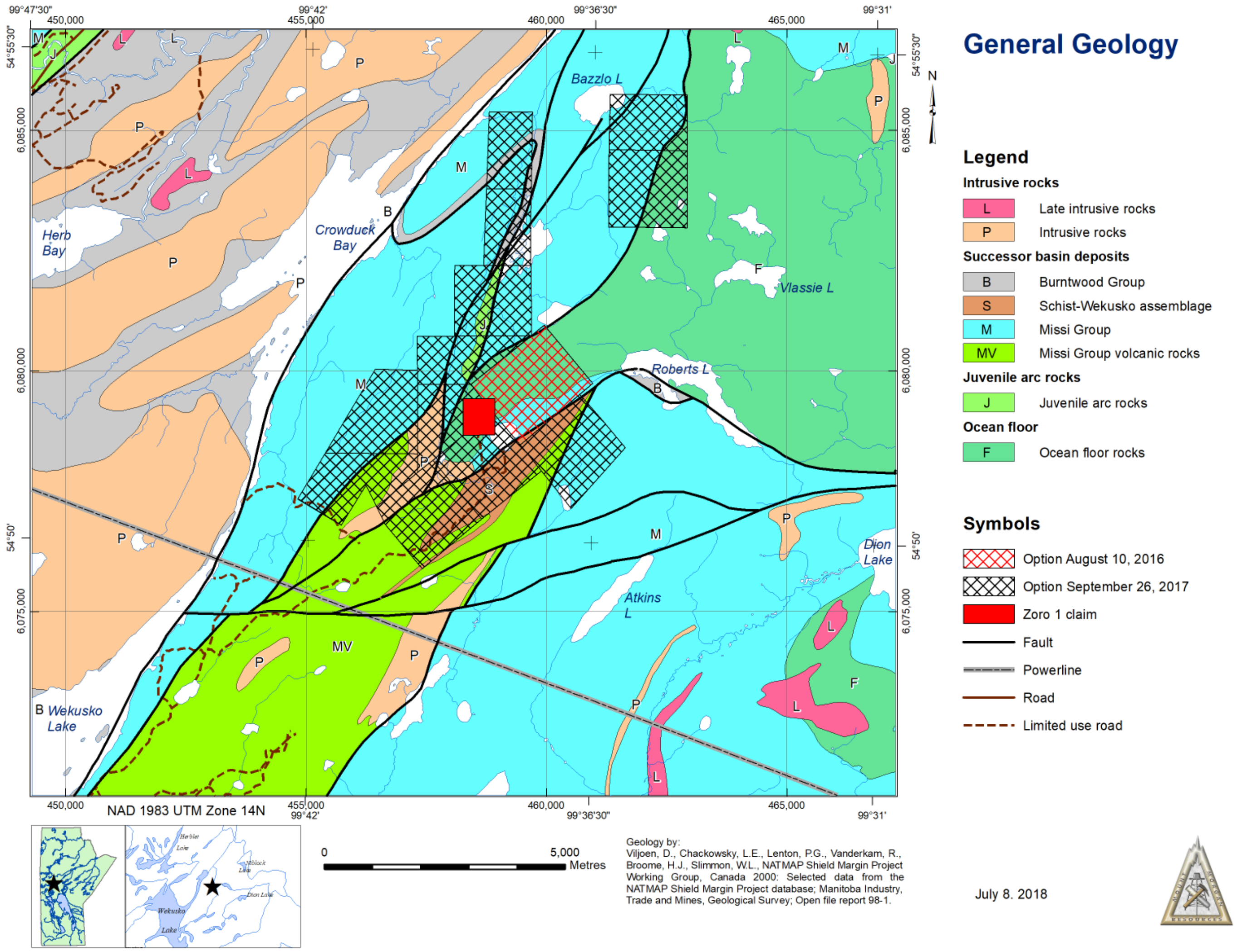
Supracrustal rocks of the HLB are dominated by metavolcanic and metasedimentary rocks. Volcanism and sedimentation are coeval from 1910 to 1880 Ma, with sedimentation continuing to at least 1850 Ma. Volcanic strata are dominantly tholeiites and include pillowed basalt overlain by intermediate to felsic flows and volcaniclastic rocks. Also present is a large felsic hypabyssal intrusive/extrusive complex. The volcanic assemblage is in contact with calc-silicate-carbonate-rich strata, silicate-facies iron formation, and polymictic conglomerate, and overlain by psammitic greywacke and mafic wacke (Maxeiner *et al*., 1993, 1999).

The supracrustal assemblages of the HLB are intruded by numerous synvolcanic intrusions, ranging in composition from ultramafic through gabbro and quartz diorite to rhyolitic. Large antiformal domes of migmatitic gneiss are accompanied by lit-par-lit injection into the supracrustal formations. Metamorphic grade generally increases from south to north, from upper greenschist to upper amphibolite facies, with regional metamorphism peaking between 1810 and 1806 Ma. A major folding event took place between 1860 and 1850 Ma and was followed by 1810 to 1800 Ma continental collision that caused the thrusting of this terrain over Archean basement (Ashton and Lewry, 1994). Deformation that accompanied crustal thickening and post-peak metamorphism continued until 1770 Ma.

7.3 GENERAL AND DETAILED GEOLOGY

General and detailed geology for the Zoro Lithium Project is depicted in Figures 7.3.1 and 7.3.2. Mapping by the Manitoba Geological Survey on the property documents the Zoro Lithium Project is underlain by Ocean Floor volcanic rocks of the Roberts Lake allochthon and lesser amounts of Missi Group sedimentary rocks. The Ocean Floor rocks comprise mafic volcanic and related intrusions and the Missi Group consists of sandstone, siltstone, mudstone and quartzo-feldspathic gneiss and migmatite. These lithologies are flanked to the south by Missi Group calc-alkaline and tholeiitic basalt and rhyolite to dacite ash flow tuff and flows and to the east and west more Missi Group sedimentary rocks. The Ocean Floor mafic volcanic rocks adjacent to the dykes consist of a fine- to medium-grained strongly foliated dark green lithology. These andesitic to basaltic lithologies are locally interbedded with volcaniclastic sedimentary rocks and all are intruded by a quartz-phyric granite intrusion. The flows are generally fine- to medium-grained, massive with a 50°-70° lineation and strikes of N10°-30°E and steep northwest dips. Localized quartz veins, quartz laminae and associated iron carbonate veinlets are also present in outcrop adjacent to lineaments interpreted to represent faults. Minor arsenopyrite was noted in the quartz veins and laminae. These rocks are locally rusty-weathered and crosscut by veinlets of iron carbonate and quartz. Minor arsenopyrite and pyrite was observed in the quartz veins and laminae.

The pegmatite dykes generally strike northwest to north-northwest with steep dips and crosscut the regional foliation at a low angle. The dykes tend to be concentric in internal structure and the grain size of the constituent minerals (potassium feldspar, quartz, spodumene and black tourmaline) coarsens towards the center of the dykes. This pattern may be locally interrupted by patches of saccharoidal albite, large muscovite aggregates and coarse albite stringers with garnet and beryl. Spodumene is concentrated in the cores of the dykes. Some of the dykes have been split into sub-parallel veins by post-emplacement tectonic activity.



Geology by:
Viljoen, D., Chackowsky, L.E., Lenton, P.G., Vanderkam, R., Broome, H.J., Slimmon, W.L., NATMAP Shield Margin Project Working Group, Canada 2000: Selected data from the NATMAP Shield Margin Project database; Manitoba Industry, Trade and Mines, Geological Survey; Open file report 98-1.

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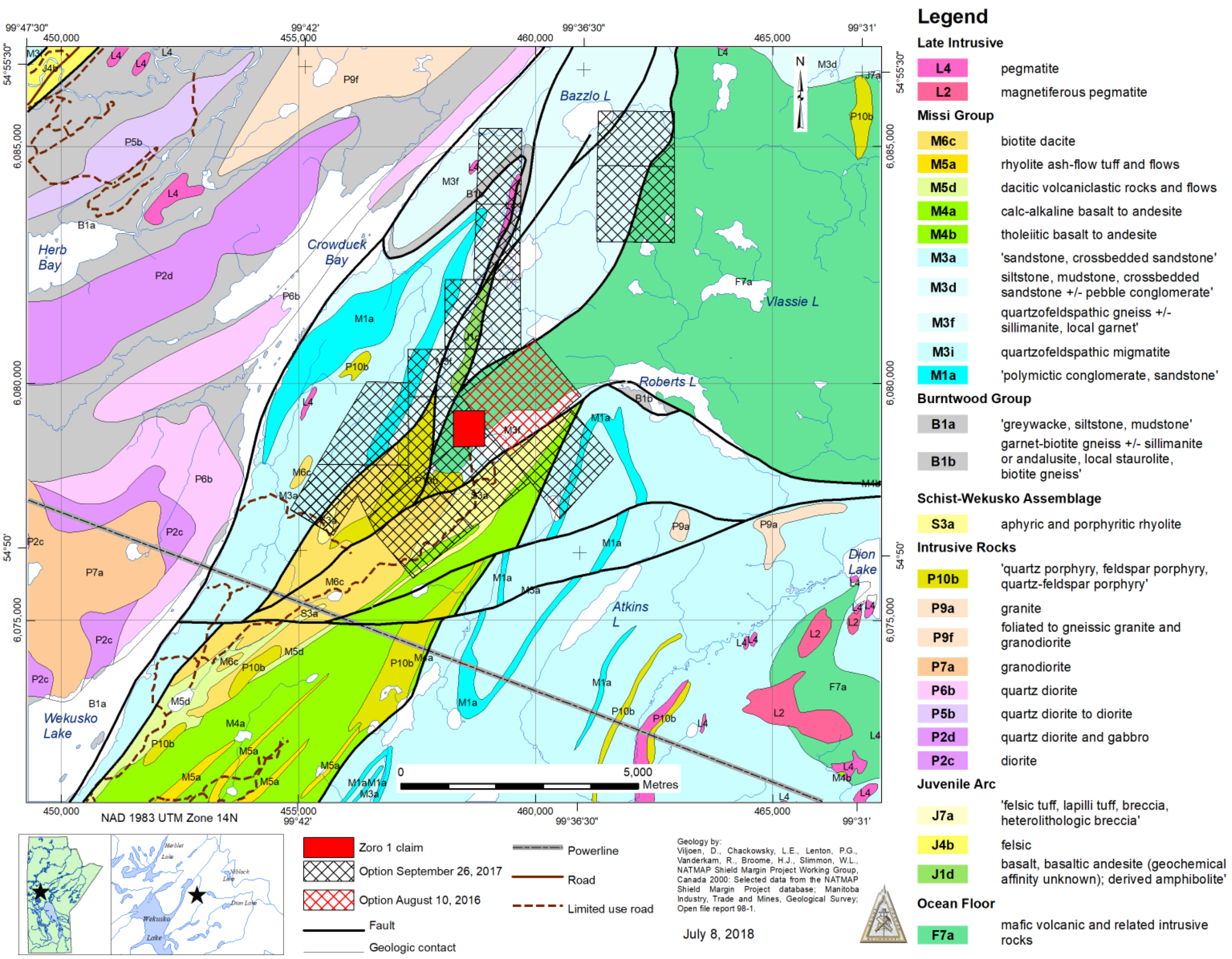
Figure 7.3.1: General geology in vicinity of the Zoro Lithium Project with claim boundaries.



Geology by:
Viljoen, D., Chackowsky, L.E., Lenton, P.G., Vanderkam, R., Broome, H.J., Slimmon, W.L., NATMAP Shield Margin Project Working Group, Canada 2000: Selected data from the NATMAP Shield Margin Project database; Manitoba Industry, Trade and Mines, Geological Survey; Open file report 98-1.

Figure 7.3.2: Detailed geology at the Zoro Lithium Project with claim boundaries.

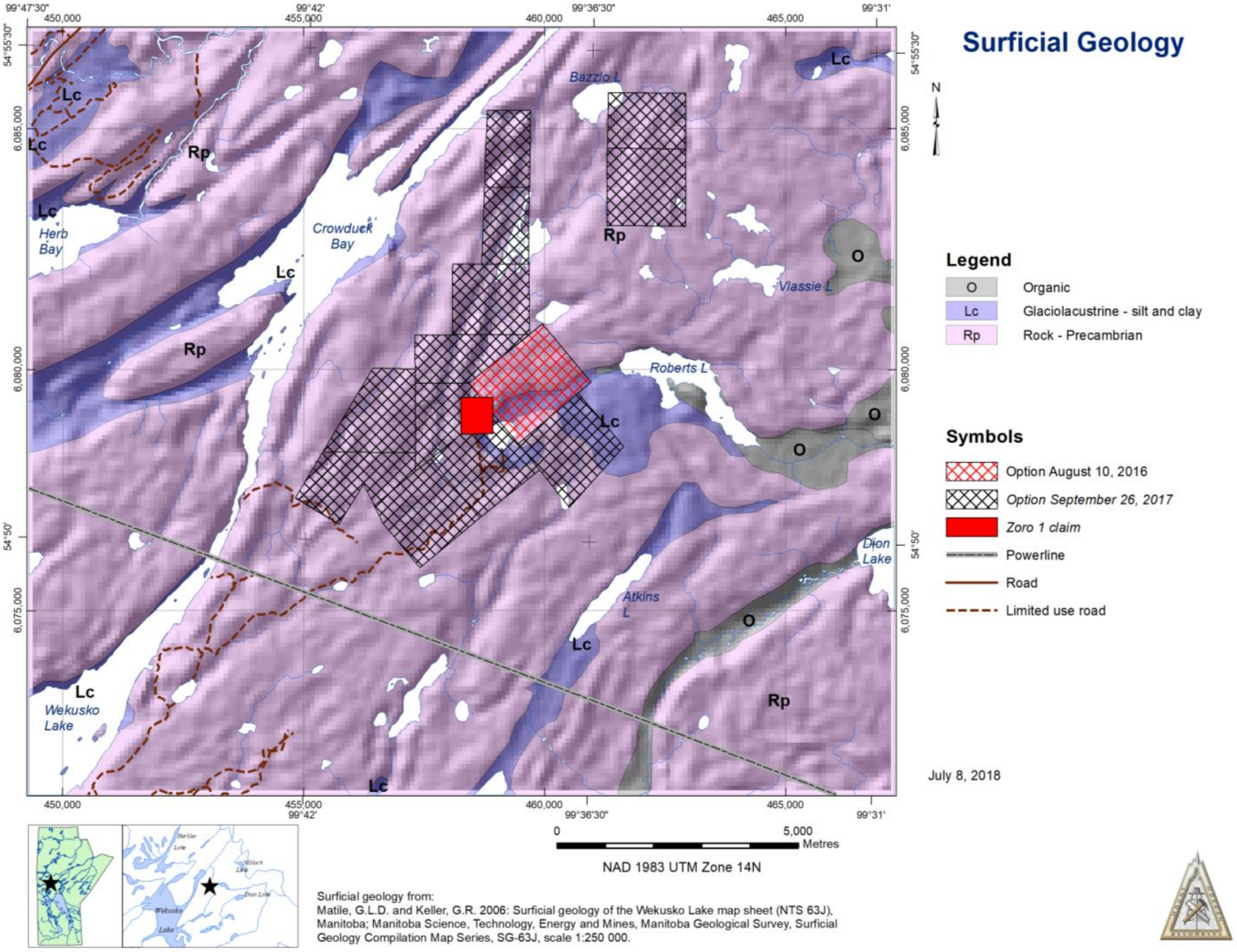
Detailed geologic observations of the pegmatites on the property were initially hampered by caved, filled and overgrown trenches (Figure 7.3.3) however subsequent to trench cleaning and mucking and the availability of drill core Martins *et al.* (2017) provide a detailed description of Dyke 1. It is a north-trending, near-vertical body that extends for at least 280 m along strike, with a maximum thickness of approximately 35 m. The presence of country-rock alteration was not noted in historical drill logs however, the mineral holmquistite ($Li_2(Mg, Fe^{2+})_3Al_2Si_8O_{22}(OH)_2$) was recently identified in the mafic volcanic host rock during field examinations, indicating metasomatic alteration associated with pegmatite intrusion. Rock and mineral analyses demonstrate that a broad metasomatic geochemical and mineralogical halo is present. The development of holmquistite-bearing assemblages is controlled by the introduction of Li into the country rock during pegmatite emplacement. These assemblages reflect greenschist-facies metamorphic conditions and are only found in amphibolitic wallrock, usually replacing hornblende, pyroxene or biotite (Heinrich, 1965; London, 1986). Based on historical and recent field and laboratory work zonation in the Dyke 1 pegmatite can be defined as follows:

1. the wall zone, composed predominantly of quartz, microcline and muscovite, with accessory tourmaline, hornblende, biotite and rare beryl and spodumene;
2. the intermediate zone, with medium-sized crystals of microcline, albite, quartz, muscovite and spodumene (<5%);
3. the central zone, with abundant spodumene (locally up to 50% but more commonly varying between 10% and 30%), albite, quartz and locally pollucite, and accessory apatite, tourmaline, pyrrhotite, lepidolite, columbite-group minerals and Fe-Mn–phosphate minerals;
4. the core zone, composed mainly of quartz with small- to medium-grained spodumene crystals (although locally 15–20 cm crystals of spodumene are observed) in a quartz matrix, with minor tourmaline and muscovite.

Figure 7.3.3: Overgrown and slumped trenches at the Zoro Lithium Project.

7.4 SURFICIAL GEOLOGY

Glacial sediments including till, glaciolacustrine and glaciofluvial deposits cover the project area (Figure 7.4.1). Most small streams flow primarily over organic deposits. Reworked glacial sediments form beaches on some lakes. Till and glacio-fluvial deposits associated with the Labrador and Keewatin sectors of the Laurentide Ice Sheet were developed approximately 115,000 years ago. During late Pleistocene, the ice sheet radiated southwest from Hudson Bay, and covered most of Manitoba. An extensive cover of carbonate-rich till derived from the James Bay and Hudson Bay Lowlands has been deposited over the area. As the ice sheet receded, glacial Lake Agassiz formed and resulted in the deposition of a veneer of glaciolacustrine silt and clay.



Surficial geology from:
Matile, G.L.D. and Keller, G.R. 2006: Surficial geology of the Wekusko Lake map sheet (NTS 63J), Manitoba; Manitoba Science, Technology, Energy and Mines, Manitoba Geological Survey, Surficial Geology Compilation Map Series, SG-63J, scale 1:250 000.

Figure 7.4.1: Surficial geology of the Zoro 1 property with claim boundaries.

7.5 GEOPHYSICAL CHARACTERISTICS OF THE ZORO LITHIUM PROPERTY

The Zoro Lithium Project occurs in a generally magnetically “active” zone in an area of low Total Magnetic Intensity. The low magnetic signature of the property area is flanked on all sides by large, very strong magnetic responses. The property encompasses two north-northeast-trending magnetic anomalies which are visible in both the Total Magnetic Intensity and Vertical Gradient maps presented in Figures 7.5.1 and 7.5.2, respectively. The more westerly of these anomalies extends the full length of the property and continues well past the northern property boundary. The eastern anomaly is confined within property boundaries but also extends the entire length of the property. Historic exploration has not fully explained these anomalies that are likely due, in part, to mineralized quartz-feldspar porphyry intrusions that occur within faults on the property and to the pegmatite dykes that trend both north and northwest.

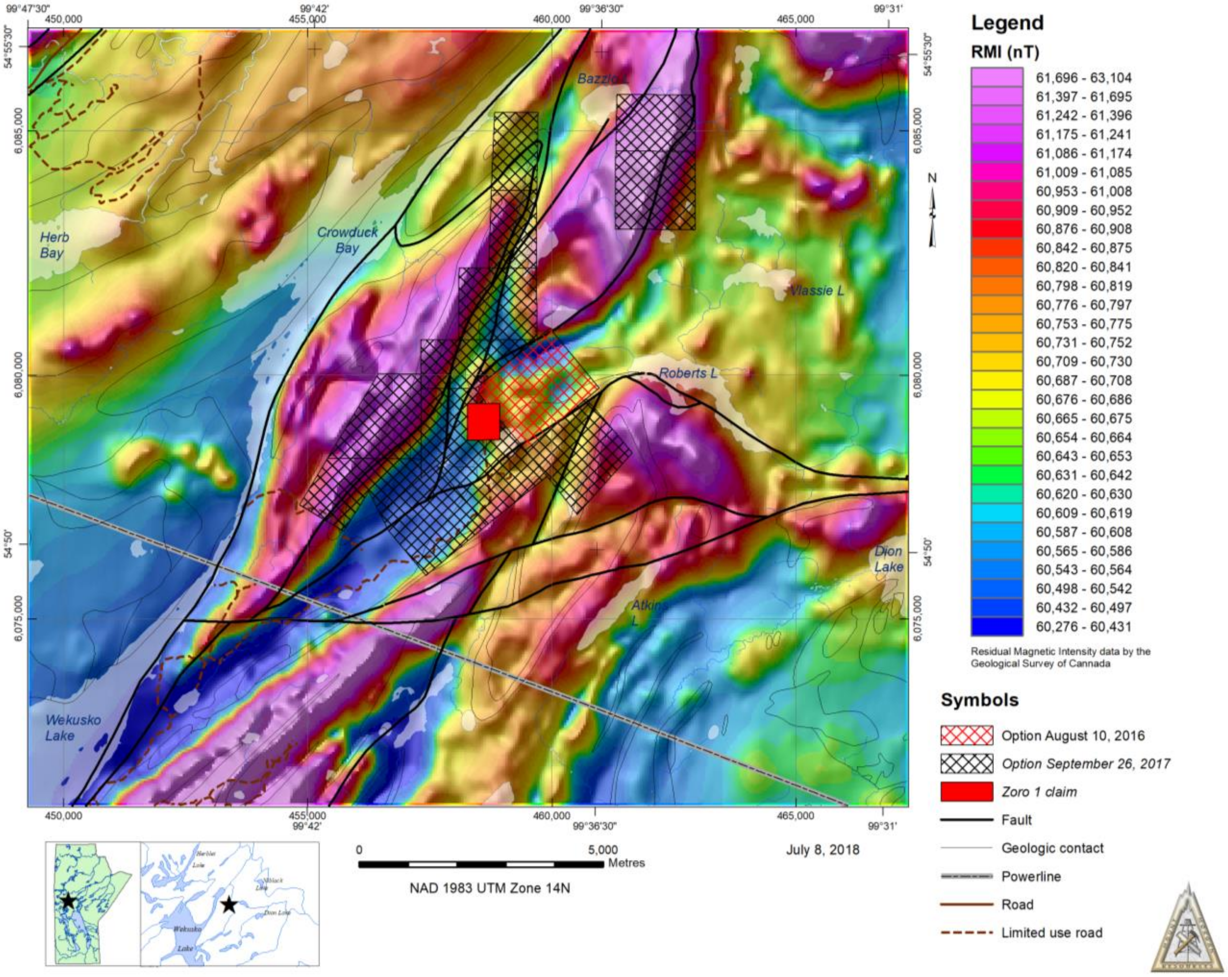


Figure 7.5.1: Regional airborne Residual Magnetic Intensity survey results of the Zoro Lithium Project, Wekusko Lake area with claim boundaries.

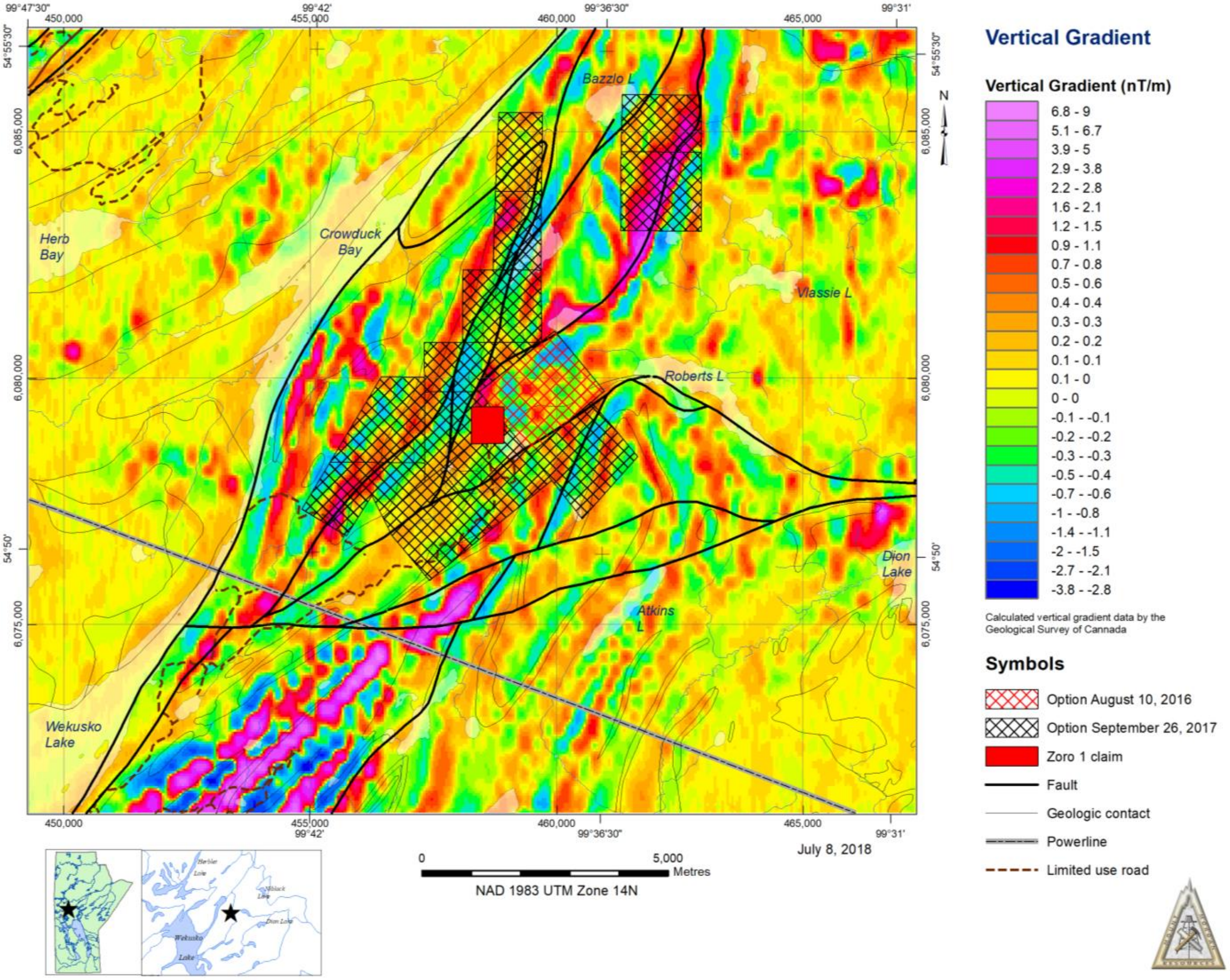


Figure 7.5.2: Regional airborne Vertical Gradient for the Zoro Lithium Project with claim boundaries.

7.6 MINERALIZATION

The major mineralizing events recognized in the Flin Flon belt took place during the three main stages of crustal development: pre-accretion, post-accretion, and continent- continent collision. The pre-accretionary stage is represented by syngenetic base metal and Au deposits. The syn- to post-accretionary stage is characterized by several examples of intrusion-hosted base and precious metal deposits, and the continental collision stage by the development of orogenic Au deposits and lithium-cesium-tantalum-enriched pegmatites. The primary mineralization of interest on the Zoro Lithium Project is spodumene which is a lithium aluminum silicate (8.0% Li_2O, 27.4% Al_2O_3, 64.6% SiO_2). Spodumene is a pegmatite mineral that has a glassy lustre and may be opaque; it is nearly white in the low-iron variety and dark green in iron-rich crystals. An example of coarse bladed spodumene in outcrop at the Zoro Lithium Project is given in Figure 7.6.1.

Figure 7.6.1: Coarse-grained bladed spodumene, trench muck sample, Dyke 1, Zoro Lithium Project.

The Zoro Lithium Project comprises a minimum of seven zoned pegmatite dykes that intrude Proterozoic Amisk Group volcanic and volcaniclastic rocks in a 2-km zone trending approximately 55° northwest (Mulligan, 1965 in Cerny *et al*., 1981; Fedikow *et. al*., 1993). The dykes strike north to northwest and dip vertically. Several have been described as gently dipping bodies (Bannatyne, 1985). The main, most westerly dyke or Dyke 1 outcrops along the west side of a ridge, 4.5 to 6 m high, and intrudes siliceous metasedimentary rocks and amphibolite (Bannatyne, 1985). It is up to 27 m (90 ft.) wide at surface and is exposed in 16 historical cross-trenches over a length of 183 m. Based on Far Resources drilling results, lithium mineralization has been defined for 265m along strike, up to 40m wide and to a depth of 265m, Individual dykes have lengths of approximately 244 m. The outer zones of the pegmatite dykes contain pink aplite and coarse feldspar, locally green muscovite, tourmaline, and occassionally beryl. Spodumene, quartz, cleavelandite, and tourmaline form core zones with interstitial coarse feldspar. Spodumene is usually coarse-grained and is sometimes altered. It is most prevalent in the central 9 m (30 ft.) of the main dyke. In this dyke, spodumene crystals (up to 35 cm long) occur either in clusters, over widths of 6 m or more, or associated with coarse tourmaline and perthite megacrysts; some spodumene crystals show a preferred orientation of 45° to 55° (Bannatyne, 1985). One of two parallel dykes south of the main outcrop, is 5 m wide, and contains spodumene crystals in pods (up to 33 cm across). In other dykes, coarse grained spodumene is abundant in lenticular bands and fine-grained spodumene is distributed through aplitic patches (Bannatyne, 1985). Beryl occurs as white, anhedral to subhedral crystals less than 1 inch (2.5 cm) in diameter in three of the seven dykes. Chemical analyses for selected minerals from the Zoro Lithium Project dykes are presented in Table 7.6.1 and discussed later in this report. Columbite-tantalite and sparse minute grains of pyrite and chalcopyrite were found in thin sections (Green Bay Mining & Exploration Ltd., Corporation File).

Table 7.6.1: The compositional characteristics of selected minerals from Dyke 1, Zoro lithium property (Cerny, 1981).

Blocky Potassium Feldspar (n=19 samples)	**Arithmetic Mean (weight %) Standard Deviation**			**Rb**	**K/Rb**	**Cs**
				0.392	33.7	0.0603
				+/-0.144	+/-15.1	+/-0.0383
Range:				0.183-0.668	17.1-62.8	0.0158-0.1404
Core Muscovite (n=4)	**Li**	**Rb**	**K/Rb**	**Cs**	**Be**	
	0.171	0.792	12.5	0.0702	0.0021	
	+/-0.122	+/-0.432	+/-5.1	+/-0.0497	+/-0.0003	
	0.025-0.460	0.473-1.42	5.9-17.7	0.025-0.123	0.0017-0.0023	
Late Beryl (n=9)	**Li**	**Na**	**Na/Li**	**Cs**		
	0.331	0.903	2.83	0.939		
	+/-0.080	+/-0.101	+/-0.50	+/-0.452		
	0.202-0.460	0.720-1.01	2.00-3.56	0.334-1.67		
Spodumene (n=3)	**Na_2O**	**Fe as Fe_2O_3**				
	0.230	0.943				
	+-0.104	+/-0.446				
	0.16-0.35	0.428-1.21				
Garnet (n=1)	**FeO**	**MnO**	**MnOx100/MnO+FeO**		**CaO**	**MgO**
	24.80	17.10	40.81		0.46	0.49

Utilizing historical descriptions and recent preliminary petrographic work, Martins *et al.* (2017) were able to distinguish at least three different stages of spodumene growth in the dykes. These included greenish spodumene with characteristics typical of a primary phase; spodumene-quartz intergrowths, possibly after petalite breakdown (Cerny and Ferguson, 1972); and late bands of very fine grained spodumene that crosscut other mineral phases or surround feldspar and muscovite grains. Locally, spodumene crystals are surrounded by fine-grained mica, possibly Li-mica or lepidolite. This could be indicative of a late Li-enriched fluid episode such as autometasomatism that could have produced late Li-enriched mica. Acicular opaque minerals of the columbite group are present, and late bands of fluorite occur locally in fractures. The latest event, identified in thin section, produced Fe-rich, quartz-calcite stringers with no preferred orientation crosscutting the pegmatite. In thin section, feldspar and muscovite show evidence of deformation such as kink bands in muscovite, suggesting that pegmatite emplacement occurred prior to the latest stages of regional deformation.

7.7 MINERAL CHEMISTRY

Mineral-chemistry data for both muscovite and K-feldspar from Dyke 1 is similar to results reported by Cerny *et al*. (1981) for pegmatites from the Green Bay group of the Wekusko Lake pegmatite group. The full dataset of electron microprobe results can be found in (Martins and Linnen, 2017). Observations for muscovite and K-feldspar mineral chemistry are given in (Martins and Linnen, 2017) and are reviewed below.

7.7.1 MUSCOVITE

Two generations of mica were identified in Dyke 1 but only results for primary muscovite are described. Primary muscovite is identified by the presence of sharp grain boundaries, subhedral to euhedral shape, grain size comparable to other magmatic minerals, absence of reactions with other minerals, absence of alteration in surrounding minerals and relative abundance, are described.

The mica compositions are all close to the stoichiometric dioctahedral muscovite end-member within the expected values for spodumene-subtype pegmatite (e.g., Selway *et al*., 2005; Martins *et al*., 2012). With respect to major element variability, muscovite has minor variation in Si and Al content, and Fe contents vary between 0.60 and 4.70 wt. % FeO. Trace-element concentrations include: variable F from below detection limit to 1.53 wt. % F, Rb from 0.18 to 0.81 wt. % Rb_2O, and Cs from below detection limit to 0.36 wt. % Cs_2O. The K/Rb ratio values of muscovite in Dyke 1 vary between 10.99 and 28.73, comparable to moderately evolved pegmatites from

Ontario (Tindle *et al*., 2002; Selway *et al*., 2005) but higher than the highly evolved Tanco pegmatite, in which mica has ratio values varying from 2.9 to 10.6 (Cerny, 2005). These results indicate that the Dyke 1 pegmatite is less fractionated than the Tanco pegmatite. Moreover, the K/Cs ratio values of muscovite vary from 27.89 to 871.48 and corroborate the lower level of fractionation of Dyke 1 compared to Tanco, in which mica ratio values vary from 14 to 93 (Cerny, 2005).

7.7.2 K-FELDSPAR

Selected K-feldspar grains were initially classified as primary but petrography and backscattered imagery revealed the k-feldspar had likely been albitized and as such the analyzed grains might not be good indicators of high-temperature primary crystallization processes. The stoichiometry of the analyzed K-feldspar is slightly non-ideal, which is typical for K-feldspar in granitic pegmatites (Cerny *et al*., 2012; Brown *et al*., 2017). Major elements do not vary significantly however trace elements have significant variability. Rubidium varies from below detection limit (<LLD) to 0.70 wt. % Rb_2O and Cs varies from <LLD to 0.27 wt. % Cs_2O. The values obtained indicate a moderate level of fractionation relative to pegmatites from Ontario (Figure GS2017-5-7b; Tindle *et al*., 2002; Selway *et al*., 2005). The K/Rb ratio values vary from 13.45 to 43.92, higher than the values listed for Tanco feldspar (4.0 to 14.2; Cerny, 2005) but typical for spodumene-type pegmatites in Ontario (Tindle *et al*., 2002). The K/Cs ratio values of K-feldspar from Dyke 1 vary from 48.26 to 584.62, well above the values reported for the Tanco pegmatite (6 to 26; Cerny, 2005), corroborating the lower degree of fractionation of Dyke 1 (Figure 7.7.1).

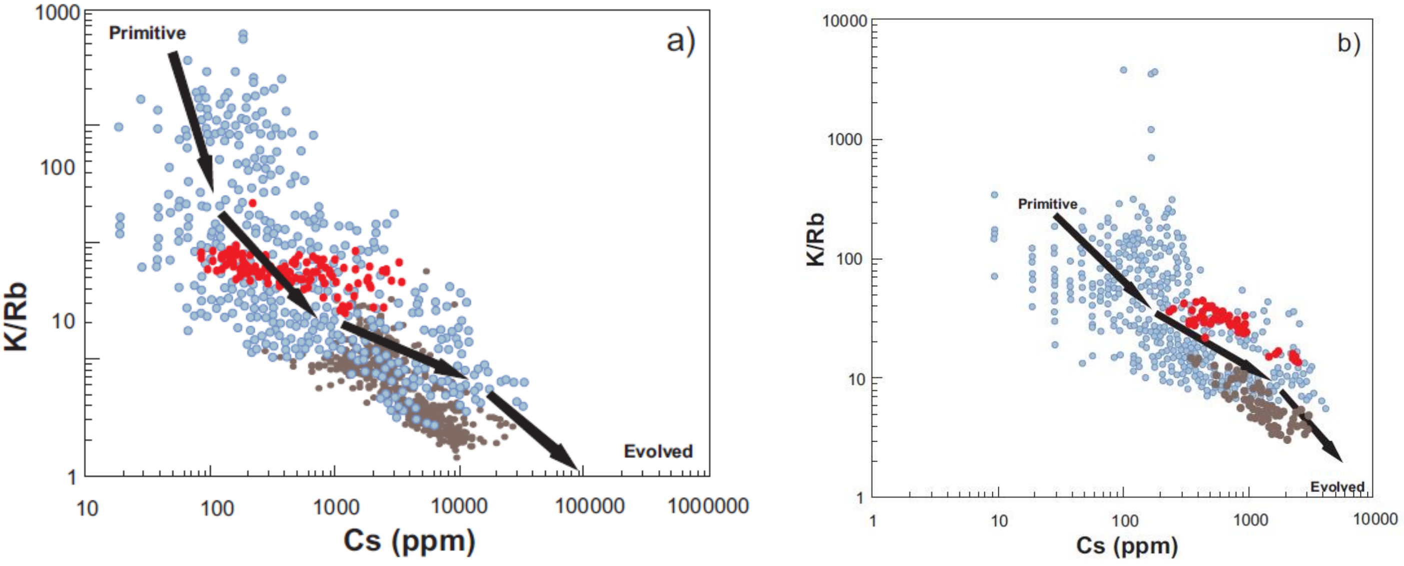


Figure 7.7.1: Mineral-chemistry results for muscovite and K-feldspar from Dyke 1 (From Martins *et al.,* 2017): a) general fractionation trend (arrows) for micas from Ontario pegmatites (blue dots; data from Selway et al., 2005) and those from the Tanco pegmatite (brown dots; S. Margison, unpublished data) and Dyke 1 (red dots); b) general fractionation trend (arrows) for K-feldspar from Ontario pegmatites (blue dots; Selway *et al*., 2005), Tanco pegmatite (brown dots; data from Brown, 2001) and Dyke 1 (red dots).

7.8 WHOLE ROCK GEOCHEMISTRY

A whole rock geochemical study based on three drill holes from the Dyke 1 pegmatite was recently undertake (Martins *et al.* (2017)). Sixty-nine samples of mafic-volcanic country rock and one sample of biotite-garnet-muscovite schist were collected from the hangingwall and footwall of the pegmatite dyke. Sample spacing was 5 m close to the contact with the pegmatite, and 10 m and 20 m apart farther away from the contact. The samples consisted of about 20 cm of split drill core from drill holes FAR16-001, FAR16-005 and FAR17-010 (Figure 7.8.1). Analyses were performed by Activation Laboratories (Ancaster, Ontario) using a sodium-pyrophosphate fusion technique, followed by multi-element ICP-MS. Selected samples of muscovite and K-feldspar from Dyke 1 were analyzed using a JEOL JXA-8530F field-emission electron microprobe at Western University. Analytical details are provided in DRI2017004 (Martins and Linnen, 2017).

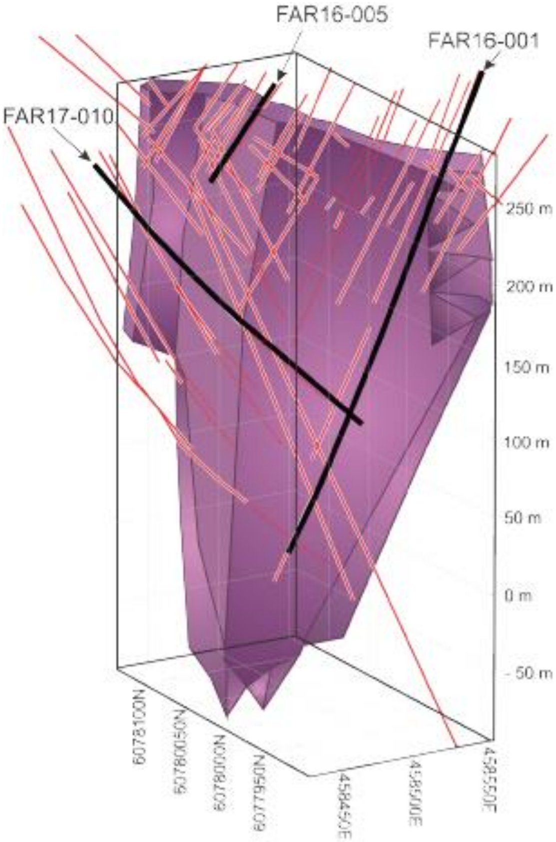


Figure 7.8.1: Location of drill holes FAR16-001, FAR16-005 and FAR17-010 used in the whole rock geochemistry study, in relation to Dyke 1.

7.8.1 BACKGROUND VALUES

Values for Li in Dyke 1 country rock are generally not available in the literature because this element is not routinely analyzed. Lithium is a moderately incompatible trace element in magmatic systems and its abundance in the mantle is estimated to be about 1.9 ppm (Ryan and Langmuir, 1987). These authors reported the world-wide range in Li content for mid-ocean ridge basalt (MORB) is 3–17 ppm and >8 ppm Li for evolved Fe-Ti basalt; andesite and dacite from the East Pacific Rise contain up to 30 ppm Li, indicating that Li increases with differentiation. Given that Dyke 1 country rock has flat rare earth element (REE) profiles characteristic of MORB (not plotted; data from Gilbert and Bailes, 2005b), the assumption for this study is that the background concentrations of Li should be low (<8 ppm) in non-metasomatized country rock to Dyke 1. Values for Rb and Cs are more readily available in the literature. Samples from an equivalent unit to the Dyke 1 country rock at south Wekusko Lake contain <7 ppm Rb and <0.03 ppm Cs (Gilbert and Bailes, 2005b). Thus, based on available data, values >6 ppm Rb and >0.02 ppm Cs are considered anomalous (twice the values of the standard deviation of data from Gilbert and Bailes, 2005b). For Li background, values are interpreted as anomalous at >16 ppm (double the maximum value for non-evolved MORB defined by Ryan and Langmuir, 1987).

In the absence of available Li analyses from the host rocks to the Dyke 1 pegmatite previous work conducted by Linnen *et al*. (2009, 2015) have been used for comparison with results obtained for the host rocks to Dyke 1. For the country rock to the Dyke 1 pegmatite, some of the highest values attained are 1900 ppm Li, 196 ppm Rb and 225 ppm Cs adjacent to the upper contact of the pegmatite.

7.8.2 RESULTS

The Rb and Cs values for Dyke 1 host rocks are well above what is reported for non-metasomatized ocean-floor mafic volcanic rocks from the same area (Gilbert and Bailes, 2005b). They are comparable to values obtained by Linnen *et al*. (2009) in country rock at the upper contact of the Dibs pegmatite (southeast Manitoba) and include values up to 2256 ppm Li, 184.5 ppm and 72.4 ppm Cs. For the Dibs pegmatite, values of Li, Rb and Cs in the country rock increase substantially toward the contact of the pegmatite (Linnen *et al*., 2009, 2015). For Dyke 1, the maximum concentrations for each element occur primarily in the country rock adjacent to the pegmatite contacts although the increase in concentration approaching the contact is not consistent. Within the same drillhole (FAR17-010), values at 11 m for Li, Rb and Cs are 48, 39.1 and 1.1 ppm, respectively (all values above background; Figure 7.8.2.1). These values close to surface are higher than at roughly 70 m downhole (14 ppm Li, 1 ppm Rb, 0.2 ppm Cs). Below 70 m there is a steady increase of Li, Rb and Cs as the contact with the pegmatite is approached at 163 m with values of 922 ppm Li, 51 ppm Rb, 23.9 ppm Cs; (Figure 7.8.2). This Li, Rb and Cs halo is likely related to the presence of fractures, the size or shape of the pegmatite and consequently the metasomatic halo, and the location of the Li, Rb or Cs mineralization within the pegmatite (i.e. the zonation). The Li, Rb and Cs halo can be measured up to 150 m away from the pegmatite contact. Elements such as Nb and Ta are low and are not enriched at the contacts with the pegmatite (Nb <5 ppm; Ta <2 ppm, with only one analysis as high as 8 ppm). Values for Sn are usually <4 ppm, with a few higher values of up to 91 ppm that occur close to the contact with the pegmatite and values for Tl are usually below detection limit but locally vary up to 6 ppm. High concentrations of As up to 6450 ppm are present in the country rock of Dyke 1 and may be related to processes responsible for the formation of gold mineralization known to occur in the area (Galley *et al*., 1989).

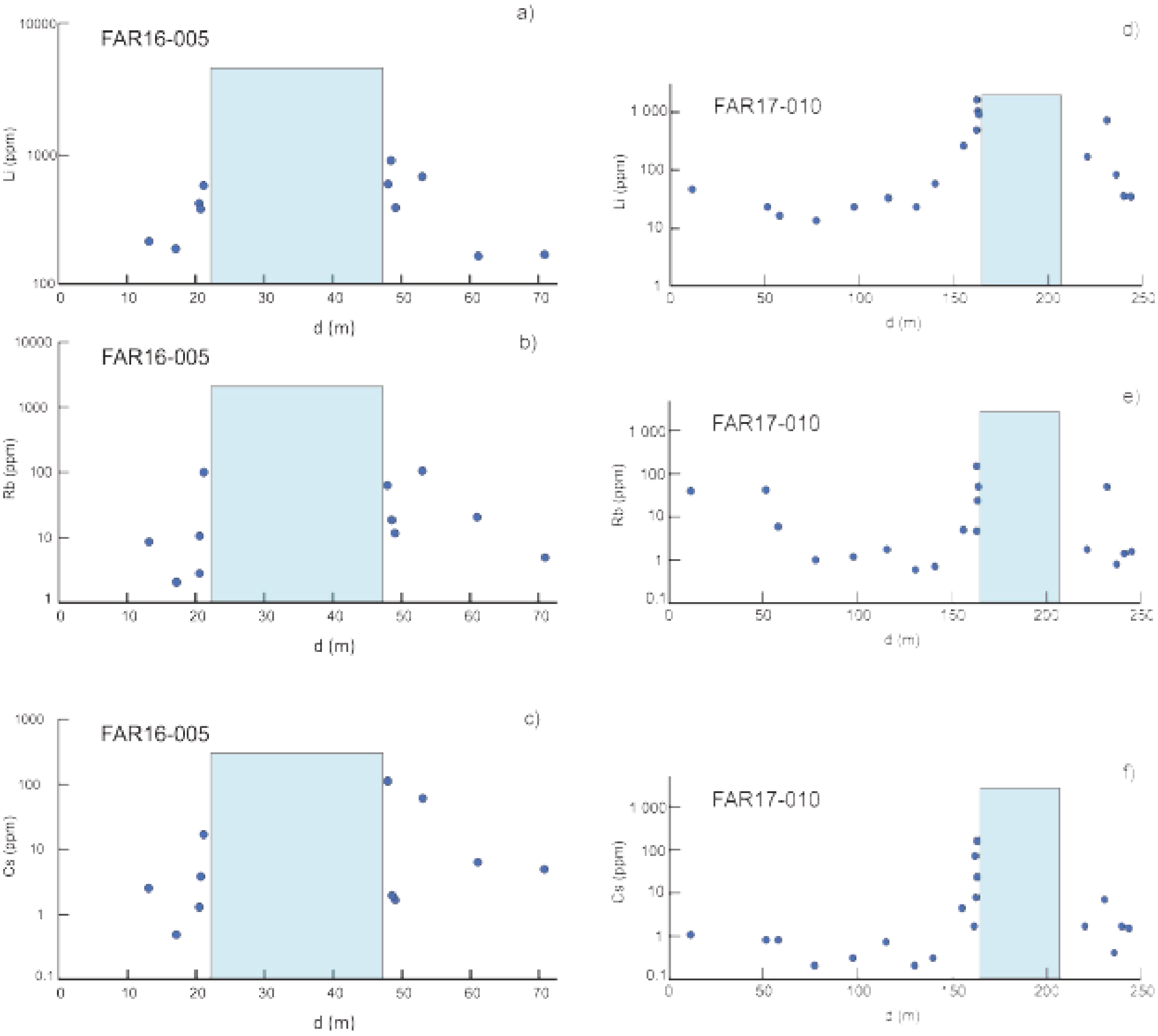


Figure 7.8.2: Element distribution diagrams showing variations along the length of the studied drill holes from Dyke 1: a) Li for drillhole FAR16-005; b) Rb for drillhole FAR16-005; c) Cs for drillhole FAR16-005; d) Li for drillhole FAR17-010; e) Rb for drillhole FAR17-010; f) Cs for drillhole FAR17-010. Shaded areas mark the location of the pegmatite. Plots from Martins *et al.* (2017).

7.9 OTHER MINERALIZATION

During historic and recent diamond drill testing of the Zoro 1 Dyke 1 pegmatite disseminated pyrrhotite, chalcopyrite and arsenopyrite were observed in andesitic wall rocks adjacent to pegmatite and in isolated zones throughout the host andesite, usually associated with quartz-carbonate veins. Green Bay Mining and Exploration presented assay results from the 1476'-1485' segment of drill hole 22A that indicated approximately 0.17 ounces per ton gold. Assay results are summarized in Table 7.9.1 and indicate that uniformly low values were documented from similar mineralization observed in Far Resources drill core. Representative samples of this core were sampled occasionally and assayed for gold during their drill programs.

Table 7.9.1: Summary of historic gold assay results from drill core at Dyke 1. Mineralization consisted of fine-grained, disseminated pyrrhotite, chalcopyrite and arsenopyrite.

DDH#	Sample Width (feet)	Gold Assay (opt or ounces per ton)
7	3	Trace
7	0.9	Trace
8	2.2	Trace
8A	1.5	0.01
16	5.5	0.01
16	2	0.02
16	3	0.02
16	4.2	0.02
16	2.6	0.01
24	9	0.03
24	3.9	0.01

Recent assays by Far Resources of intervals of drill core with disseminated arsenopyrite +/- pyrite and chalcopyrite within a pervasive front of silicification and quartz-carbonate +/- feldspar veinlets document the presence of gold in the core. The abundance of gold in these intervals is typified by the results for drill holes FAR17-015 and -019 in Table 7.9.2.

Table 7.9.2: Typical recent gold assay results from DDH FAR17-015 and -019.

Drill Hole #	Sample	From To (m)	Width (m)	Fire Assay (AAS) Parts per billion
DDH FAR17-015	705458	26.95-27.45	0.5	32
	705459	27.45-28.45	1	486
	705460	28.45-29.0	0.55	168
	705461	29.0-30.0	1	130
	705462	30.0-31.0	1	94
	705463	31.0-31.64	0.64	34
	705464	31.64-32.14	0.5	34
DDH FAR17-019	762268	9.6-10.55	0.95	367
	762269	10.55-11.3	0.75	153
	762271	11.3-12.0	0.7	24
	762272	12.0-13.0	1	119
	762273	13.0-14.0	1	68
	762274	14.0-15.0	1	233
	762275	15.0-16.0	1	51
	762276	16.0-17.0	1	38
	762277	17.0-18.0	1	76
	762278	18.0-19.0	1	101
	762279	19.0-19.5	0.5	110

8.0 DEPOSIT TYPES

8.1 RARE ELEMENT LITHIUM-CESIUM-TANTALUM PEGMATITES

Pegmatite deposits belong to a category of granite-related ore deposits distinct from magmatic ores with disseminated mineralization in granites. Cerny and Ercit (2005) provided an update for the classification of pegmatite deposits based on geochemical composition and the geological location. Using these criteria five classes were recognized:

1. Abyssal;
2. Muscovite;
3. Muscovite – rare element;
4. Rare-element;
5. Miarolitic.

Most of these classes can be further subdivided using geochemical and geological characteristics. The subclasses can also be broken down into types and subtypes using more subtle differences in geochemical signatures or pressure and temperatures conditions of solidification, evidenced by different accessory mineral assemblages. Cerny and Ercit (2005) propose a petrogenic subdivision developed for pegmatites derived by igneous differentiation from plutonic parents. Three groups are distinguished:

1. Progressive accumulation of Nb, Y and F (including Be, REE, Sc, Ti, Zr, Th and U), fractionated from sub-aluminous to meta-aluminous A- and I-type granites that can be generated by a variety of processes involving depleted crust or mantle contributions;
2. A peraluminous LCT group with prominent accumulation of Li, Cs and Ta (including Rb, Be, Sn, B, P and F), derived mainly from S-type granites, less commonly from I-type granites, and:
3. A mixed group with diverse origins, such as contamination of Nb-Y-F plutons by digestion of undepleted supracrustal rocks.

Using the Cerny and Ercit model the Zoro pegmatites are classified as rare-element LCT-type.

8.2 GENERAL CHARACTERISTICS OF RARE ELEMENT LITHIUM-CESIUM-TANTALUM PEGMATITES

Cerny et al. (2005) document rare-element pegmatite deposits of the LCT family from orogens spanning from early Archean to very recent (Pezzotta, 2000). The granite-pegmatite suites are syn- to late orogenic and related to fold structures, shears and fault systems. The pegmatites vary greatly in form and are controlled by the competency of the enclosing rocks, the depth of emplacement, and the tectonic regime during and after emplacement. The pegmatites rarely occur within their parent granites, but in three scenarios pegmatites form swarms or networks of fracture-filling dykes hosted by contraction fractures or structures generated by post-consolidation stresses (Ginsburg *et al*., 1979). Many of the deposits are hosted by schists and gneisses, and their shapes vary from lenticular, ellipsoidal, turnip- or mushroom-like forms in plastic environments, to fracture fracture-filling dykes and stocks in brittle host rocks (e.g. Cameron *et al*., 1949). The length of a mineralized pegmatite intrusion is typically tens to hundreds of metres, but they may be up to several km (Greenbushes, Australia; Partington *et al.*, 1995), and interconnected dyke systems are known to be up to 12 km long (Manono, Zaire; Thoreau, 1950).

Generally, a zoning sequence for individual pegmatite districts (Cameron *et al*., 1949; Cerny *et al*., 2005) has been recognized such that minerals present in each zone decrease in number from the margins (border and wall zones) to the central or latest primary unit, termed the core. Assemblages of the border and wall zones typically consist of quartz-plagioclase-microcline-biotite-garnet-tourmaline-(beryl-apatite), and the internal zoning sequence ends with nearly monomineralic masses of microcline followed by a monomineralic quartz core.

Cerny *et al*. (2005) note the shape and attitude of pegmatite intrusions have considerable control over the internal structure of the deposits such that homogeneous bodies are unusual. The pegmatites are largely concentrically zoned or layered, or they display a combination of both features (Cameron *et al.*, 1949; Cerny, 1991b). Concentric patterns typical of three-dimensional bodies can be extensively disturbed in flat pegmatites. Sub-vertical dykes commonly exhibit telescoping of strongly asymmetric zoning patterns, with the inner zones prominently shifted upward. The zoning progresses from finer grained zones of approximate granitic composition on the outside to inner zones that exhibit enrichment in rare-element mineralogy and textural diversity, some are also near-monomineralic.

In conjunction with the accumulation of rare-element mineralization in the inner zones, complex pegmatites also show inwardly increasing geochemical fractionation in rock-forming minerals (e.g. Cerny *et al*., 1985; Cerny, 2005; London, 2005) which serves as an important exploration guide (e.g. Cerny, 1992).

8.3 PEGMATITE EMPLACEMENT

Passive emplacement of pegmatite magma was historically advocated by many authors, but structural-geological analysis appears to contradict this interpretation (Cerny *et al*., 2005). Forcible intrusion is indicated in all closely examined cases (Brisbin, 1986). Beus (1966) determined empirically a distance of 2 km for the maximum distance of a pegmatite from its parent granite. Baker (1998), however considers the magma pressure in the parental chamber sufficient to propel low-viscosity pegmatite melts up to 10 km from the source.

In the presence of higher contents of Li, B., P, F and H_2O polymerization is reduced, fluidity and mobility are increased, and the thermal stability of pegmatite melts to lower temperatures is enhanced (Cerny *et al*., 2005). The result is that pegmatite melts enriched in volatiles and rare-elements can travel the farthest from the source (Figure 8.3.1). This explains the regional zoning of the rare-element pegmatites around parental granites (Cerny, 1992). The lithium-rich complex pegmatites in general and in particular the lepidolite-subtype dykes are invariably the most distal ones relative to the parent plutons (Cerny *et al.*, 2005). These categories of LCT rare-element pegmatites appear to be separated from granites by the interplay of hosts structures and erosional exposure.

Pegmatite dykes commonly occur as groups of similar pegmatite-types that originated from the same parent granite intrusion. A pegmatite field can occur over territories of hundreds to thousands of square km when favorable conditions are met.

Cerny *et al.* (2005) provide the following definitions for various groupings of pegmatites:

1. **Pegmatite Group**: a spatially and genetically coherent pegmatite population, generated by differentiation from a single granitic pluton. Pegmatite dykes interior, marginal, and exterior to a particular fertile granite intrusion maybe neatly distributed

around the plutonic parent, although asymmetric arrays are much more common (Figure 8.3.1.; Beus, 1966; Cerny, 1989b; 1990; 1991c; Cerny *et al*., 2005). In many examples radiometric dating confirms the link between fertile granites and surrounding pegmatite dykes (e.g. Baadsgaard and Cerny, 1993; Breaks *et al*., 2005). The pegmatites tend to show different kinds and degrees of mineralization in a regional zonal pattern concentric to unidirectional. The common progression from proximal to distal pegmatites is from barren to Be, Be-Nb-Ta, Li-Be-Ta-Nb and Li-Cs-Be-Ta-(F) assemblages, with B, P., and Sn appearing at locally different stages. The zoning tends to be strongly developed, with the most evolved pegmatites at the top of the three-dimensional array. Locally, the more evolved pegmatites are relatively late, as they crosscut the primitive dykes (Cerny, 1991c; 1992b).

2. **Pegmatite Fields**: result from favourable conditions for partial melting that generate fertile granites, are regional in scale, and commonly lead to intrusions and differentiation of multiple fertile plutons over territories of hundreds to thousands of square km (Cerny *et al*., 2005). The ensuing pegmatite fields contain granite-pegmatites suites that are more or less closely related, having been mobilized and differentiated from related or identical metamorphic protoliths during a single anatectic event. This results in similarities in mineral assemblages and geochemical signatures of the granite-pegmatite groups.
3. **Pegmatite Provinces**: are huge terranes characterized by commonality of geologic history that tend to generate arrays of pegmatite fields that are at lease loosely related in time, structural style, and mode of origin; geologic provinces locally represent rare-element pegmatite provinces of enormous dimensions (Cerny 1991a; c).

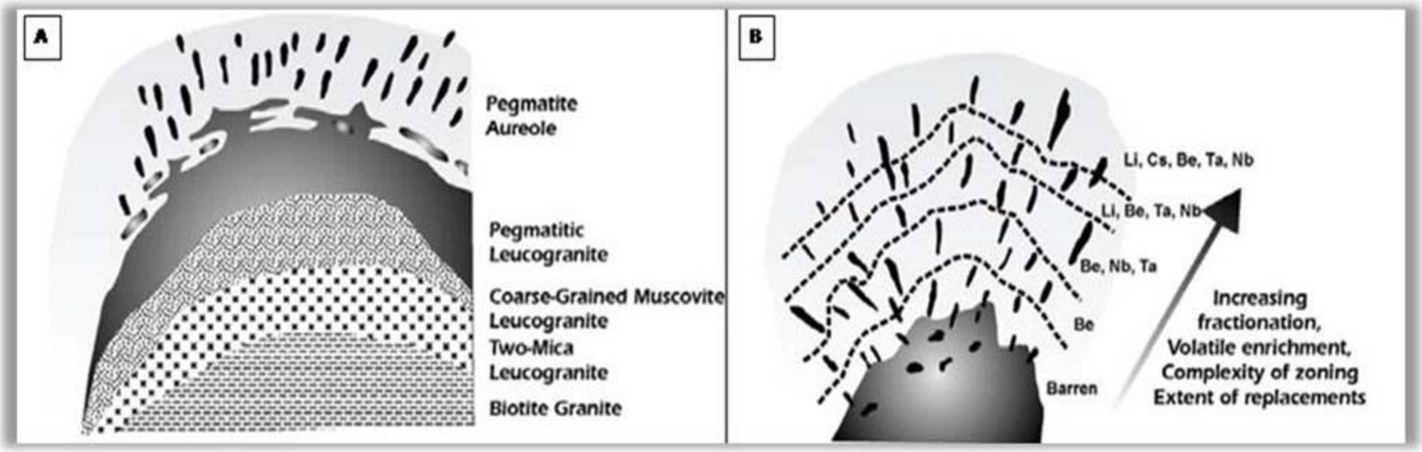


Figure 8.3.1: Regional Zoning in Fertile Granites and Pegmatites (after Cerny, 1991b and Selway *et al.*, 2005). A) Regional zoning of a fertile granite (outwardly fractionated) with an aureole of exterior lithium pegmatites; B) Schematic representation of regional zoning in a cogenetic parent granite and pegmatite group. Pegmatites increase in the degree of fractionation with increasing distance from the parent granite.

8.4 THE TANCO PEGMATITE

8.4.1 GENERAL

The Tanco deposit (Figure 8.4.1) is situated in the Archean Superior Province of the Canadian Shield in southeast Manitoba. The following description is taken from Cerny *et al*. (1998), Cerny (2005), Cerny *et al*. (2005) and Martins *et.al*. (2013). The pegmatite has an age of 2640 Ma and is completely blind forming a sub-horizontal lenticular undeformed bilobate intrusion consisting of four concentric and five layered zones with approximate dimensions of 1520 m in length, 1060 m in width and 100 m in thickness. Its parental granite is not exposed, however, nearby pegmatite groups of similar character show a clear connection to pegmatitic leucogranites. The deposit is not exposed as it lies beneath Bernic Lake and accordingly geological observations are based on examination of drill core and exposures created during underground mining. Tanco is a highly fractionated lithium-cesium-tantalum (LCT) pegmatite with an extensive mineralogy of >100 minerals from nine internal zones. The zonation consists of concentric outer zones and segmented and complex layered inner zones.

8.4.2 ZONATION AND MINERALOGY

The Tanco pegmatite is a complexly zoned intrusion described in detail by various researchers (Cerny, 2005; Cerny *et al.*, 1996, 1998; Stilling *et al*., 2006). Table 8.4.1 summarizes zonation, mineralogy, texture and geochemistry for the deposit.

8.4.3 GEOCHEMISTRY

The Tanco pegmatite is a peraluminous LCT-type pegmatite intrusion belonging to the Rare Element-Lithium subclass, complex type, subtype petalite of Cerny and Ercit (2005). Based on the work of Stilling et. al., 2006 the bulk mode of Tanco closely approximates a muscovite granite with the exception of 8 weight % petalite, 2.8 weight % lithian mica and 1 weight % primary spodumene. The percentages of all other silicate and phosphatic mineral phases are very low. Tanco therefore approximates a moderately silicic, high phosphorus Na>K granite enriched in Li, Rb, Cs and F, moderate contents of Tl, Be, B, Ga, Sn, Nb and Ta and highly depleted in Fe, Mn, Mg, Ca, Ba, Sc, Ti and Zr. A high degree of fractionation is indicated by values for K/Rb of 4.7, K/Cs of 9.3, Rb/Cs of 2.0, Rb/Tl of 137, Fe/Mn of 0.63, Mg/Li of 0.02, Al/Ga of 917, Zr/Hf of 2.6, Zr/Sn of 0.21 and Nb/Ta of 0.19.

8.4.4 LITHIUM-CESIUM-TANTALUM MINERALIZATION

Tanco has been a major producer of spodumene (Li), pollusite (Cs) and tantalite (Ta). High tantalum zones in the deposit are spatially associated with metagabbro rafts however textural and geochemical studies (van Lichtervelde *et al*., 2007) indicate the tantalum mineralization is likely of magmatic origin.

Table 8.4.1: Zonation, mineralogy, textures and geochemistry of the Tanco pegmatite (Cerny, 2005).

Zone	Main constituents	Characteristics subordinate (accessory) and rare minerals	Textural and structural characteristics	Geochemistry important major & minor elements
Exomorphic unit	Biotite, tourmaline, holmquistite	Arsenopyrite	Fine-grained reaction rims and diffuse veins	K, Li, B (P, F)
(10) Border zone	Albite, quartz	Tourmaline, apatite, (biotite), beryl, triphylite	Fine-grained layers	Na, (B, P, Be, Li)
(20) Wall zone	Albite, quartz	Beryl, (tourmaline), muscovite, Li-muscovite, microcline-perthite	Medium-grained, with giant K-feldspar crystals	K, Na, (Li, Be, F)
(30) Aplitic albite zone	Albite, quartz, (muscovite)	Muscovite, Ta-oxides, beryl, (apatite, tourmaline, cassiterite), ilmenite, zircon, sulfides	Fine-grained undulating layers, fracture fillings, rounded blebs, diffuse veins	Na, (Be, Ta, Sn, Zr, Hf, Ti)
(40) Lower intermediate zone	Microcline-perthite, albite, quartz, spodumene,	Li-muscovite, lithiophilite, lepidolite, petalite, Ta-oxides	Medium- to coarse-grained; heterogeneous	K, Na, Li, P, F, (Ta)
(50) Upper intermediate zone	Spodumene, quartz, amblygonite	Microcline-perthite, pollucite, lithiophilite, (albite, Li-muscovite), petalite, eucryptite, Ta-oxides	Giant crystal size of major and most of the subordinate minerals	Li, P, F, (K, Na, Cs, Ta)
(60) Central intermediate zone	Microcline-perthite, quartz, albite, muscovite	Beryl, (Ta-oxides), zircon, ilmenite, spodumene, sulfides, lithiophilite, apatite, cassiterite	Medium- to coarse-grained	K, (Na, Be, Ta, Sn, Zr, Hf, Ti)
(70) Quartz zone	Quartz	Spodumene, amblygonite	Monomineralic	Si, (Li)
(80) Pollucite zone	Pollucite	Quartz, spodumene, petalite, muscovite, lepidolite, albite, microcline, apatite	Almost monomineralic	Cs, (Li)
(90) Lepidolite zone	Li-muscovite, lepidolite, microcline-perthite	Albite, quartz, beryl, (Ta-oxides, cassiterite), zircon	Fine-grained	Li, K, Rb, F, (Na, Be, Ta, Sn, Zr, Hf, Ga)

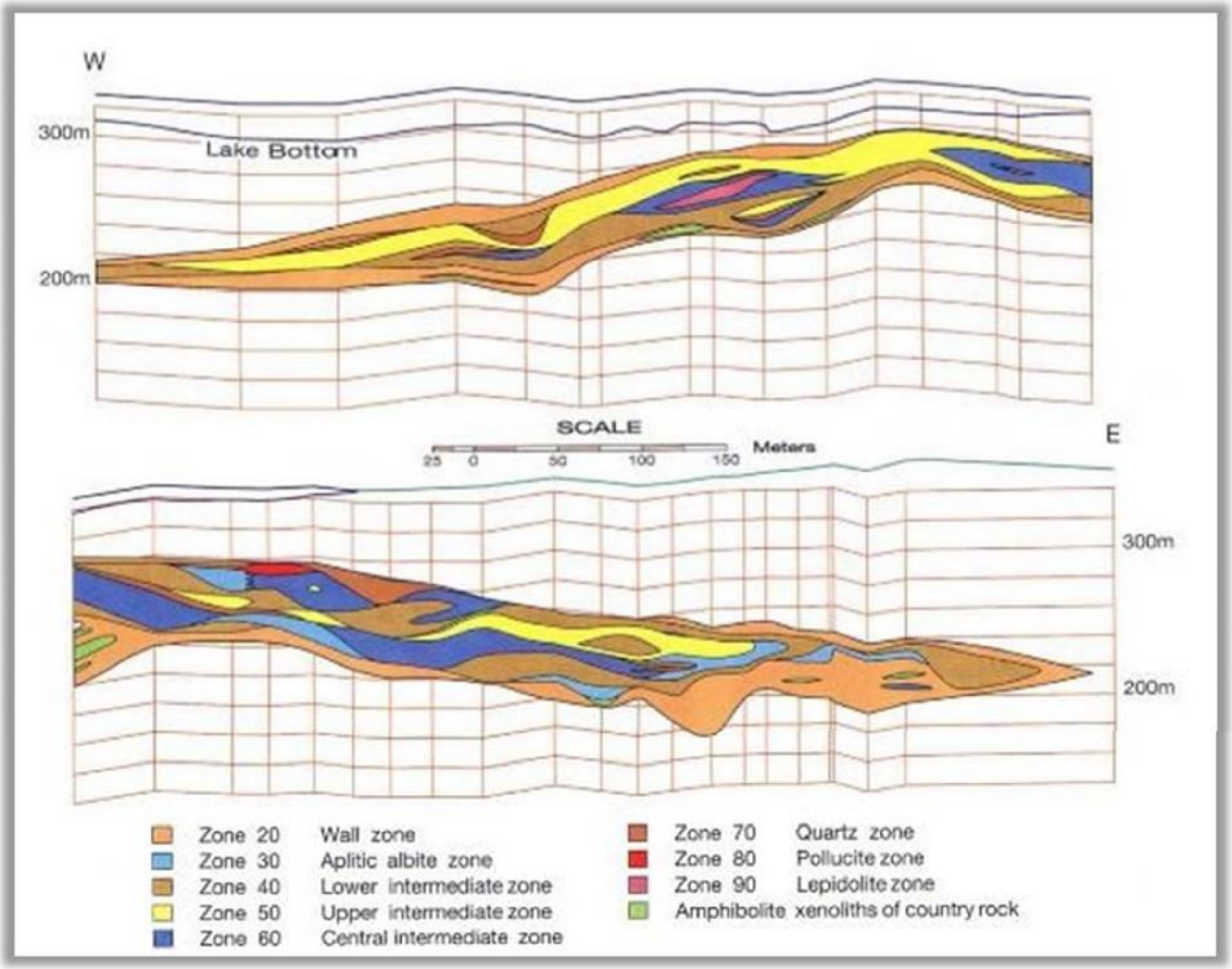


Figure 8.4.1: West to east longitudinal section through the Tanco pegmatite. (Modified from Stilling *et al*., 2006; Cerny *et al*., 2005). *Note:* The border zone (Zone 10) is not shown at this scale.

8.4.5 *ALTERATION AND ELEMENT DISPERSION*

Alteration haloes resulting from metasomatism have been documented around many pegmatites, with Cabot Corporation’s Tanco mine in the Bird River greenstone belt of southeastern Manitoba being the most studied example in the province (e.g., Trueman, 1978; Morgan and London, 1987; Halden *et al*., 1989). At Tanco, this type of country-rock metasomatism has been utilized for exploration (Trueman, 1978), and this methodology has since been applied throughout the Bird River greenstone belt (Galeschuk and Vanstone, 2005; 2007; Linnen *et al*., 2015). Lithium anomalies define the widest haloes adjacent to Li-Cs-Ta (LCT) pegmatites (Linnen *et al*., 2012) and, in the case of Tanco, Li haloes have been recognized to extend more than 100 m away from the pegmatite body (Cerny, 1989). However, dispersion of other elements such as Rb and Cs seems to be more restricted (e.g., Cerny, 1989; London, 2008). This type of country-rock alteration is caused by the influx of pegmatite magma and coexisting fluids rich in incompatible elements. The composition of the fluid phase is related to the magma composition; therefore, the diagnostic elements of the alteration aureoles are related to element enrichments and mineralogy of the associated pegmatite intrusion (Beus, 1960). In the case of evolved LCT pegmatites, the adjacent country rock is altered by an influx of alkali rare elements (e.g., Li, Rb and Cs) and subsequent interaction between the fluid phase and the country rock, forming a dispersion halo. This interaction results in a change of the composition of pre-existing mineral assemblages in the country rock and stabilization of exotic mineral assemblages. Metasomatism by Li-enriched fluids can produce holmsquistite-bearing assemblages in amphibolitic country rock, as has been documented at several locations, including the Edison pegmatite in the Black Hills of South Dakota (Shearer *et al*., 1986; Shearer and Papike, 1988) and the Tanco pegmatite in Manitoba (Morgan and London, 1987; Selway *et al*., 2000). These alteration assemblages can be a good exploration tool and have been used in many pegmatite districts (e.g., Beus, 1960; Truman and Cerny, 1982; Norton, 1984; London, 1986). This study focuses on alteration haloes caused by 1) the Dibs LCT pegmatite from the Cat Lake–Winnipeg River pegmatite field in the Archean Bird River greenstone belt; and 2) the Dyke 1 LCT pegmatite from the Wekusko Lake pegmatite field in the Paleoproterozoic Flin Flon–Snow Lake greenstone belt. Although the ages differ, both bodies intrude metamorphosed volcanic rocks and the premise for this study is that both would be associated with above-normal background values for elements that are enriched in the pegmatite. Factors that could influence the metasomatic halo around Li-bearing pegmatites, include 1) the relationship between dyke thickness and the size of the metasomatic halo; 2) the shape of the halo related to the location of the Li mineralization within the pegmatite; 3) fluid pressures at time of emplacement; 4) structural permeability; 5) country-rock composition; 6) emplacement history; and 7) overprinting by later structural, metamorphic or hydrothermal events.

8.5 RARE-ELEMENT PEGMATITES FROM THE SUPERIOR PROVINCE

Selway *et al*. (2005) reviewed rare-element pegmatites in the Superior Province of Ontario and Manitoba and determined rare-element pegmatite dykes within this Province cluster to form pegmatite fields that contain one or two large and highly fractionated pegmatites and numerous small pegmatite dykes. An example would be the Bernic Lake pegmatite group, part of the Cat Lake-Winnipeg River pegmatite field in southeastern Manitoba, that includes the Tanco pegmatite (1.99 km long x 1.06 km wide x 100 m thick; Stilling, 1998) and eight other smaller, less-fractionated pegmatite dykes (Cerny *et al*.,1981). The Separation Rapids pegmatite group lies to the east of the Cat Lake–Winnipeg River pegmatite within the same Bird River–Separation Lake metavolcanic belt (Breaks *et al*., 1975). The Separation Rapids pegmatite group contains two large highly fractionated pegmatites: Big Whopper (350 m in strike length x 60 m thick) and Big Mack (30 x 100 m; Breaks and Tindle,1997 Breaks *et al*., 1999). The Big Whopper and Big Mack pegmatites are members of the Southwestern pegmatite subgroup, which contains at least 23 additional smaller pegmatite dykes. Additional large pegmatite fields in the Superior Province of Ontario with economic potential include: the Dryden pegmatite field, which includes the highly-fractionated Fairservice pegmatite dykes and Tot Lake pegmatite, and the Seymour Lake pegmatite group, which includes the highly-fractionated North Aubry and South Aubry pegmatites (Breaks *et al*., 2003). These pegmatites contain elevated Rb, Cs, Be, and Ta contents. The Case pegmatite in northeastern Ontario is unique in that it is a large fractionated pegmatite with no identified associated smaller pegmatite dykes, likely due to thick overburden (Breaks *et al*., 2003).

Selway et al. (2005) noted geological features common amongst pegmatites of the Superior Province of Ontario (Breaks and Tindle, 2001; Breaks *et al*., 2003) and Manitoba (Cerny *et al*., 1981; Cerny *et al*., 1998):

1. The pegmatites tend to occur along sub-province boundaries. For example, Tanco (Manitoba) and Separation Rapids (Ontario) pegmatites within the Bird Lake-Separation Lake metavolcanic belt occur along the boundary between the English River and Winnipeg River sub-provinces; the beryl-phosphate Sandy Creek and McCombe pegmatites, and the Lilypad Lake pegmatite field occur along the Uchi–English River sub-provincial boundary; the Dryden pegmatite field occurs within the Sioux Lookout Domain along the Winnipeg River–Wabigoon sub-provincial boundary; and the North Aubry, South Aubry, and Tebishogeshik pegmatites occur along the English River–Wabigoon sub-provincial boundary north of Armstrong.
2. Most pegmatites in the Superior Province (in Ontario and Manitoba) occur along sub-province boundaries, except for those that occur within the metasedimentary Quetico Subprovince. Examples of pegmatites occurring in this area from west to

east are: Wisa Lake (south of Atikokan), the Georgia Lake pegmatite field (north of Nipigon), and the Lowther Township (south of Hearst) pegmatites.

3. Pegmatites are present at greenschist to amphibolite metamorphic grade. In Ontario and Manitoba, pegmatites are absent in the granulite terranes of the Quetico and English River sub-provinces.
4. Most pegmatites in the Superior Province (Ontario and Manitoba) are genetically derived from fertile parent granite. The Cat Lake–Winnipeg River pegmatite field (Manitoba) contains six leucogranite intrusions (Greer Lake, Eaglenest Lake, Axial, Rush Lake, Tin Lake, and Osis Lake) emplaced along east-trending faults, which are parents to numerous pegmatites (Cerny *et al*., 1981; Cerny et al., 1998). In contrast, the Tanco pegmatite has no fertile granite outcropping in reasonably close vicinity that could be its potential parent (Cerny *et al*., 1998). The peraluminous Separation Rapids pluton (4 km wide) is the parent to the Separation Rapids pegmatite field, including Big Whopper and Big Mack pegmatites, north of Kenora, Ontario. The peraluminous Ghost Lake batholiths (80 km wide) is the parent to the Mavis Lake pegmatite group, including the Fairservice pegmatite dykes, north of Dryden, Ontario.
5. Highly fractionated spodumene- and petalite- subtype pegmatites are commonly hosted by mafic metavolcanic rocks (amphibolite) in contact with a fertile granite intrusion along sub-provincial boundaries, whereas numerous beryl-type pegmatites are hosted by metasedimentary rocks (metawacke or metapelite) of the Sioux Lookout Domain. Pegmatites within the Quetico Subprovince are hosted by metasedimentary rocks or their fertile granitic parents.
6. Biotite and tourmaline are common minerals within metasomatic aureoles in mafic metavolcanic host rocks to pegmatites. Tourmaline, muscovite, and biotite are common within metasomatic aureoles in metasedimentary host rocks.
7. Most of the pegmatites of the Superior Province contain spodumene and/or petalite as the dominant lithium mineral, except for the Lilypad Lake, Swole Lake, and Lowther Township pegmatite (all in Ontario), and the Red Cross Lake lithium pegmatite (Manitoba), which have lepidolite as the dominant Li mineral. Amblygonite- and elbaite-dominant pegmatites have not yet been found in the Superior Province, although amblygonite and elbaite occur in the Tanco pegmatite.
8. Cesium-rich minerals only occur in the most extremely fractionated pegmatites. Pollucite occurs in the Tanco, Marko's, and Pakeagama petalite-subtype pegmatites, the Tot Lake spodumene-subtype pegmatites, and the Lilypad Lake

lepidolite-subtype pegmatites (Teertstra and Cerny, 1995). The Pakeagama pegmatite in northwestern Ontario occurs along the Sachigo-Berens River subprovincial boundary. Cesium-rich beryl occurs in the spodumene-subtype North Aubry, South Aubry, Case, Tot Lake, and McCombe pegmatites and the lepidolite-subtype Lowther pegmatite, all in Ontario, and in the Tanco pegmatite, Manitoba.

9. Most of the pegmatites in the Superior Province contain ferro-columbite and mangano-columbite as the dominant Nb-Ta-bearing minerals. Some pegmatites contain mangano-tantalite as the dominant Ta-oxide mineral, for example the North Aubry, South Aubry, Fairservice, Tot Lake, and Tebishogeshik pegmatites. The Tanco pegmatite contains wodginite as the dominant Ta-oxide mineral. Tantalum-bearing cassiterite is relatively rare in pegmatites of the Superior Province, except for the Separation Rapids and Tanco pegmatites.
10. Fine-grained tantalum-oxides (e.g. manganotantalite, wodginite, and microlite) commonly occur in the aplite, albitized K-feldspar, mica-rich, and spodumene core zones in pegmatites in the Superior province. At Tanco, tantalum mineralization occurs in the albitic aplite zone (30), central intermediate muscovite-quartz after microcline zone (60), and lepidolite zone (90).

8.6 THE ZORO PEGMATITES

Dyke 1 of the Zoro pegmatite dykes is classified as a rare metal spodumene pegmatite (Rudenko *et al*. (1975)) though it may be marginal to the lepidolite sub-formation (Table 8.6.1). It shares many of the characteristics of Superior Province pegmatites described by Selway *et. al*. (2005) including the Tanco pegmatite.

Table 8.6.1: Pegmatite classification abbreviated after Rudenko *et al.* (1975) showing the possible field of the Dyke 1 spodumene pegmatite (from Cerny, 1982).

Pegmatite Formation	Pegmatite Sub-formation	Pegmatite Type	Mineralization	Parent granites	Level of emplacement	Metamorphic grade of host
Ceramic		Ceramic	Ceramic: U, REE	Normal biotite (?); leucocratic and alaskitic	Very deep	Granulite; sill. – alm. amphibolite
Silica-bearing		Muscovite-bearing	Muscovite: U, REE, Be		Deep	Kyanite-alm. amphibolite
Rare-metal	Beryllium-bearing	Be, Ta, Nb (Sn)	Be (Ta, Nb)	Leucocratic, alaskitic	Intermediate (rarely shallow)	Staurolite-alm. Amphibolite, epidote-amphibolitic, greenschist
	Pollucite-bearing	Ta, Cs, Be, Li, Rb (Sn)	Be, Li, Cs, Ta, Rb	F, Li-bearing		
	Spodumene (Li)	Li, Be, Ta, Nb (Sn)	Li, Ta, Be (Nb)			

	Lepidolite (F, Li)	Li, Ta, Be (Sn)	(Ta, Li, Cs, Be)			
	Rare Earth-element-bearing	REE, U, Th, Nb	(REE, Nb)	Alkalic		
Rock -crystal-bearing		No cavities, quartz phyroblasts	Rock crystal	Leucocratic, alaskitic	(Intermediate) shallow	–
		Cavities – Murzinka-type				
		Cavities – Kazakhstan-type				

The Far Resources exploration program will utilizing the well-documented characteristics of lithium-enriched pegmatite and related alteration phenomena in the host wallrocks to advance the property. Exploration will monitor the geochemical signature of underlying spodumene-bearing pegmatites in overlying soils using Mobile Metal Ions Technology with follow-up diamond drilling.The association of pegmatite with late fractures provides pathways for element migration from lithium source to surface where analysis of soil samples will define drill targets.

9.0 EXPLORATION

9.1 FAR RESOURCES LTD.

Commencing in 2016 Far Resources undertook prospecting, geological mapping, soil geochemical surveys, mineral and rock geochemical surveys and collaborative research all leading to four diamond drill campaigns. In support of these activities the limited existing historic databases were utilized to produce preliminary three-dimensional models of the spodumene-bearing dykes on the property to assist drill targeting. These models are presented in Figures 9.3.1 through 9.3.5 and Table 9.3.1 summarizes physical characteristics from historic descriptions of the dykes as exposed in outcrop.

Table 9.3.1: Summary of historic physical characteristics of dykes 2 through 7 (Assessment File AF95362).

DYKE	LENGTH (feet)	WIDTH (feet)	DECLINATION	BULK SAMPLE ASSAYS % Li_2O
Dyke 2				
Northwest Dyke	150	2-7	North 32 Degrees West	1.69
Northeast Dyke	25	N/A	North 25 Degrees West	1.66
South Dyke	>250	6-10	North 48 Degrees West	
Dyke 3				
North Dyke	200	4-6	North 40 Degrees West	
West Dyke	100	3-5	North 55 Degrees West	
East Dyke	175	5-15	North 40 Degrees West	
Dyke 4	500	2-10	North 35-45 Degrees West	1.12
				1.16
Dyke 5	600	5-35	North 60 Degrees West	2.26
				2.22
Dyke 6	350	1-5	North 25 Degrees West	0.46
				0.50

Dyke 7	250	3-10	North 73 Degrees West	

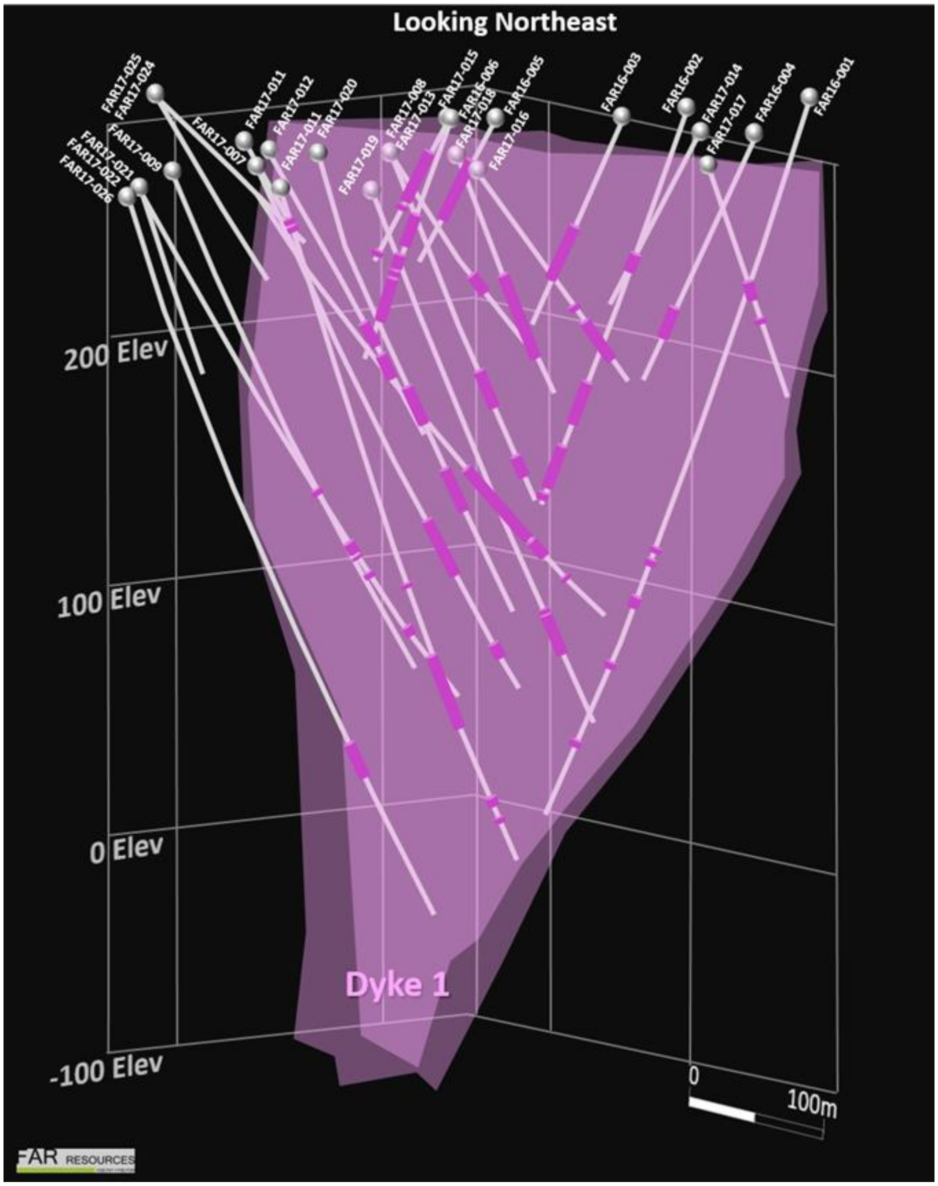


Figure 9.3.1: Reconstruction of Dyke 1 based on historic and recent drill information including spodumene intercepts.

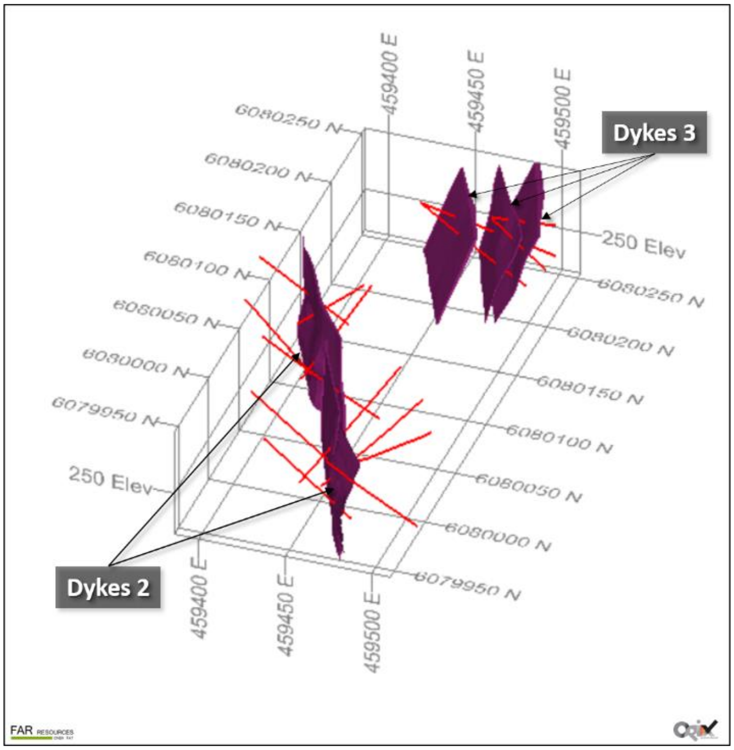


Figure 9.3.2: Reconstruction of Dykes 2 and 3 based on historic drill information.

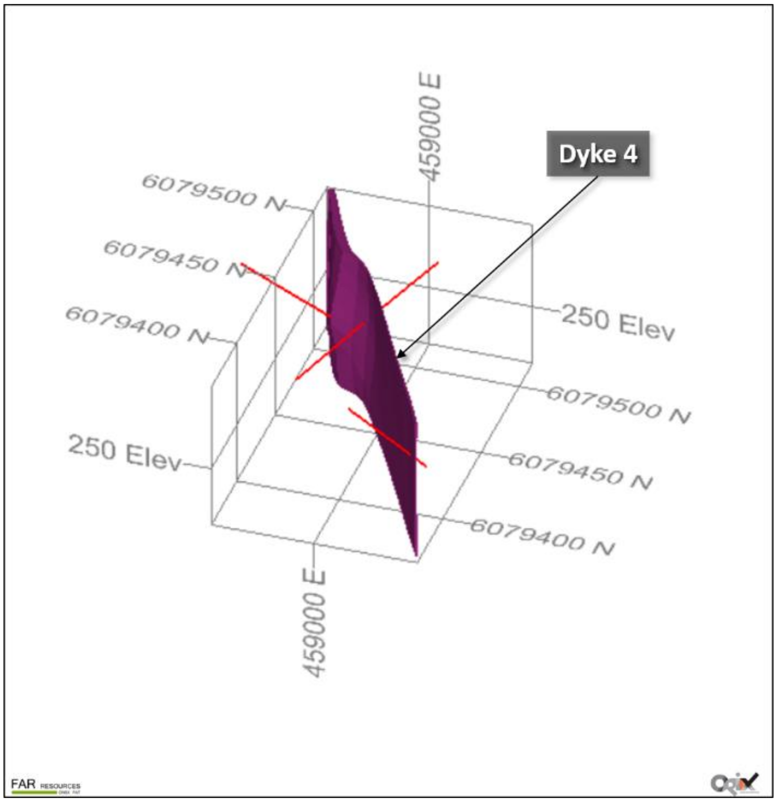


Figure 9.3.3: Reconstruction of Dyke 4 based on historic and recent drill information.

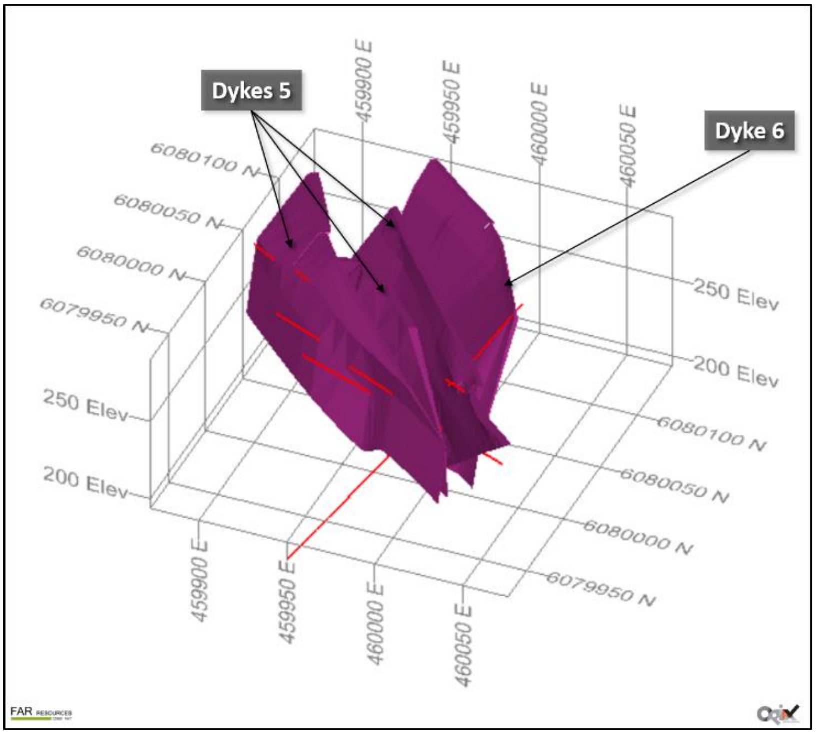


Figure 9.3.4: Reconstruction of Dykes 5 and 6 based on historic drill information.

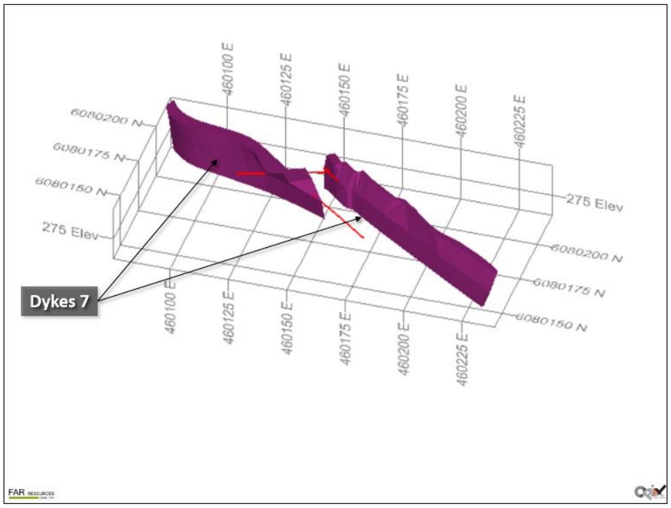


Figure 9.3.5: Reconstruction of Dyke 7 based on historic drill information.

9.1.1 PROSPECTING

Prospecting teams active on the Zoro property have prepared existing outcrop exposures for sampling and mapping and have documented new outcrop exposures for follow-up. An expanded property position has resulted in part from these activities. The teams have also undertaken rock chip and soil geochemical (Mobile Metal Ions) surveys to quantify and qualify historic lithium assays and assess overburden covered terrain for geochemical signatures of buried lithium-bearing pegmatite. Glacial and organic sediments are extensive on the property and blanket areas where the potential for additional pegmatite dykes and the delineation of potential drill targets is high. Results are discussed below in their respective sections.

9.1.2 GEOLOGICAL MAPPING AND PETROGRAPHY

Geological mapping of the entire Zoro Lithium Project has not been undertaken. Limited mapping has been undertaken near dykes 5, 6 and 7 to determine the attitude of the pegmatite dykes and assess them for possible drill targets. Results are presented in Figure 9.3.6 and illustrates what is tentatively interpreted to be a dyke swarm in this area. Outcrop exposure is poor due to glacial sediments, organic soils and lichen cover on available outcrop.

Dyke 1 pegmatite is the largest and best studied dyke on the property. It is a north-south trending, near vertical body that extends for at least 280 m in length and a maximum known thickness of approximately 35 m. An apparent lack of alteration in the country rock is commonly described in the historical drill logs with only a local description of brecciation of the mafic volcanic rocks associated with a quartz network of veins. Recent field work identified holmquistite in the mafic volcanic country rock, indicating metasomatic alteration associated with the pegmatite intrusion, and lithogeochemical analyses demonstrate that a broad metasomatic halo was developed. Holmquistite-bearing assemblages are a function of the activity of Li introduced into the pegmatite’s wall rock. These assemblages reflect greenschist facies metamorphic conditions and are only found in amphibolite wall rock usually replacing hornblende, pyroxene or biotite (Heinrich, 1965; London, 1986). Based on historical and recent drill log descriptions the zonation in the Dyke 1 pegmatite can be defined as follows:

1) the wall zone is at the contact and is predominately composed of quartz, microcline and muscovite, with accessory tourmaline, hornblende, biotite and rare beryl and spodumene;
2) intermediate zone with medium sized crystals of microcline, albite, quartz, muscovite and spodumene (<5%);
3) central zone with abundant spodumene (locally up to 50% but more commonly varying between 10 and 30%), albite, quartz and locally pollucite and tantalite, and accessory apatite, tourmaline, pyrrhotite, lepidolite, columbite group minerals and Fe-Mn phosphate minerals;

4) core zone is mostly composed of quartz with small to medium grained spodumene crystals (locally 15-20 cm crystals of spodumene are observed) in a quartz matrix, with minor tourmaline and muscovite.

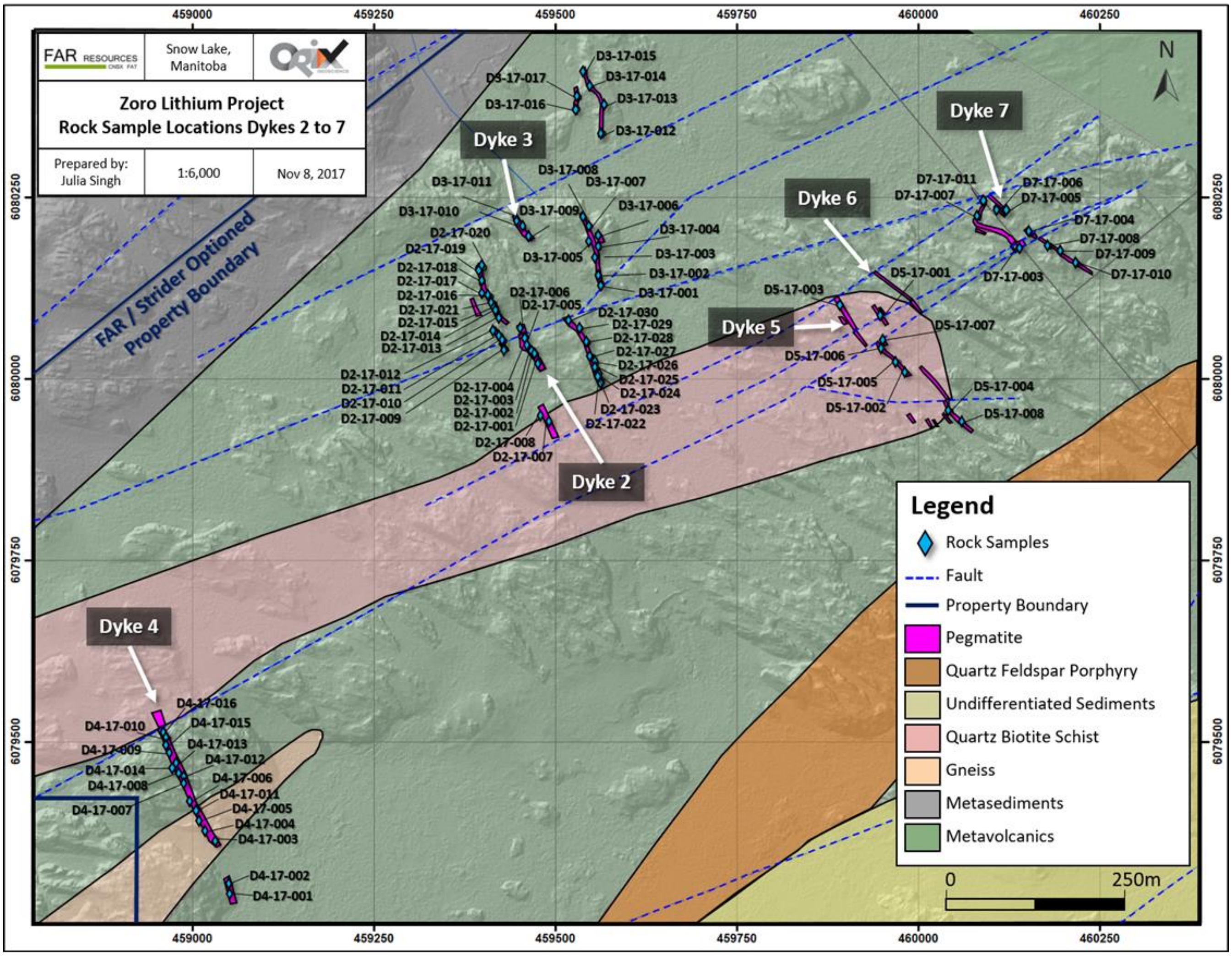


Figure 9.3.6: Rock chip sample location map for the assessment of lithium and multi-element compositions of the seven historic dykes on the Zoro Lithium Project. The results of geological mapping and LIDAR surveys form the base for this figure.

Based on historical descriptions and recent preliminary petrographic work three stages of spodumene growth can be identified. The first is a primary phase greenish spodumene and the second spodumene plus quartz intergrowths possibly after petalite breakdown (Cerny and Ferguson, 1972). The third phase consists of late bands of very fine-grained spodumene crystals that cross cut other mineral phases or surround feldspars and muscovite grains. Locally, spodumene crystals can be surrounded by fine-grained mica described in historical drill logs as purple, possibly Li-mica or lepidolite. This could be indicative of a late Li-enriched episode (possibly auto-metasomatism) responsible for the formation of the later mica. Acicular opaque minerals of the columbite group are present, and locally late bands of fluorite were reported in historic assessment files associated with fractures in the pegmatite. The latest event identified in thin section is characterized by late Fe-rich, quartz-calcite stringers with no definite direction that crosscut the pegmatite. Deformation is visible in thin section in the feldspars and muscovite (kink bands in muscovite are commonly observed) suggesting that the pegmatites are pre- to syn-deformational.

9.1.3 ROCK GEOCHEMICAL SURVEYS

Rock chip samples were collected as composite chips over the exposed area of the dykes and are illustrated in Figure 9.3.6. The purpose of the surveys was to confirm historic lithium assays presented in Manitoba government assessment files and build a database for use in focused exploration including diamond drilling. Select assay results are given in Figure 9.3.7 for dykes 2 and 3, in Figure 9.3.8 for dyke 4, and in Figure 9.3.9 for dykes 5, 6 and 7. Assay results for samples collected from Dyke 1 are discussed in section 9.2 under Force Energy. Persistent elevated grades of lithium (Li_2O) are documented from all dykes on the property in association with light green to white spodumene.

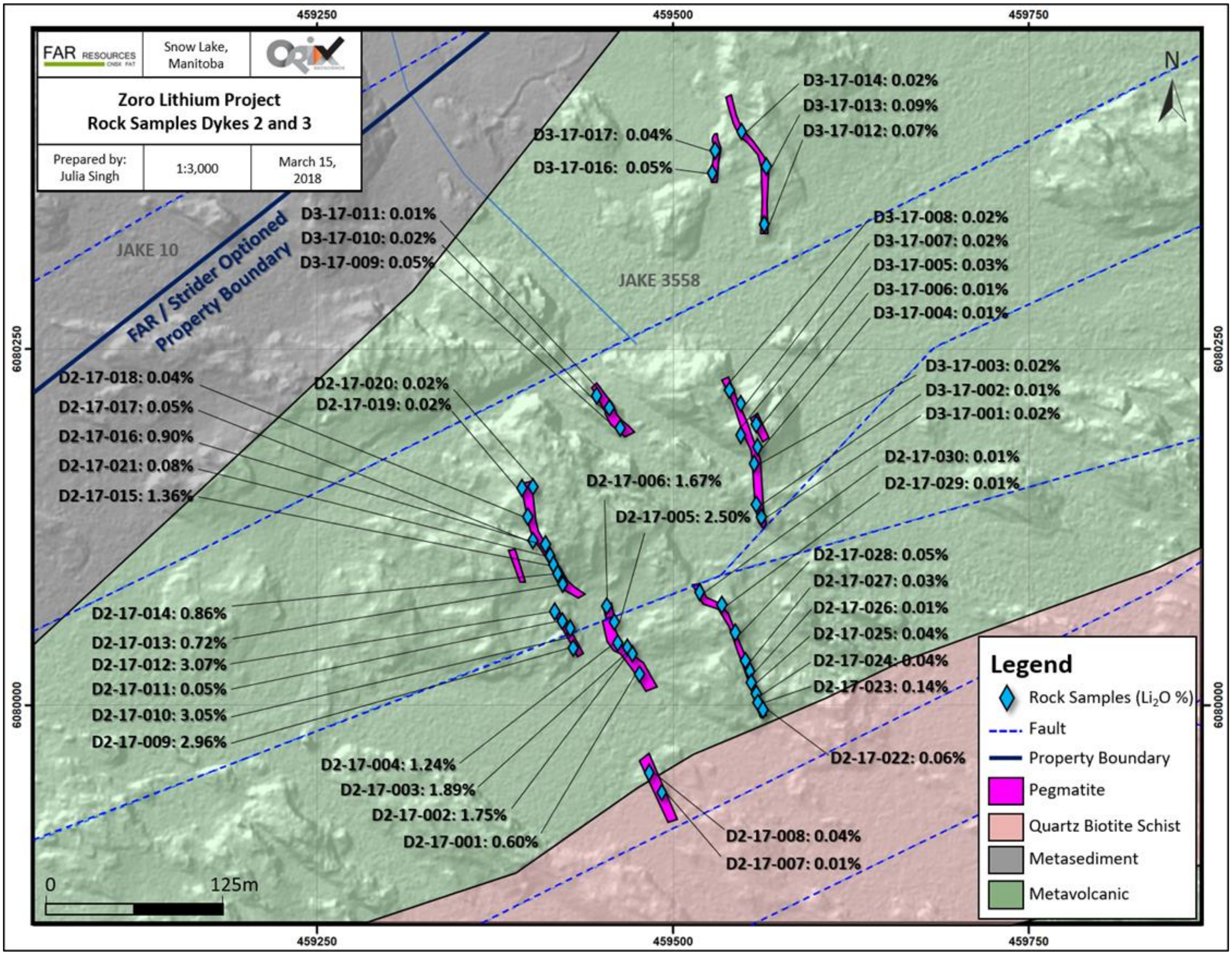


Figure 9.3.7: Assay results for rock chip samples at pegmatite dykes 2 and 3.

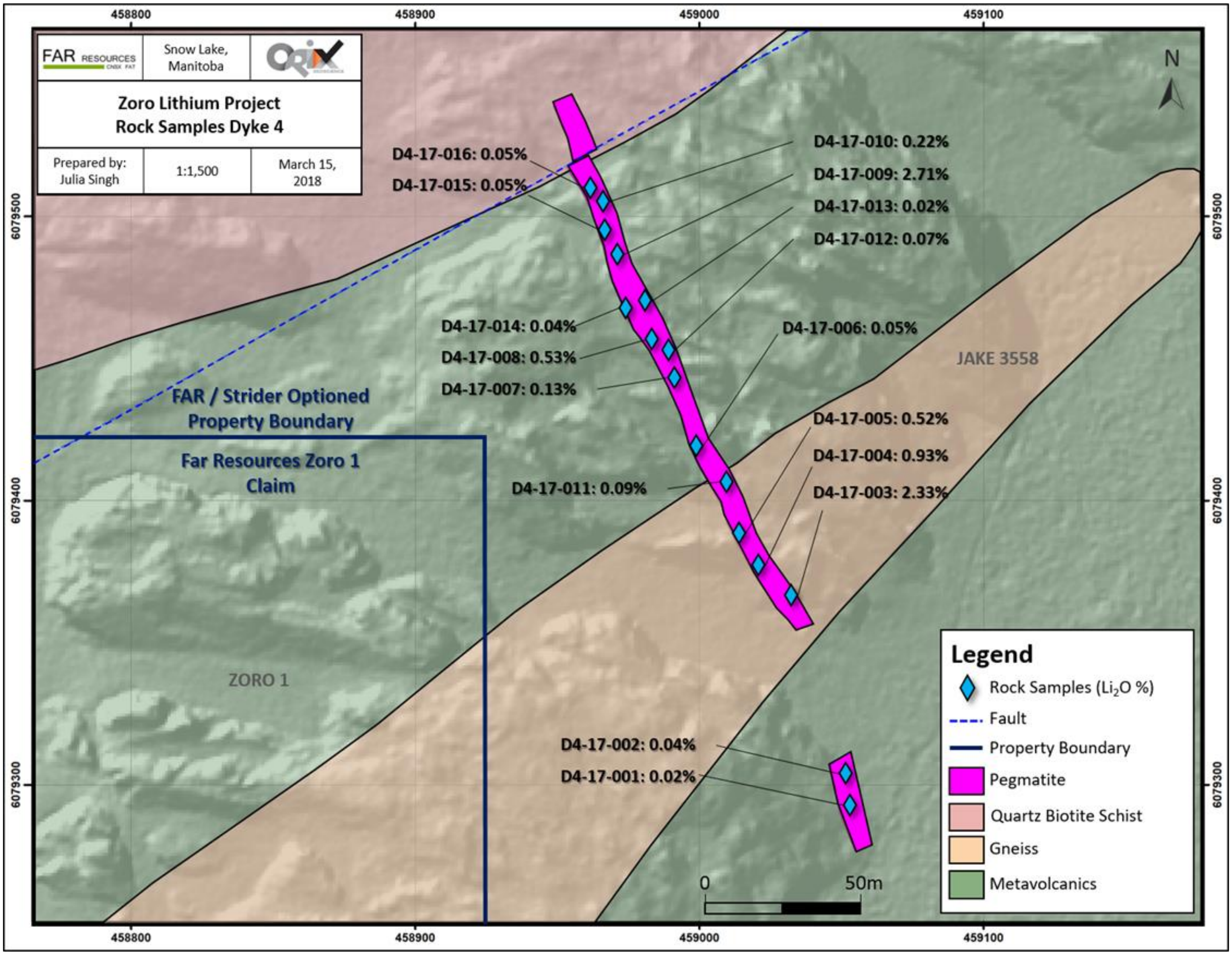


Figure 9.3.8: Assay results for rock chip samples at pegmatite dyke 4.

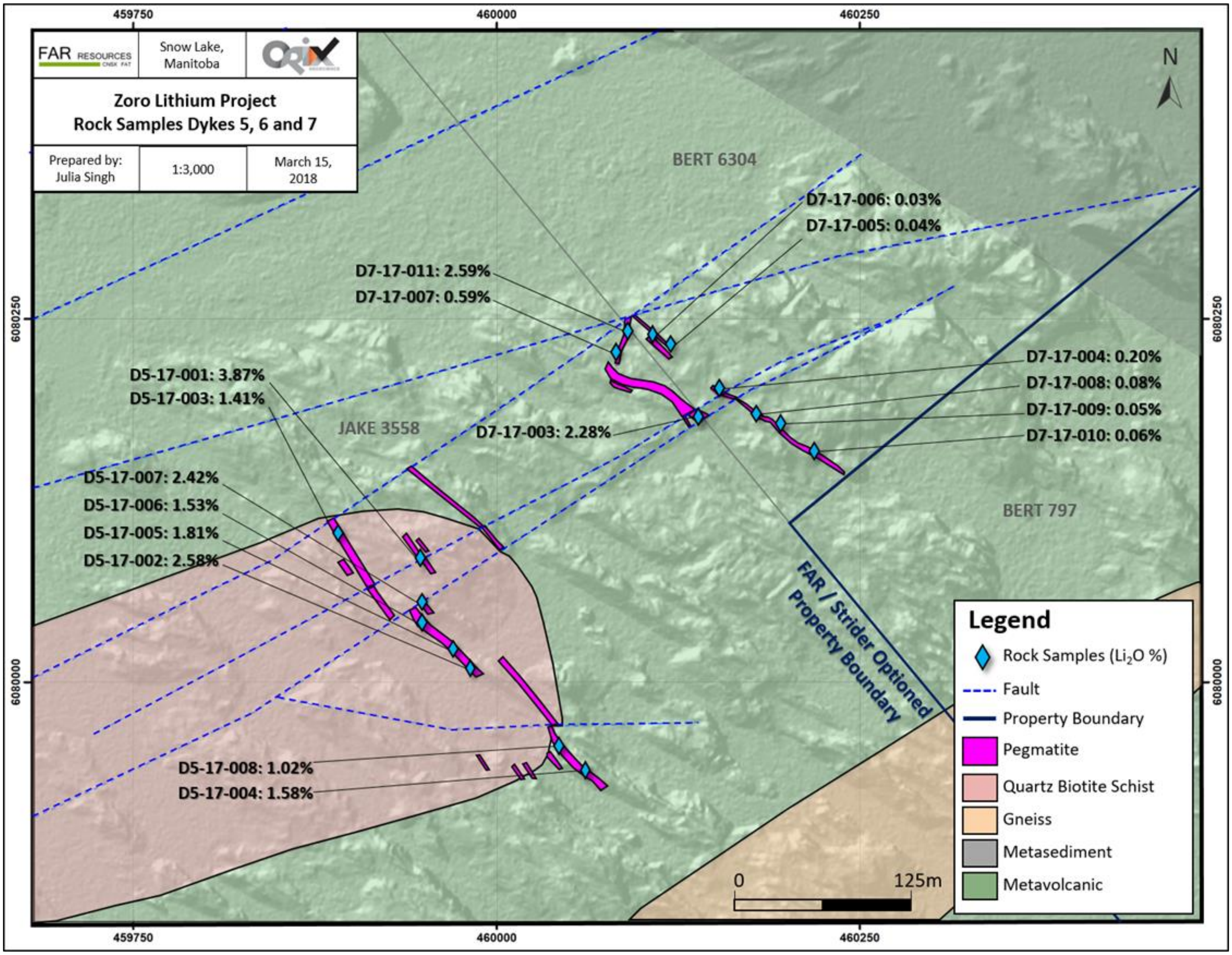


Figure 9.3.9: Assay results for rock chip samples at pegmatite dykes 5, 6 and 7.

9.2 SOIL GEOCHEMICAL SURVEYS

The abundance of glacial inorganic soil and organic soil cover on the property is problematic for exploration. The use of Mobile Metal Ions (MMI) soil geochemistry was utilized to assess overburden covered areas for extensions of known lithium pegmatite and to assess areas on the property where potential for additional pegmatite occurs but for which there is no outcrop. Samples were collected at 25 m sample sampling between 10 and 25 cm below the contact between organic and inorganic soil using a Dutch auger. Sample locations are given in Figures 9.4.1 and 9.4.2.

9.2.1 RESULTS

Bubble plots depicting the variation in MMI-extractable lithium and tantalum are given in Figures 9.4.3 through 9.4.6. The results indicate good correlation between MMI Li and Ta and known Li-Ta-bearing dykes. In some areas elevated MMI Li and Ta extend outwards into overburden covered terrain and this is interpreted to indicate the potential for the extensions of the dykes and as such exploration targets. Elevated Li and Ta are also noted from areas where there are no outcrops and these anomalies represent exploration drill targets. Results for Cs and Rb correspond with those for Li and Ta.

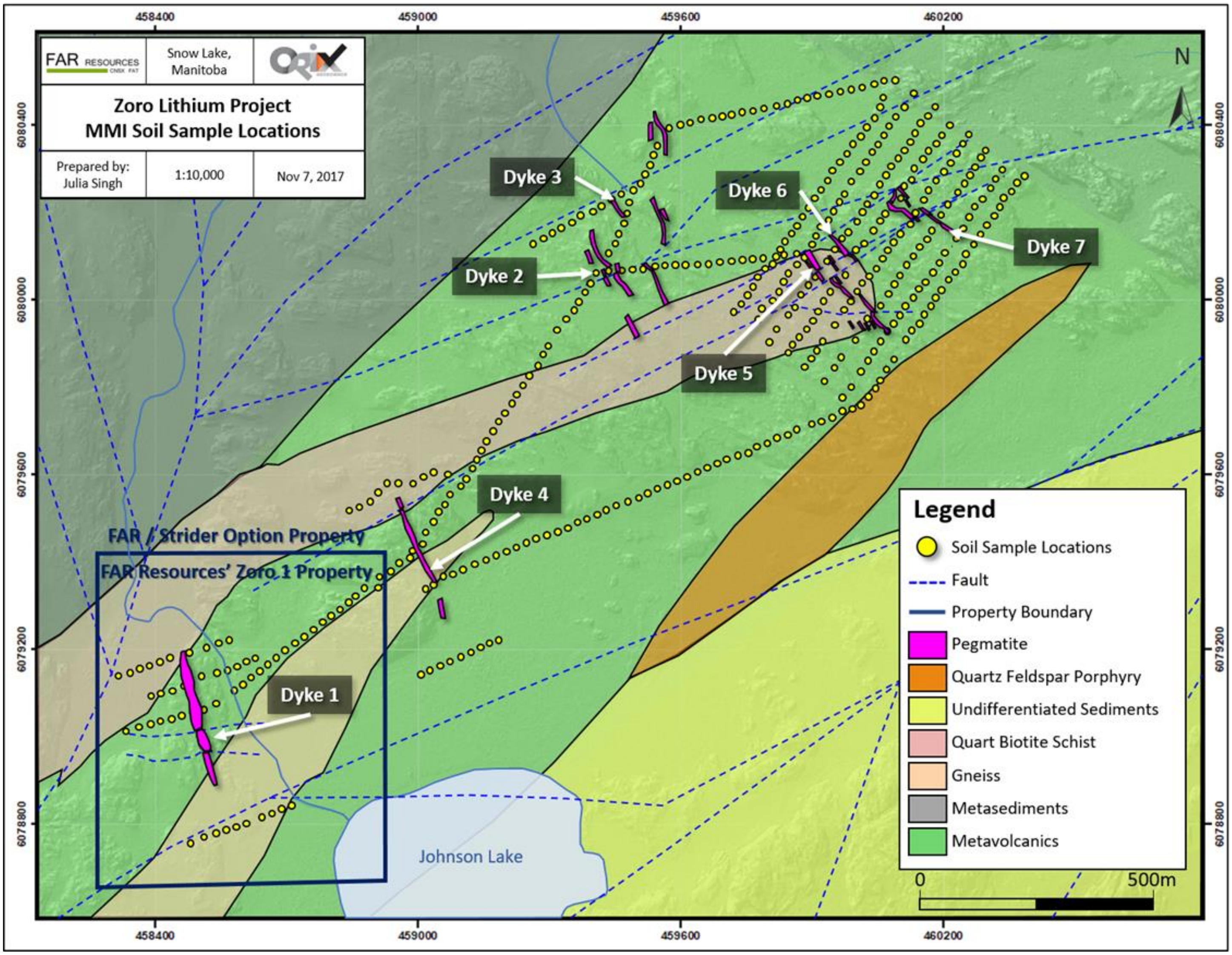


Figure 9.4.1: Sample location map for soil samples collected to assess areas of overburden cover for buried pegmatite dykes and areas of possible extensions of known pegmatite dykes. The base for this figure is compiled from historic and recent geological mapping.

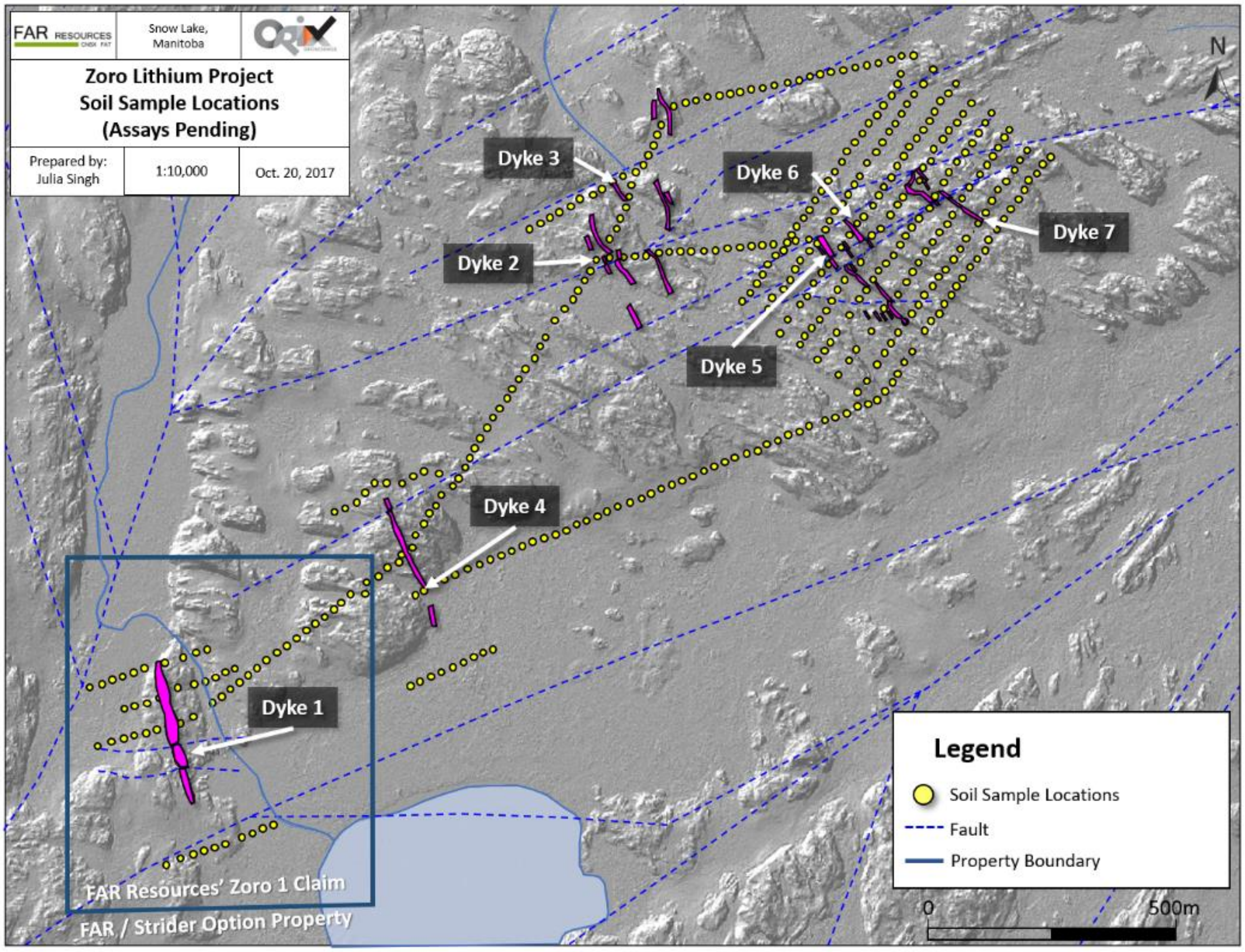


Figure 9.4.2: Sample location map for soil samples collected to assess areas of overburden cover for buried pegmatite dykes and areas of possible extensions of known pegmatite dykes. The base for this figure is from recent LIDAR surveys.

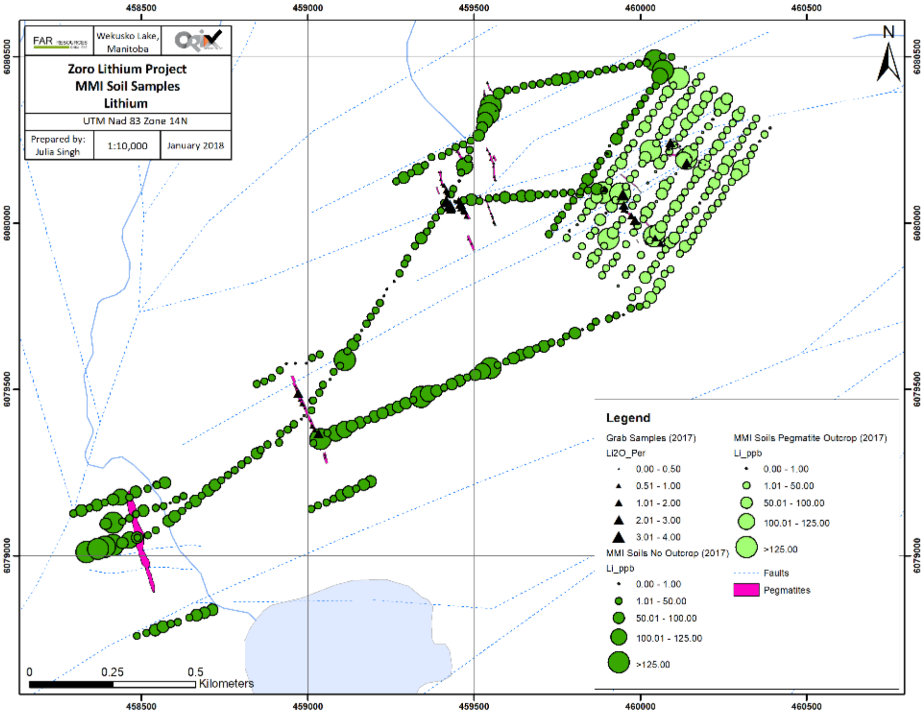


Figure 9.4.3: Bubble plot depicting results for lithium in Mobile Metal Ion soil geochemical data, Zoro Lithium Project. Rock chip assay data depicted as black triangles. Soil data presented for areas of no outcrop and in areas of known lithium mineralization.

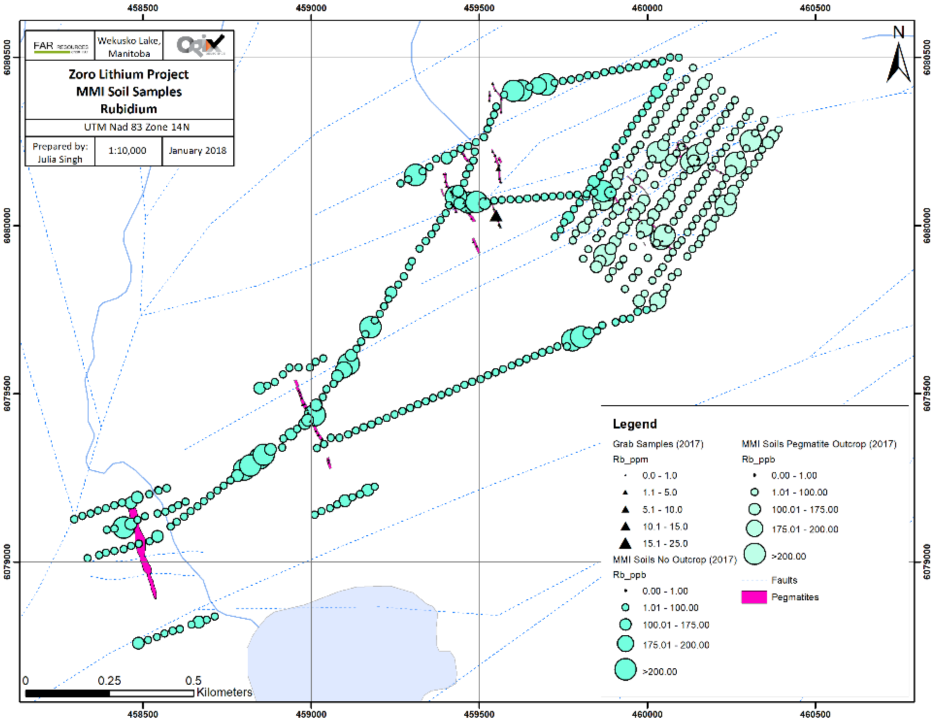


Figure 9.4.4: Bubble plot depicting results for rubidium in Mobile Metal Ion soil geochemical data, Zoro Lithium Project. Rock chip assay data depicted as black triangles. Soil data presented for areas of no outcrop and in areas of known lithium and tantalum mineralization.

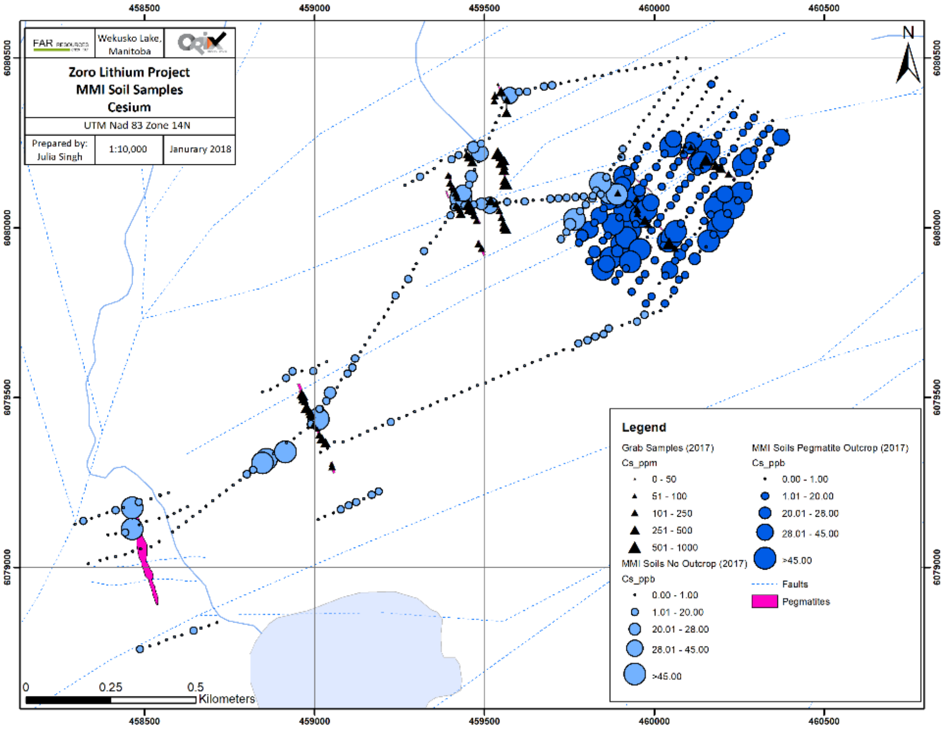


Figure 9.4.5: Bubble plot depicting results for cesium in Mobile Metal Ion soil geochemical data, Zoro Lithium Project. Rock chip assay data depicted as black triangles. Soil data presented for areas of no outcrop and in areas of known lithium-tantalum mineralization.

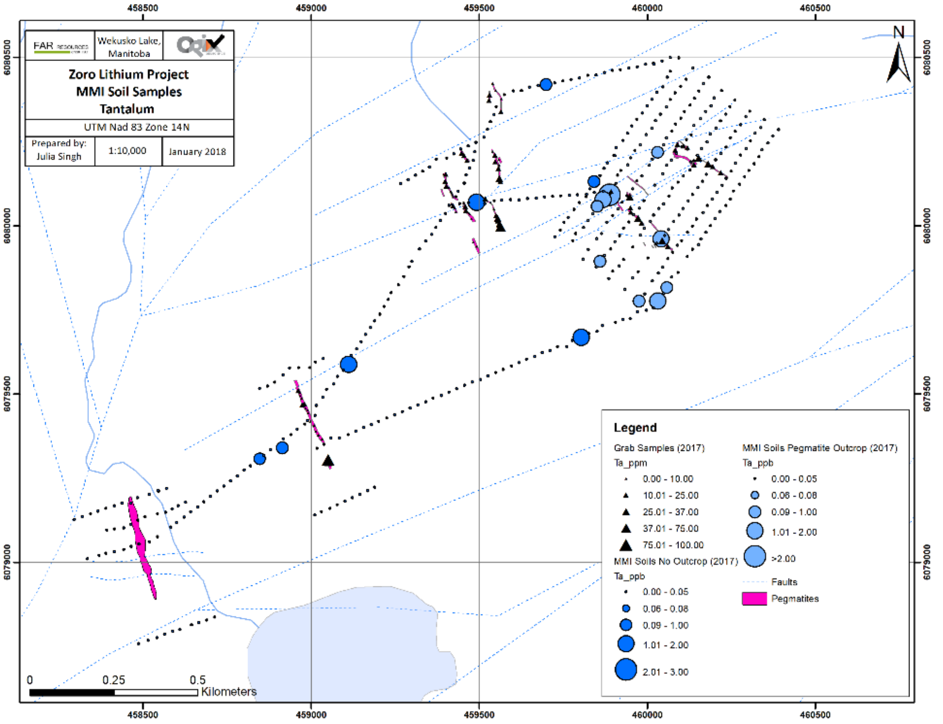


Figure 9.4.6: Bubble plot depicting results for tantalum in Mobile Metal Ion soil geochemical data, Zoro Lithium Project. Rock chip assay data depicted as black triangles. Soil data presented for areas of no outcrop and in areas of known lithium and tantalum mineralization.

9.3 DATA QUALITY

9.3.1 *ROCK SAMPLES*

Channel and rock chip samples acquired from the Zoro 1 pegmatite were analyzed at Activation Laboratories (ACTLABS), Ancaster (Ontario). Soil Samples

10.0 DRILLING

10.1 FAR RESOURCES DRILLING

After the historical drillhole locations were determined, an excel drillhole database was compiled. All 78 historical PDF drill logs were entered into the database by hand. All relevant data was captured including survey data, major and minor lithologies, alteration, pegmatite mineral composition, structure, and assay data when available. A legend was created and rock codes were assigned. Lithologies were consolidated when necessary to model the dykes in 3 dimensions.

Historical drillholes that intersected Dyke 1 were de-surveyed in Studio EM (Datamine). Seventeen cross-sections (spaced at 25m intervals perpendicular to Dyke 1) and 7 plan maps (spaced at 50m elevation intervals below surface) were created for interpretation. Interpretations of the pegmatite, gneiss, schist, metavolcanics, and overburden were done by hand drawing. These interpretations were then scanned, digitized in Autocad, and wireframed in Studio EM. A model of the Dyke 1 pegmatite was created from the cross-sections and plan maps, which served as the basis for Far Resources' drill targeting.

Between November fall of 2016 and the November of 2017, Far Resources conducted three small drilling programs on Dyke 1 to validate historical drilling results and to test the pegmatite farther along strike and at depth. A total of 19 drillholes summing 2,920.4m were drilled through these three exploration phases focused exclusively on Dyke 1. In the Winter of 2018, Far Resources conducted a larger drilling program which not only continued the definition of Dyke 1 but also began drill testing other lithium bearing dykes on the property (Dykes 2, 4, 5 and 7) as well as a strong lithium MMI anomaly. Figure 10.2.1 displays Far Resources collar locations at Dyke 1. Figure 10.2.2 displays Far Resources 2018 drilling program on multiple lithium bearing dykes (phase 4).

Gogal Air Services located at 494 Lakeshore Drive., Snow Lake, Manitoba provided helicopter support and supplied the core shack, equipment, and core storage for Far Resources drilling programs. Jake Ziehlke of Strider Resources based in Oakbank Manitoba, performed the ground truthing and drill pad cutting and Richard Stoltz based in Snow Lake Manitoba performed the core cutting and aided in drill pad cutting. Mark Fedikow (Independent consultant, Phase 1), Mike Kilbourne (Orix Geoscience, Phase 2), Chris Watters (Orix Geoscience, Phase 3) and Paul Nagerl (Orix Geoscience, Phase 4) were the project geologists on site and were responsible for spotting drillholes, drill site inspections, logging core, sampling, and ensuring that samples were properly bagged and shipped to Activations Laboratory, an ISO accredited laboratory. Westcore Drilling performed drilling for all three phases. A single drill rig was used throughout the programs and water was supplied by pump and hose sourced from several nearby water bodies including a local creek which runs a few meters east of Dyke 1. This local creek feeds into Johnson Lake approximately 200m south of Dyke 1 outcrops. Orix Geoscience Inc. performed data entry, database management, geological interpretations, 3D modeling, and drill targeting. Orix continued support of the drill programs by completing drillhole status update documents, continuously updating the database, and performing QAQC checks on the drilling sample programs.

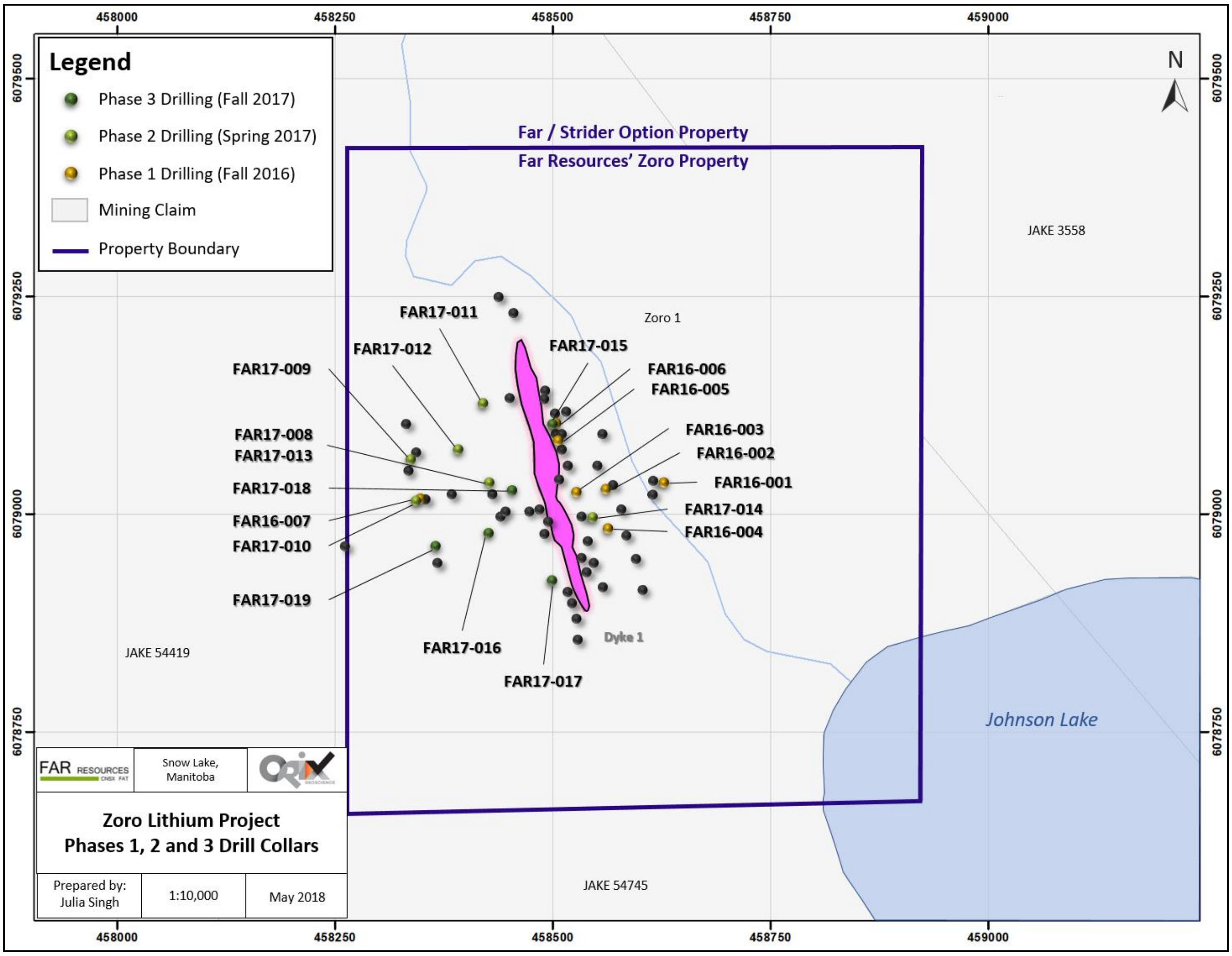


Figure 10.2.1: Phases 1, 2 and 3 drill collar locations at Dyke 1.

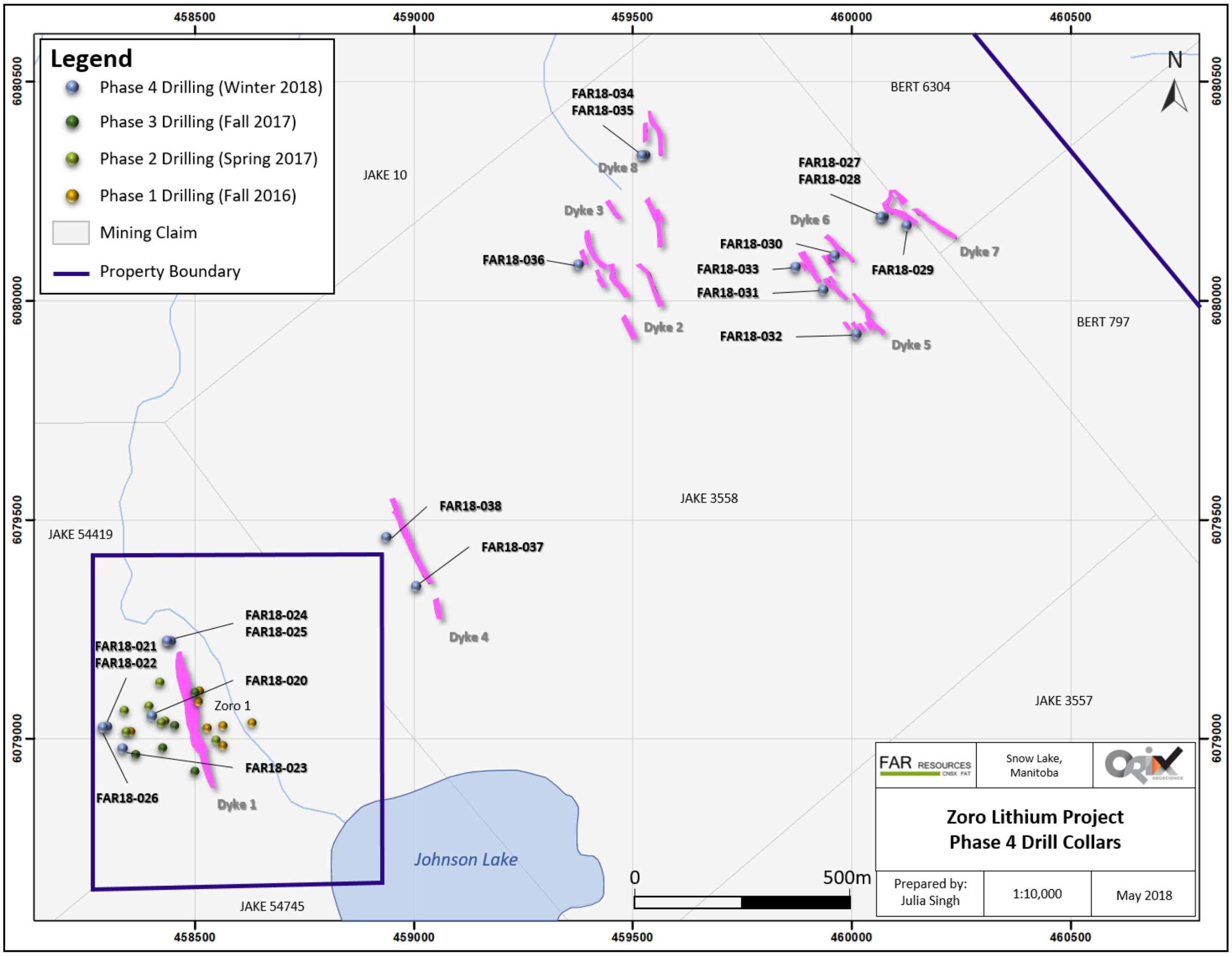


Figure 10.2.2: Phase 4 drill collar locations on multiple lithium bearing dykes.

10.1.1 PHASE 1 DRILLING

Phase 1 drilling program occurred in the November of 2016 and comprised 7 drillholes totaling 1,142.0 m of NQ core. The program was designed to validate results from historical drilling by attempting to twin holes that intersected wide zones of pegmatite and where good visual spodumene was mentioned. Historical holes zl-56-005 / 013 / 017 and 021 were twinned to validate the historical grade composites and three other holes targeted near other significant intersections.

A total of 143 drill core samples of pegmatite were cut and sent for assay. A summary of the drillholes and significant assay results are listed in Table 10.1.5. Highlighted results of the program include 1.1% Li$_2$O over 23.4 m in FAR16-007, which was successful in twinning historical hole zl-56-021. Furthermore, although FAR16-001 failed to intersect a similar wide zone as historical target zl-56-017 / 078, it intersected 0.117% Ta_2O_5, the first recorded occurrence of tantalite at Dyke 1. Phase 1 drilling information including all significant assay results are listed in Table 10.2.1. Phase 1 drilling was overall successful in validating historical results and prompted drilling to be planned for 2017.

Table 10.1.5: Far Resources Phase 1 Drilling.

BHID	UTM E	UTM N	Azimuth	Dip	EOH (m)	Target	Pegmatite Intersections (m)	Significant Assay Results
FAR16-001	458627	6079036	245.9°	-55.8°	323.00	Twin zl-56-017 / 078	202.30-203.84 206.14-208.97 224.08-229.00 254.00-255.64 290.57-292.00	1.5% Li_2O over 1.6m 0.1% Ta_2O_5 over 1.6m
FAR16-002	458562	6079028	252.9°	-57.7°	170.00	Twin zl-56-013	118.00-123.30 123.30-134.36 146.10-163.74 167.20-170.00	1.1% Li_2O over 8.6m
FAR16-003	458527	6079024	250.5°	-50.1°	92.00	Twin zl-56-005	48.84-73.07	1.1% Li_2O over 11.1m
FAR16-004	458563	6078983	249.8°	-48.9°	116.00	Testing zl-56-007 intercept	82.14-94.18	0.6% Li_2O (Max value)
FAR16-005	458507	6079083	251.9°	-45.1°	71.00	Testing between zl-56-001 and 071	21.31-47.90	1.3% Li_2O (Max value)
FAR16-006	458500	6079103	252.0°	-45.1°	70.91	Testing between zl-56-069	19.28-43.24 45.61-46.61 67.73-69.15	0.8% Li_2O over 11.0m 1.3% Li_2O over 1.4m
FAR16-007	458345	6079015	73.0°	-57.0°	275.00	Twin zl-56-021	186.18-213.81 251.81-258.11	1.1% Li_2O over 23.4m 1.2% Li_2O over 4.1m

10.1.2 PHASE 2 DRILLING

Based on successful results from Phase 1, a second small drilling program was planned for the March of 2017. The two-week Phase 2 drilling program commenced in March of 2017 with 7 drillholes totaling 1,088.0 m of NQ core. Drillhole information from Phase 1 was used to refine the Dyke 1 pegmatite model. Although historical assay results were lacking, spodumene occurrence was modeled to attempt to identify mineralized zoning within the pegmatite which further refined the drill targets. The second program was designed to test Dyke 1 within 150m from surface along strike. The holes were designed to infill previously untested areas of the pegmatite (approximately 30m to 50m from historical and Phase 1 drill intersections).

A total of 167 drill core samples of pegmatite were cut and sent for assay. Highlighted results for the program include 1.2% Li_2O over 38.3m in FAR17-010. Phase 2 drilling information including all significant assay results are listed in Table 10.2.1. The Phase 2 program was successful in intersecting spodumene-bearing pegmatite and confirming the Dyke 1 model, as well as indicating a pinch out of the Dyke toward its northern exposed boundary when in contact with the gneissic unit.

Table 10.2.1: Far Resources Phase 2 drilling.

BHID	UTM E	UTM N	Azimuth	Dip	EOH (m)	Target	Pegmatite Intersections (m)	Significant Assay Results
FAR17-008	458426	6079037	75.0°	-50.0°	161.00	Testing near zl-56-020	101.30-117.70 141.83-149.77	1.1% Li_2O over 2.4m
FAR17-009	458338	6079063	82.0°	-65.0°	264.20	Testing near zl-56-018	165.18-166.97	No sig. mineralization
FAR17-010	458343	6079014	68.0°	-48.0°	56.00	Testing near zl-56-029	162.92-205.45 207.00-216.62 230.10-230.91	1.2% Li_2O over 38.3m 2.3% Li_2O over 4.6m 2.6% Li_2O over 2.1m 1.4% Li_2O over 7.7m
FAR17-011	458419	6079129	69.0°	-58.0°	256.00	Testing near zl-56-003	43.63-45.31 46.74-48.00	1.3% Li_2O over 1.3m
FAR17-012	458392	6079075	250.0°	-45.0°	155.00	Testing near zl-56-071	103.40-115.55 123.06-124.44 125.44-133.23	1.7% Li_2O over 10.7m 4.1% Li_2O over 0.4m 2.1% Li_2O over 5.1m
FAR17-013	458426	6079037	75.0°	-51.0°	114.30	Testing near zl-56-004	71.56-82.30	1.0% Li_2O over 1.7m
FAR17-014	458545	6078996	78.0°	-50.0°	86.00	Testing near zl-56-006	57.95-65.91	No significant mineralization

10.1.3 PHASE 3 DRILLING

In September of 2017, another drill program was executed to follow up on drilling results from Phases 1 and 2. Phase 3 drilling comprised 5 drillholes totaling 710.0 m of NQ core. Based on Far Resources previous drill programs the pegmatite Dyke 1 model was revised and a preliminary internal block model was created to be used for identification of additional target areas within Dyke 1. The purpose of the Phase 3 drill program was to increase grade and tonnage of the internal block model moving toward a 43-101 resource estimate by infilling gaps in the drilling, to a distribution of 25 m radius reportage.

A total of 207 drill core samples of pegmatite were cut and sent for assay. Highlighted results for the program include 1.4% Li_2O over 20.6 m in FAR17-018 and 1.2% Li_2O over 12.4 m in FAR17-019. Phase 3 drilling information including all significant assay results are listed in Table 10.2.2. The Phase 3 program was successful in intersecting spodumene-bearing pegmatite and further confirming the Dyke 1 model.

Table 10.2.2: Far Resources Phase 3 drilling.

BHID	UTM E	UTM N	Azimuth	Dip	EOH (m)	Target	Pegmatite Intersections (m)	Significant Assay Results
FAR17-015	458499	6079104	250.0°	-59.2°	104.00	Infill	40.50-43.71 43.71-48.38 48.38-52.92 52.92-53.82 57.76-60.44 60.44-64.00 65.72-67.17 68.00-86.81	1.1% Li_2O over 1.2m 0.8% Li_2O over 3.3m 1.4% Li_2O over 1.0m 1.0% Li_2O over 1.5m 1.0% Li_2O over 6.1m 1.0% Li_2O over 1.1m
FAR17-016	458426	6078978	73.8°	-45.0°	132.00	Infill	95.81-108.50 108.5-118.18	0.7% Li_2O over 0.9m 0.6% Li_2O over 1.0m
FAR17-017	458499	6078923	80.4°	-65.0°	104.00	Infill	52.40-58.68	0.5% Li_2O over 3.0m
FAR17-018	458452	6079028	74.0°	-63.0°	110.00	Infill	54.05-58.75 58.75-80.55 80.55-93.86	1.4% Li_2O over 20.5m 2.2% Li_2O over 4.0m 3.1% Li_2O over 1.0m
FAR17-019	458365	6078963	68.5°	-60.0°	260.00	Infill	206.83-226.60	1.2% Li_2O over 12.4m

10.1.4 PHASE 4 DRILLING

In January of 2018, Far Resources initiated a Phase 4 drilling program. Phase 4 comprised 19 drillholes totaling 2,472.0 m, establishing it the largest drilling program designed by Far Resources to date. Phase 4 drilling consisted of multiple targets, including further definition of Dyke 1 and its possible northern extension, modern drill testing of Dykes 2, 4, 5 and 7, and drill testing of a strong MMI anomaly.

A total of 306 drill core samples of pegmatite were cut and sent for assay. Highlighted results from the program are detailed below in relation to the target of interest. Phase 4 drilling information including all significant assay results are listed in Table 10.2.3. Phase 4 drilling was successful intersecting Dyke 1 deeper than previous programs and successfully drill tested a strong MMI anomaly

DYKE 1

Five drillholes were designed to further define Dyke 1 at previously untested depths and to infill the north-central portion of the dyke model. Highlighted results include 1.8% Li_2O over 15.7 m in FAR18-020 and two significant intersections of 1.6% Li_2O over 5.0 m and 0.7% Li_2O over 16.2 m in FAR18-023. FAR18-022 was designed to test below FAR18-021 from the same collar location, however, the drillhole was cancelled before intersecting the interpreted target interval, interpreted to be related to the pinching nature of the pegmatite in FAR18-021. The drilling results of Dyke 1 confirmed continuation of pegmatite and spodumene mineralization at depth.

Two drillholes, FAR18-024 and FAR18-025, were designed to target the northern extension of Dyke 1 previously identified in a historical trench map reported by Green Bay Uranium Group (assessment report #93562). Far Resources personnel initially confirmed a small surface exposure of pegmatite near the target area (internally referred to as the "Dogleg"). However, neither drillhole intersected pegmatite. Additional field work is suggested to further define this target for future drilling.

DYKE 7

Three drillholes were designed to target Dyke 7 based on a review of historical drilling data, as well as results from the 2017 preliminary reconnaissance program in which surface grab samples resulted up to 2.6% Li_2O. FAR18-027 and FAR18-028 tested a vertical fan of the northern extent of known surface exposure of spodumene in of Dyke 7. FAR18-029 targeted below 15-20% surface-exposed spodumene proximal to historical hole zl-56-056, which intersected 5.2m of pegmatite. All three drillholes intersected multiple pegmatite dykes with minor spodumene. Highlighted results for Dyke 7 include 0.6% Li_2O over 1.5m in FAR18-028.

DYKE 5

Four drillholes were designed to target Dyke 5, based on a review of historical drilling data, results from the 2017 preliminary reconnaissance program in which surface grab samples resulted up to 3.9% Li_2O, and results from the 2016 surface chip sample program in which resulted up to 6.4% Li_2O. FAR18-030 and FAR18-031 targeted below high-grade surface results; both were successful intersecting lithium mineralization at depth, although pegmatite widths were narrower than observed at surface. FAR18-032 targeted below the southern surface exposure of Dyke 5 and test an MMI anomaly which resulted 313 ppb Li. The drillhole intersected the highlighted result of 1.6% Li_2O over 1.3m. FAR18-033 targeted the northernmost surface exposure of Dyke 5 and an MMI anomaly which resulted 264 ppb Li; this drillhole did not intersect significant mineralization. Spodumene mineralization was identified in all four drillholes, however, the pegmatites measured to be narrower at depth than what was observed at surface.

MMI SOIL ANOMALY – NEW DISCOVERY

Two drillholes were designed to target an anomaly from the Summer 2017 MMI survey. This was the first drilling to assess a target not exposed at surface. MMI soil surveys are used as an effective tool for lithium exploration on buried or undiscovered lithium-bearing pegmatites. The targeted anomaly is identified as two 25m spaced samples along an east-west grid which both returned high MMI soil results (379ppb and 971ppb). The drillholes were designed as a vertical fan to intersect midway between the two anomalous soil samples; success on the shallow hole triggered the drilling of a steeper hole. Both drillholes were successful in intersecting lithium-bearing pegmatite, therefore identifying a new discovery. Highlighted results include 0.7% Li_2O over 4.9m (including 1.3% Li_2O over 0.9 m) in FAR18-034 and multiple high-grade intersections, including 1.1% Li_2O over 12.3 m, in FAR18-035. This newly discovered zone, identified as “Dyke 8”, will be a focus for future drilling on the property (Figure 10.2.3).

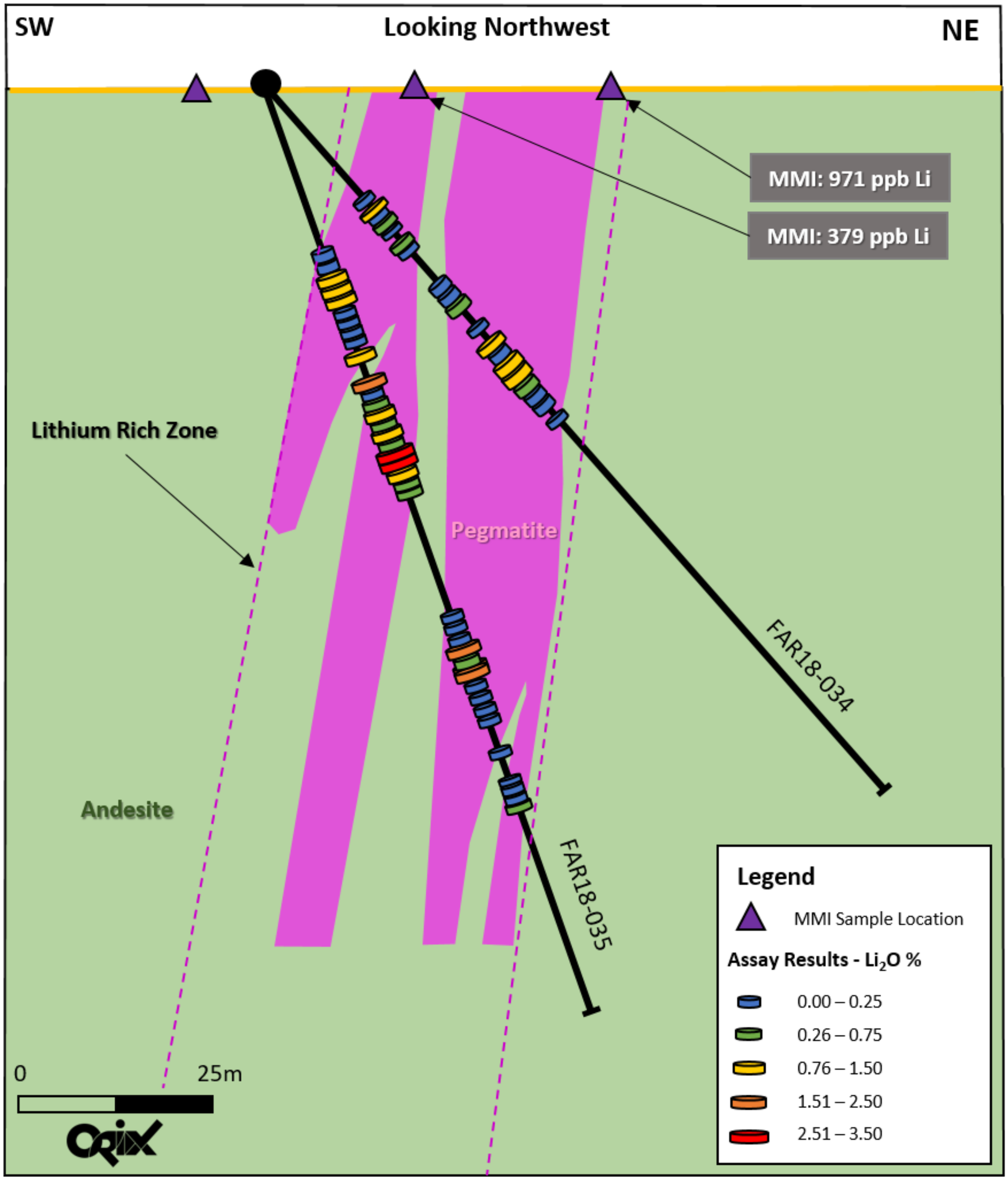


Figure 10.2.3: Drill results from testing a Mobile Metal Ions Li soil geochemical anomaly.

DYKES 2 AND 4

Three drillholes were designed to target Dykes 2 and Dykes 4 to corroborate historical drilling results. FAR18-036 tested Dyke 2, targeting a historical drillhole intersection of 9.2 m of pegmatite in zl-56-038. Although FAR18-036 intersected 2 m of pegmatite, there was no significant mineralization.

FAR18-037 and FAR18-038 targeted the north and south extensions (respectively) of Dyke 4, proximal to historical intersections. Highlighted results include 0.5% Li_2O over 0.8m in FAR18-037. FAR18-038 failed to intersect significant mineralization. Further surface work is suggested to better define future drilling targets on Dykes 2 and 4.

Table 10.2.3: Far Resources Phase 4 drilling.

BHID	UTM E	UTM N	Azimuth	Dip	EOH (m)	Target	Pegmatite Intersections (m)	Significant Assay Results
FAR18-020	458400	6079051	73.3	-65.0	209.0	Dyke 1	106.22-123.20 145.08-163.73	1.8% Li_2O over 15.7m
FAR18-021	458297	6079026	73.0	-51.0	305.0	Dyke 1	210.00-214.39 216.65-216.83 228.9-229.21 260.72-263.02	0.5% Li_2O over 1.4m
FAR18-022	458297	6079026	73.0	-61.0	92.0	Dyke 1	Drilling Cancelled	NA
FAR18-023	458332	6078979	70.7	-62.0	326.0	Dyke 1	193.66-193.87 227.76-260.70 299.07-301.21 306.97-307.76 308.11-308.25	0.7% Li_2O over 16.2m 1.6% Li_2O over 5.0m
FAR18-024	458442	6079225	112.6	-45.0	74.0	Dogleg	None	NA
FAR18-025	458442	6079225	112.6	-60.0	86.0	Dogleg	None	NA
FAR18-026	458289	6079025	80.8	-65.0	350.0	Dyke 1	265.92-282.36	0.5% Li_2O over 3.0m
FAR18-027	460071	6080190	31.4	-50.0	101.0	Dyke 7	8.52-12.95 13.53-13.95 53.40-53.78 66.40-66.78 67.66-68.30	No significant mineralization
FAR18-028	460071	6080190	31.4	-50.0	128.0	Dyke 7	8.40-13.50	0.6% Li_2O over 1.5m
FAR18-029	460132	6080168	3.9	-50.0	71.0	Dyke 7	4.00-4.90 5.51-6.75	No significant mineralization
FAR18-030	459959	6080102	221.4	-54.0	74.0	Dyke 5	44.34-47.00 47.46-48.20	1.2% Li_2O over 1.0m
FAR18-031	459934	6080026	41.0	-52.0	71.0	Dyke 5	11.90-12.10 17.70-19.84 24.40-24.88 59.00-59.20	1.2% Li_2O over 1.3m
FAR18-032	460012	6979926	40.6	-52.0	101.0	Dyke 5	16.00-17.20 25.34-25.59 34.90-38.52 68.28-68.77 82.52-85.62 86.51-87.23	1.6% Li_2O over 1.3m
FAR18-033	459871	6080076	43.6	-50.0	86.0	Dyke 5	24.22-24.70 27.40-27.77 69.35-71.00	No significant mineralization
FAR18-034	459527	6080332	60.3	-48.0	85.0	Dyke 8*	14.57-19.54 24.58-40.32	0.7% Li_2O over 4.9m
FAR18-035	459527	6080332	60.3	-70.0	89.0	Dyke 8*	15.34-22.10 24.44-39.21 51.22-64.10 67.00-69.25 73.32-73.47	1.2% Li_2O over 4.4m 1.1% Li_2O over 12.3m 1.5% Li_2O over 2.2m

FAR18-036	459378	6080081	61.9	-45.0	74.0	Dyke 2	46.10-48.20	No significant mineralization
FAR18-037	459003	6079348	67.1	-45.0	74.0	Dyke 4	30.20-30.56 36.00-36.10 45.69-48.78	0.5% Li_2O over 0.8m
FAR18-038	458936	6079459	67.4	-45.0	76.0	Dyke 4	57.63-58.45	No significant mineralization

***Dyke 8 newly discovered in 2018 drill program.**

10.2 CORE HANDLING, SAMPLING METHODS, AND APPROACH

10.2.1 HISTORICAL WORK

Historically, work on the project area in 1956/1957 was not reported in a method that was not done to CIMM standards and guidelines. Core storage was situated on the property itself and has since deteriorated with historical core being lost. Acid dip tests for survey data were recorded in the drill logs but did not indicate azimuth deviation, therefore it is uncertain where the drillholes deviated at depth. Furthermore, the sampling protocol did not include insertion of QAQC reference materials (blanks and standards) into the sample stream. The sampling methods are not described in the drill logs or related assessment report. Ultimately, the historical drilling was used in a general capacity to provide a base understanding of the pegmatite dyke, mineralization, and local geology. After Far Resources completed sufficient drilling on Dyke 1 to confirm the model (Phase 3), the historical drillholes were removed from the wireframes and internal block model.

10.2.2 FAR RESOURCES SURVEY METHOD

Far Resources used a Reflex Multi-shot tool to survey the drillholes. Survey readings were taken at 30-50m intervals downhole, recording the azimuth, dip, temperature, and magnetics for each reading. All survey data was included in the database. Survey readings aided in calculating expected deviation for subsequent drilling programs.

Magnetic interference occurs on the west side of Dyke 1, therefore, an APS unit (Azimuth Pointing System) was added in Phase 3 to record the collar location, azimuth, and dip at surface in attempt to avoid magnetic interference. The APS unit is not affected by local magnetic interference and may also be used after drilling is complete to verify the starting azimuth and dip. Previous phases of drilling utilized a Garmin GPS unit and orienteering for spotting drillholes and for recording a final collar site after drilling had been completed.

10.2.3 CORE HANDLING, LOGGING AND SAMPLING METHODS

The core was first retrieved at the drill site as the drill helper removes the core from the core tube into the drill box. The drillers were provided with a drill status sheet prior to the start of each hole to ensure that the boxes were labelled with the appropriate hole number. Should an error be noted by the project geologist on site, the box numbers were corrected immediately. A wooden block was inserted after each 3m core run recorded with the meterage down hole. The core remains in the custody of the drillers until it is flown back by helicopter sling after each 12-hour drilling shift (weather permitted), where it was immediately inspected by the project geologist. The core, when not being attended to, is stored at Gogal Air services helicopter hanger on core racks and pallets. The core was then transported into the core logging facility where it was teched, logged, and sampled.

Core was laid out on core tables to be initially inspected for correct meter blocks and drillhole ID. The project geologist was responsible for recording RQD measurements and ensuring that the core is in continuous correct order. In Phase 3, a logging and sampling procedure document was created and implemented for subsequent programs. Logging was recorded in an excel template detailing intervals with associated major lithology, minor lithology, pegmatite minerals, structure, alteration, and samples. Generally, major lithologies were considered any unit greater than 2m and minor lithologies less than 2m. Special attention was placed on logging the pegmatite intervals, ensuring that mineralogical zoning was noted.

Once the geological information for each hole was recorded, the project geologist identified core to be sampled. Sample number identifiers were written on the core in red grease pencil, marking each sample with a starting arrow and ending arrow to indicate length. A corresponding sample tag was filled out for each sample, including the “from” and “to” intervals and a brief description was recorded in the drill logs. One third of the sample tag remained in the sample book to be retained as reference, a second tag was placed underneath the remaining portion of core interval to be sampled, and a third tag was placed in the sample bag with the sample portion. Sample intervals did not cross lithological boundaries significant alteration zones, or mineralogical zoning of the pegmatite; samples were selected of homogenous content. Sample interval lengths were greater than 0.3m but less than 1.5m. Shoulder samples of the host rock were not required as spodumene mineralization does not occur outside of pegmatites. QAQC controls were inserted into the sample stream at regular intervals and are discussed in detail in section 11.

11.0 SAMPLING PREPARATION, ANALYSES AND SECURITY

11.1 SAMPLE COLLECTION AND SECURITY

Outcrop samples for the analysis of lithium and other elements were collected depending on the provenance of the sample. Channel samples were cut from available pegmatite exposures in mucked out and washed trenches on the Zoro Lithium Project. Channels were cut with a STIHL rock saw and were 6 cm wide and 6 cm deep on average. Sample length was variable owing to available exposure. In this manner 165 channel samples were acquired from 16 trenches with an additional 5 representative rock chips collected where channel samples could not be obtained.

Samples were removed from the channel using a rock pick and a standard chisel, labelled and placed in plastic sample bags. At the end of each work day all samples were transported by helicopter in sealed rock pails to a locked storage facility in Snow Lake. All samples were stored in this manner until the end of the sampling project when all samples were shipped from Snow Lake via Gardewine Transport to ACTLABS in Ancaster (Ontario) for analysis.

Drill core samples were collected after washing and core logging. Sample intervals from drill core varied according to the presence of spodumene in the pegmatite. Once the project geologist had completed sampling procedures, the core cutter was responsible for sawing identified core in half. Core was bisected into two halves with a rock saw with one half of the core returned to the core box in the correct interval and the other half was placed in a sample bag. Each sample bag was labelled with the sample number corresponding to the sample tag and sealed using a zip tie. The project geologist then assembled the sample bags to ensure no samples were missed and placed into rice bags and sealed for shipment. The project geologist filled out a sample submittal form for each batch of samples. All samples were stored in a locked core logging facility located adjacent to Gogal Air Services (Snow Lake) compound until shipped. The sample batches were shipped by ground using Gardewine to Activation Laboratories in Ancaster, Ontario. Blanks, duplicate samples and when available lithium standards were included in each sample batch forwarded to the laboratory.There is a standard client relationship that exists between the issuer and the analytical laboratories.

It is the author’s opinion that quality assurance and quality control methods adopted for both rock and soil sample analyses are adequate for the purposes of undertaking an exploration program.

Soil samples collected for Mobil Metal Ions analysis were collected with a Dutch auger. Proper collection procedures are vital to the success of an MMI™ orientation survey. Four samples must be taken from each hand dug pit to obtain a broad cross section of data sufficient to capture the optimal sampling depth. First, the interface or depth to begin sampling must be located. Typically, this interface is defined by the top of the humified organic layer lying just below the stratum containing leaf litter and organic material with visible structure (i.e. decomposing leaves, bark, twigs and peat). Below this interface, four depths are marked out (0-10 cm, 10-20 cm, 20-30cm, and 30-40 cm) and samples are carefully taken from each, beginning at the bottom and working upwards.

Samples are taken from the bottom to the top of the pit to minimize the contamination of lower samples with soil from higher in the profile. Using a plastic or vinyl scoop, a cross section of material was taken from each layer, ensuring each sample contains 200-300g of soil and is placed in a snap-seal plastic bag (e.g. ZIPLOC). Samples are not dried or sieved, and no sample preparation is required other than ensuring the sample is not contaminated. MMI™ geochemistry measures metallic mobile ions in parts per billion (ppb) or subparts per billion. At these concentrations contamination can easily overwhelm metal ion counts and strict adherence to survey cleanliness is required to ensure accuracy and repeatability. Cleanliness practices that must be followed during an MMI™ Orientation Survey include: (i) Sampling equipment brushed clean and flushed with soil from the new sample site before digging to eliminate residue from previous samples. (ii) During sample collection and handling, no jewelry (watches, rings, bracelets, chains etc.) can be worn, as this can be a major source of contamination. (iii) Sampling pits must be excavated with “clean” shovels that are paint and rust free. (iv) Vertical pit surfaces must be scraped clean to remove any debris and potential contaminants. (v) Sampling equipment must be made of plastic or vinyl only unless samples are collected with a Dutch auger.

All soil sample bags were double checked by the Far Resources sampler prior to shipping. A sample shipment form was filled out and the samples were stored in a locked facility next to the

Gogal Air Services (Snow Lake) compound until shipped. Blanks, duplicate field and laboratory samples and standard reference materials were included with each batch of samples prior to shipping to the laboratory.

11.2 HISTORIC SAMPLING METHODS

The specifics of drill core and sample collection including sample footages and lengths for historical drill holes from the Zoro Lithium Project are not reported in the Manitoba Government Assessment Files or other historic information sources. The assay sample intervals for gold are presented in Table 7.3 and indicate highly variable sample widths. It is not known what sampling methods were used. A description of the approach to outcrop and pit/trench sampling is available in Cancelled Assessment File 93562 and was produced by Dr. R. Banfield, consultant to Green Bay Mining and Exploration Ltd. The Banfield report (*cf.* C.A.F. 95362) noted the following factors as critical in designing an appropriate sampling scheme for the Zoro 1 pegmatite:

> *“The entire width of the pegmatite dyke required sampling. The spodumene crystals varied between “a fraction of an inch to 18 inches long”. The orientation of the spodumene crystals is mostly irregular but where there is orientation, the long axis of the crystals is parallel to the walls of the dyke.”*

Sample collection was aided by a gasoline powered portable crusher mounted on skids. Trenches approximately “2 feet wide” were blasted into the dyke and extended from wall to wall of the pegmatite at right angles to the strike of the dyke. The top one foot of rock was discarded to avoid the effects of weathering. The trench was then deepened for an additional two feet. Trenches were established every twenty-five feet. The sample that was collected for assay was two feet wide, two feet deep which at 12 cubic feet to the ton would represent a third of a ton per lineal foot of trench. It was recommended that the sample lengths be equal to five lineal feet of trench which would yield about one ton of sample.

Blasting mats were utilized to avoid scatter and loss of material and all fines were reserved and included with coarser fractions for assay. Rock fragments were crushed to maximum diameters of one inch. Subsequent to sizing the material is shoveled into a cone-shaped pile on a metal or wooden platform with subsequent material added to the top of the cone. The sample was then flattened to a thickness of 1 foot and spread out to form a ring with no material in the center of the ring. The central cone was then re-established by shoveling material back into the center of the ring. This procedure was repeated twice and then the cone is flattened to a thickness of one foot and divided into four quadrants. Two of the four quadrants were then “coned” again using the above procedure and a second set of quadrants produced. This procedure was repeated until a total of 40 pounds of sample remained. This sample was once again quartered, divided into two halves and twenty pounds were bagged, labelled and sent to the analytical facility for assay. The remaining half was archived for future use.

11.3 SAMPLE PREPARATION

11.3.1 ROCK

Rock samples, including channel (Figure 11.3.1) and representative rock chip, were prepared for analysis by crushing to a nominal minus 10 mesh (1.7 mm), mechanically splitting (riffle) to obtain a representative sample and then pulverizing to at least 95% minus 150 mesh (106 microns). Cleaner sand is used between each sample and the quality of crushing and pulverization is routinely checked as part of the Activation Laboratories quality assurance program. Pulverizing is done in a mild steel vessel with possible Fe contamination of up to 0.2%.

Figure 11.3.1: Channel cut for sampling of the Zoro 1 spodumene pegmatite.

11.3.2 SOIL

There is no sample preparation for soil samples collected for analysis using Mobil Metal Ion Technology.

11.4 SAMPLE ANALYSIS-ROCK SAMPLES

All drill core and rock chip samples were analyzed with Activation Laboratories (Actlabs) in Ancaster Ontario which is an ISO 17025 accredited laboratory issued by the Standards Council of Canada (SCC). The samples underwent “Ultratrace7” (UT-7) analysis. This analytical approach combines a Sodium Peroxide (Na_2O_2) fusion with ICP/OES and ICP/MS finish. All metals are solubilized. A brief description of the analytical methodology is given here:

- **ICP/MS:** Fused samples are diluted and analyzed by Perkin Elmer Sciex ELAN 6000, 6100 or 9000 ICP/MS. Fused blank is run in triplicate for every 22 samples. Controls and standards fused with samples are run after the 22 samples. Fused duplicates are run every 10 samples. Instrument is recalibrated every 44 samples.
- **ICP/OES:** Samples are analyzed with a minimum of 10 certified reference materials for the required analyte, all prepared by sodium peroxide fusion. Every 10th sample is prepared and analyzed in duplicate; a blank is prepared every 30 samples and analyzed. Samples are analyzed using a Varian 735ES ICP or a Thermo 6500 ICAP. Results are reported in parts per million (ppm).

The reported elements and their detection limits are given in Table 11.4.1.

Table 11.4.1: Ultratrace-7 analysis: elements and detection Limits (ppm) unless otherwise indicated.

Element	Detection Limit	Reported By
Ag	13	ICP/MS
Al	0.01%	ICP
As	5	ICP/MS
B	10	ICP/MS
Ba	3	ICP/MS
Be	3	ICP/MS
Bi	2	ICP/MS
Ca	0.01%	ICP
Cd	2	ICP/MS
Ce	0.8	ICP/MS
Co	0.2	ICP/MS
Cr	30	ICP/MS
Cs	0.1	ICP/MS
Cu	2	ICP/MS
Dy	0.3	ICP/MS
Er	0.1	ICP/MS
Mo	1	ICP/MS
Nb	2.4	ICP/MS
Nd	0.4	ICP/MS
Ni	10	ICP/MS
P	0.005%	ICP
Pb	0.8	ICP/MS
Pr	0.1	ICP/MS
Rb	0.4	ICP/MS
S	0.01%	ICP
Sb	2	ICP/MS
Se	0.8	ICP/MS
Si	0.01%	ICP
Sm	0.1	ICP/MS
Sn	0.5	ICP/MS
Sr	3	ICP/MS
Ta	0.2	ICP/MS

Eu	0.1	ICP/MS
Fe	0.05%	ICP
Ga	0.2	ICP/MS
Ge	0.7	ICP/MS
Gd	0.1	ICP/MS
Hf	10	ICP/MS
Ho	0.2	ICP/MS
In	0.2	ICP/MS
K	0.1%	ICP
La	0.4	ICP/MS
Li	3	ICP/MS
Mg	0.01%	ICP
Mn	3	ICP/MS
Tb	0.1	ICP/MS
Te	6	ICP/MS
Th	0.1	ICP/MS
Ti	0.01%	ICP
Tl	0.1	ICP/MS
Tm	0.1	ICP/MS
U	0.1	ICP/MS
V	5	ICP/MS
W	0.7	ICP/MS
Y	0.1	ICP/MS
Yb	0.1	ICP/MS
Zn	25	ICP/MS

The Quality Control System at ACTLABS is accredited to international quality standards through the International Organization for Standardization /International Electrotechnical Commission (ISO/IEC) 17025 (ISO/IEC 17025 includes ISO 9001 and ISO 9002 specifications) with CAN-P-1758 (Forensics), CAN-P-1579 (Mineral Analysis) and CAN-P-1585 (Environmental) for specific registered tests by the SCC. The accreditation program includes ongoing audits which verify the Quality Assurance system and all applicable registered test methods. ACTLABS is also accredited by the National Environmental Laboratory Accreditation Conference (NELAC) program and Health Canada (Activation Laboratories website). Specific details are presented in Table 11.4.2.

Table 11.4.2: Activation Laboratories scope of accreditation and inspection (ACTLABS website).

Accrediting Organization	Scope of Accreditation
Standards Council of Canada (SCC) for International Standards Organization (ISO) 17025	Forensic Tests (CAN-P-1578) Mineral Analysis/Geological Tests (CAN-P-1579) Environmental Tests (CAN-P-1585) Chemical/Physical/Mechanical Tests
Health Canada	Establishment License for Pharmaceutical Testing (#101067-A and –B)
Food & Drug Administration (FDA) Registered and Inspected	Pharmaceutical Testing (Registration #3005494188)
Ontario Ministry of Agriculture and Food (OMAFRA)	Accredited Soil Analysis Laboratory (Agriculture)

ACTLABS' Quality System is accredited to international quality standards through the International Organization for Standardization /International Electrotechnical Commission (ISO/IEC) 17025 (ISO/IEC 17025 includes ISO 9001 and ISO 9002 specifications) with CAN-P-1578 (Forensics), CAN-P-1579 (Mineral Analysis) and CAN-P-1585 (Environmental) for specific registered tests by the SCC. The accreditation program includes ongoing audits which verify the

QA system and all applicable registered test methods. ACTLABS is also accredited by Health Canada.

The quality program at ACTLABS also includes the use of standards, analytical duplicates and blanks. Table 11.4.3 presents a comparison of the results of lithium analysis for international standards (measured and certified), analytical duplicates and replicate analyses of the analytical blank for assays from Dyke 1. Review of Table 11.4.3 indicates the Zoro lithium analyses are accurate, precise and have no laboratory-based contamination introduced into the samples as monitored by the method blank. Figure 11.4.1 presents a graphical representation of duplicate analyses.

Table 11.4.3: Summary of quality control data, Dyke 1 assays, Zoro Lithium Project.

Analyte Symbol Unit Symbol Detection Limit (ppm) Analysis Method	**Li ppm 3 FUS-MS-Na2O2**
Standards	
W-2a Measured	8
W-2a Certified	9.6
NCS DC70018 Measured	28
NCS DC70018 Certified	29
BIR-1a Measured	6
BIR-1a Certified	3.6
NCS DC70014 Measured	36
NCS DC70014 Certified	39.1
DNC-1a Measured	4
DNC-1a Certified	5.2
Analytical Duplicates	
ZR-10 Original	2420
ZR-10 Duplicate Analysis	2370
ZR-20 Original	2600
ZR-20 Duplicate Analysis	2580
ZR-30 Original	643
ZR-30 Split	750
ZR-30 Original	650
ZR-30 Duplicate Analysis	636
ZR-40 Original	283
ZR-40 Duplicate Analysis	309

ZR-50 Original	722
ZR-50 Split	866
ZR-50 Original	728
ZR-50 Duplicate Analysis	717
ZR-60 Original	5450
ZR-60 Split	5860
ZR-60 Original	5400
ZR-60 Duplicate Analysis	5490
ZR-70 Original	4440
ZR-70 Duplicate Analysis	4530
ZR-80 Original	1170
ZR-80 Duplicate Analysis	1080
ZR-90 Original	6900
ZR-90 Split	7960
ZR-90 Original	6840
ZR-90 Duplicate Analysis	6960
ZR-100 Original	1840
ZR-100 Split	1930
ZR-100 Original	1840
ZR-100 Duplicate Analysis	1830
ZR-110 Original	259
ZR-110 Duplicate Analysis	256
ZR-120 Original	360
ZR-120 Split	395
ZR-120 Original	357
ZR-120 Duplicate Analysis	362
ZR-130 Original	5200
ZR-130 Duplicate Analysis	5100
ZR-140 Original	359
ZR-140 Duplicate Analysis	362
ZR-150 Original	435
ZR-150 Split	450
ZR-150 Original	433
ZR-150 Duplicate Analysis	438
ZR-151 Original	323
ZR-151 Split	330
ZR-160 Original	3380
ZR-160 Duplicate Analysis	3420
ZR-170 Original	31
ZR-170 Split	31
ZR-170 Original	30
ZR-170 Duplicate Analysis	32
Method Blank	
Method Blank	< 3

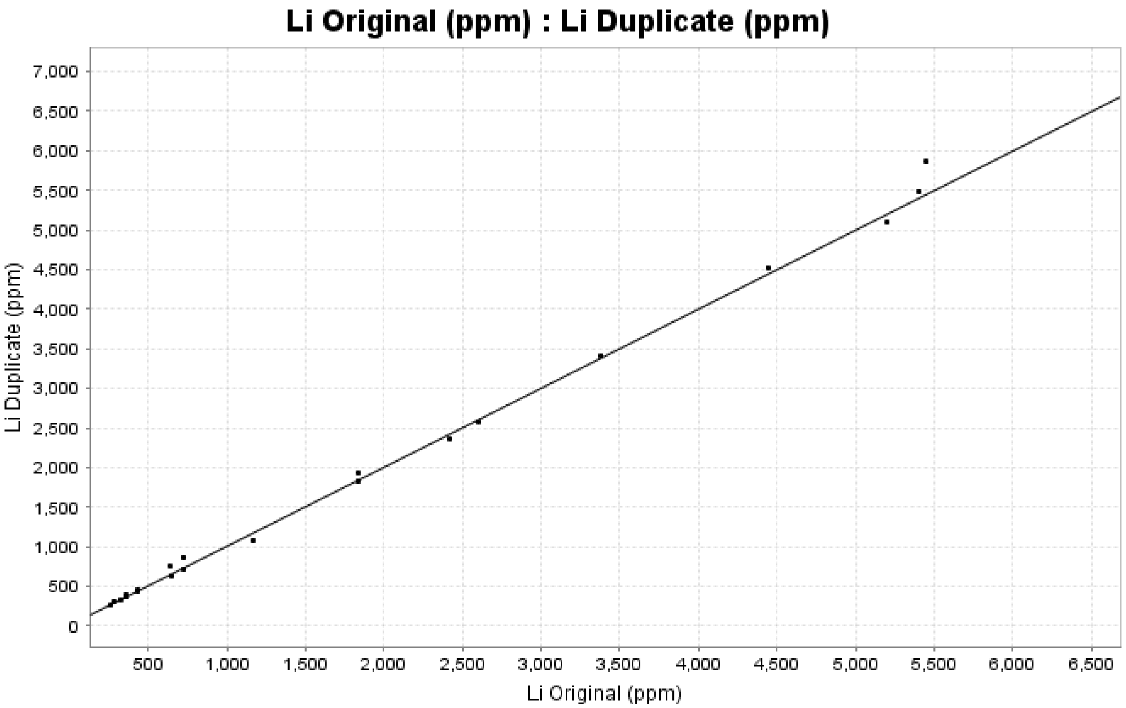


Figure 11.4.2: Graphical representation of duplicate lithium analyses, Dyke 1 trench samples.

SOIL SAMPLES

Soil samples were analyzed in the Vancouver laboratories of SGS Mineral Services (Vancouver) using their proprietary Mobil Metal Ions Technology. SGS Mineral Services is a Standards Council of Canada (SCC) accredited laboratory to ISO/IEC 17025, the international standard for testing and calibration laboratories. The management system requirements contained in ISO/IEC 17025 meet the principles of, and are aligned with, the internationally recognized quality management system standard, ISO 9001:2015.

MMI technology is an innovative analytical process that uses a unique approach to the analysis of metals in soils and related materials. Target elements are extracted using weak solutions of organic and inorganic compounds rather than conventional aggressive acid or cyanide-based digests. MMI solutions contain strong ligands, which detach and hold metal ions that were loosely bound to soil particles by weak atomic forces in aqueous solution. This extraction does not dissolve the bound forms of the metal ions. Thus, the metal ions in the MMI solutions are the chemically active or 'mobile' component of the sample. Because these mobile, loosely bound complexes are in very low concentrations, measurement is by conventional ICP-MS and the latest evolution of this technology, ICP-MS Dynamic Reaction Cell™ (DRC II™). This allows us to report very low detection limits. The method targets mobile metal ions that rise vertically through overburden from buried sources of metals. This results in few false anomalies, focused, sharp anomalies, excellent repeatability, definition of metal zones and associations, detection of deeply buried mineralization, low background values (low noise) and the detection of more regional zones of metal dispersion that can be related to mineralizing processes.

Typically field duplicate samples are collected at one sample per 30 samples during a program to monitor the reproducibility of analyses. Analytical duplicates are samples that are selected from the sample batch for re-analysis. Internal reference materials and analytical blanks are included in the analytical stream to assess accuracy and to monitor potential laboratory-based contamination.

Figure 11.4.4 is a graphical representation of analytical duplicate analyses for Li, Cs and Rb from the MMI program at the Zoro Lithium Project. Analytical reproducibility is interpreted to be excellent over a wide range of concentration.

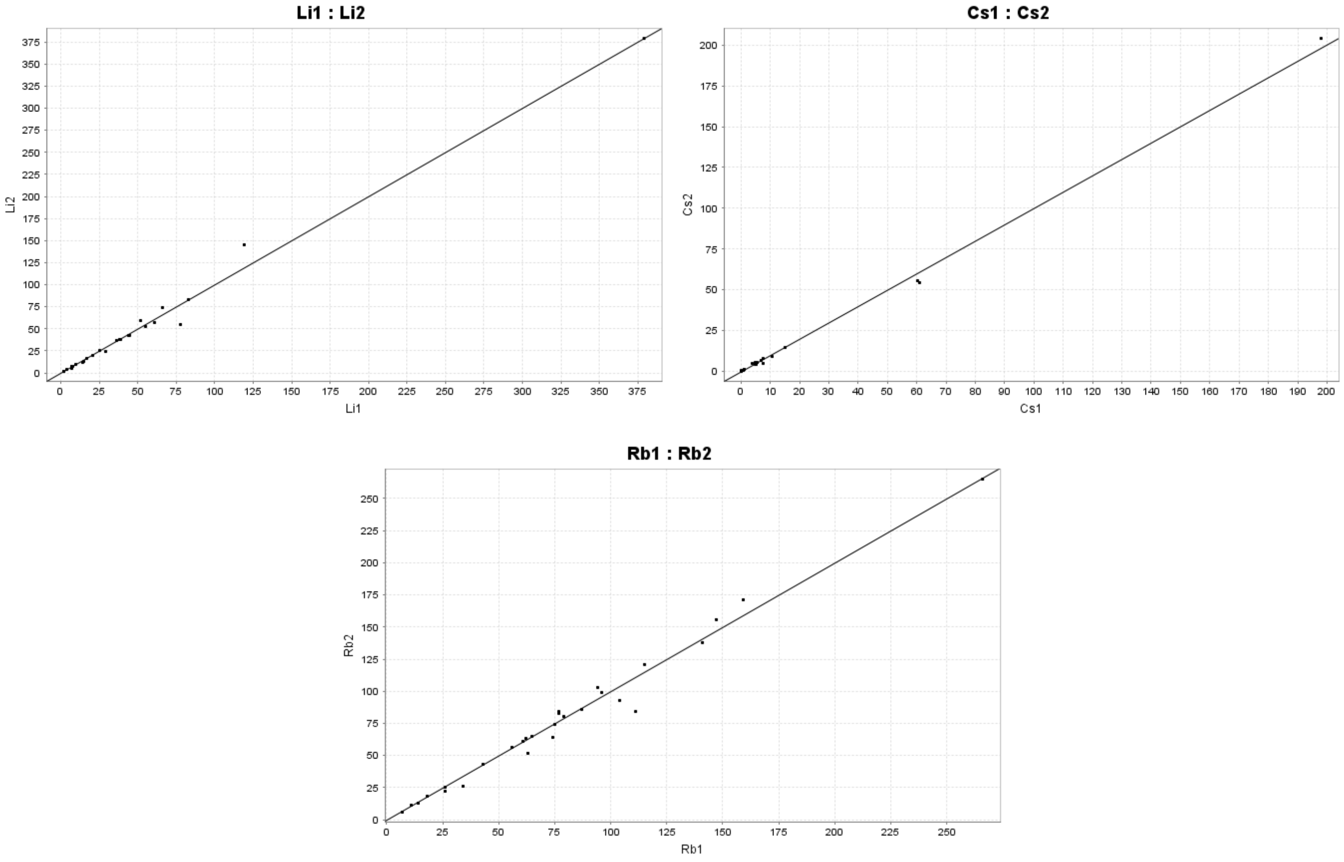


Figure 11.4.4: Plot of duplicate analyses for Li, Cs and Rb in Zoro soil samples analyzed by MMI Technology. (n=28 duplicate pairs).

MMI technology is an innovative analytical process that uses a unique approach to the analysis of metals in soils and related materials. Target elements are extracted using weak solutions of organic and inorganic compounds rather than conventional aggressive acid or cyanide-based digests. MMI solutions contain strong ligands, which detach and hold metal ions that were loosely bound to soil particles by weak atomic forces in aqueous solution. This extraction does not dissolve the bound forms of the metal ions. Thus, the metal ions in the MMI solutions are the chemically active or 'mobile' component of the sample. Because these mobile, loosely bound complexes are in very low concentrations, measurement is by conventional ICP-MS and the latest evolution of this technology, ICP-MS Dynamic Reaction Cell™ (DRC II™).

12.0 DATA VERIFICATION

12.1 MARK FEDIKOW VERIFICATION

12.1.1 HISTORIC WORK

Data verification for diamond drill core samples have not been reported in the Manitoba Government Assessment Files or other historic sources of information. A program of re-sampling drill core from DDH 5, 13, 17 and 21 was undertaken by Green Bay Exploration to assess previous assay databases (C.A.F. 93562). Samples were shipped to Ledoux and Company of New Jersey (U.S.A) but resulting assays are not reported. Historic drill core is no longer available for sampling.

12.1.2 DIAMOND DRILLING

All drill collars on the Zoro Lithium Project were surveyed using a handheld GPS and an APS unit. The surveys conducted on the Zoro Lithium Project are adequate for ongoing exploration and an eventual resource estimate. The great majority of the holes were surveyed by a REFLEX instrument (single shots approximately every 50 m).

Active drill sites have been visited during active drilling and during helicopter-assisted drill moves. Spodumene was observed in the drill core. Drill casings were observed for all drill holes.

12.1.3 OUTCROP SAMPLING

Channel and grab sample locations for Dykes 1 through 7 have been observed and documented with a hand-held GPS by the QP. Channel samples are approximately at right angles to the strike of the pegmatite and are of variable lengths.

Channel sampling was used to assess the lithium and related element contents of the pegmatites because traditional grab samples are very difficult to obtain from smooth, hard outcrop surfaces using a hammer and chisel. The channel samples are selective by nature and are likely to approximate average grades. The purpose of such sampling is to rapidly determine whether mineralization is constant throughout the outcropping pegmatite.

12.1.4 DRILL CORE

Drill core and spodumene intersections have all been observed in core racks and cross-piled drill core stored in the locked compound of Gogal Air Services. All core boxes were labelled and properly stored outside. Sample tags, located at the end of each sample, were still present in the boxes. Marks on the bottom of the box were also found, indicating sample intervals. It was possible to validate sample numbers and confirm the presence of spodumene for each of the samples in the mineralized zones.

12.1.5 ROCK AND SOIL SAMPLE COLLECTION

The goal of verification for outcrop sampling was to verify historic values for lithium reported in assessment files in the Manitoba Mining Recorders office in Winnipeg. Mineralization-level contents were obtained for all showings. The authors are satisfied that all known pegmatite occurrences described in this report contain Li mineralization. Soil sample collection was undertaken by field staff after lengthy in-field demonstrations and field experience with the MMI technique since 1997. All MMI soil samples were collected with procedures considered to be best practices.

12.2 SCOTT ZELLIGAN VERIFICATION

12.2.1 DATA VALIDATION

Validation of the drillhole and assay database was undertaken using the original drill logs and assay certificates. 100% of the database was reviewed to check for discrepancies in the collar locations, downhole survey data, and assay values. Drill logs were obtained from Orix Geoscience Inc. (who have provided database management and exploration services) as well as the original Certificates of Analysis from ACTLABS for all results. After review, there were no errors in the assay results in the database, and <5% errors in the collar/survey database, all of which were explained by incorrectly entered data in the original logs. Therefore, the only deviation from the original data were corrections to originally incorrect values.

12.2.2 SITE VISIT

On May 26th, 2018, Scott Zelligan, P.Geo., visited the Project, accompanied by Mark Fedikow, Qualified Person. The visit included flying by helicopter from Snow Lake to visit the drilling locations, as well as visiting the core logging/cutting facilities and the core farm in Snow Lake.

7 drill collar locations (with 12 collars total) were visited and measured using a Garmin GPS Map 60Csx handheld GPS. Table 12.6.1 displays the locations measured and their location according to the drill logs as compared to the validation measurement, in NAD83 (14U) Datum. The locations correspond well within the accuracy of the device (+/- 10 m). Figure 12.6.1 displays the collars visited. Additionally, one historical trench was visited (Figure 12.6.2).

The core logging and cutting facilities were visited (Figure 12.6.3) and are appropriate facilities for conducting drill logging and cutting and are maintained in excellent condition to facilitate a high-quality sampling program.

While visiting the core farm three mineralized intervals were reviewed by the authors. These mineralized intervals were selected from three different holes, one each from three of the Far drilling campaigns (locations shown in Figure 12.6.4). Table 12.6.2 shows the intervals reviewed, and Figures 12.6.5 through 12.6.7 are photos of the core reviewed. The mineralization is visually obvious and was observed as expected from the assay results for the reviewed intervals. Note some intervals were incomplete due to metallurgical sampling

Table 12.6.1: Collar locations visited on the site visit (NAD83 14U).

Drill hole ID	Site Visit Measurement (Easting, Northing)	Drill Log Location (Easting, Northing)
FAR16-007	458347 m E 6079016 m N	458345 m E 6079015 m N
FAR17-008	458429 m E 6079034 m N	458426 m E 6079037 m N
FAR17-010	458347 m E 6079016 m N	458343 m E 6079014 m N

FAR17-013	458429 m E 6079034 m N	458426 m E 6079037 m N
FAR17-018	458454 m E 6079028 m N	458452 m E 6079028 m N
FAR18-020	458400 m E 6079051 m N	458400 m E 6079051 m N
FAR18-021	458301 m E 6079027 m N	458297 m E 6079026 m N
FAR18-022	458301 m E 6079027 m N	458297 m E 6079026 m N
FAR18-023	458332 m E 6078980 m N	458332 m E 6078979 m N
FAR18-026	458292 m E 6079026 m N	458289 m E 6079025 m N
FAR18-034	459529 m E 6080336 m N	459527 m E 6080332 m N
FAR18-035	459529 m E 6080336 m N	459527 m E 6080332 m N

Figure 12.6.1: Collar locations visited on the site visit.

Figure 12.6.2: Historical trench and mineralized sample

Figure 12.6.3: Core logging and cutting facilities in Snow Lake

Table 12.6.2: Drill core intervals reviewed on site visit.

Drill hole ID	From-To (m)	Average Li_2O
FAR16-007	188.2-213.8	0.98 %
FAR17-010	162.92-171.00 181.05-192.60 201.08-213.57	0.76% 1.14% 0.46%
FAR18-020	102.1-123.5	1.56%

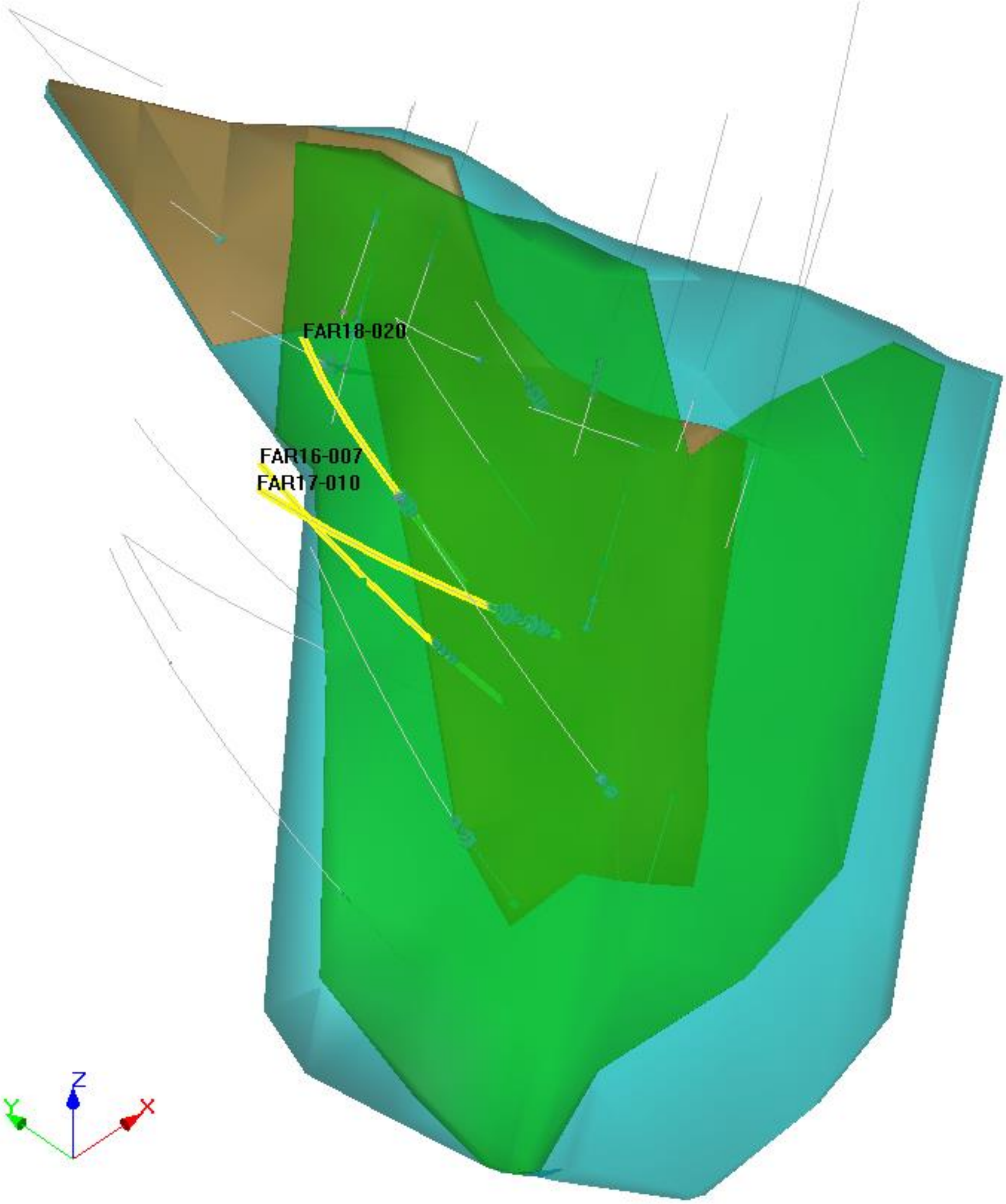


Figure 12.6.4: Location of reviewed holes in deposit (3D view looking down to the north-east).

Figure 12.6.5: Interval reviewed in FAR16-007.

Figure 12.6.6: Interval reviewed in FAR17-010.

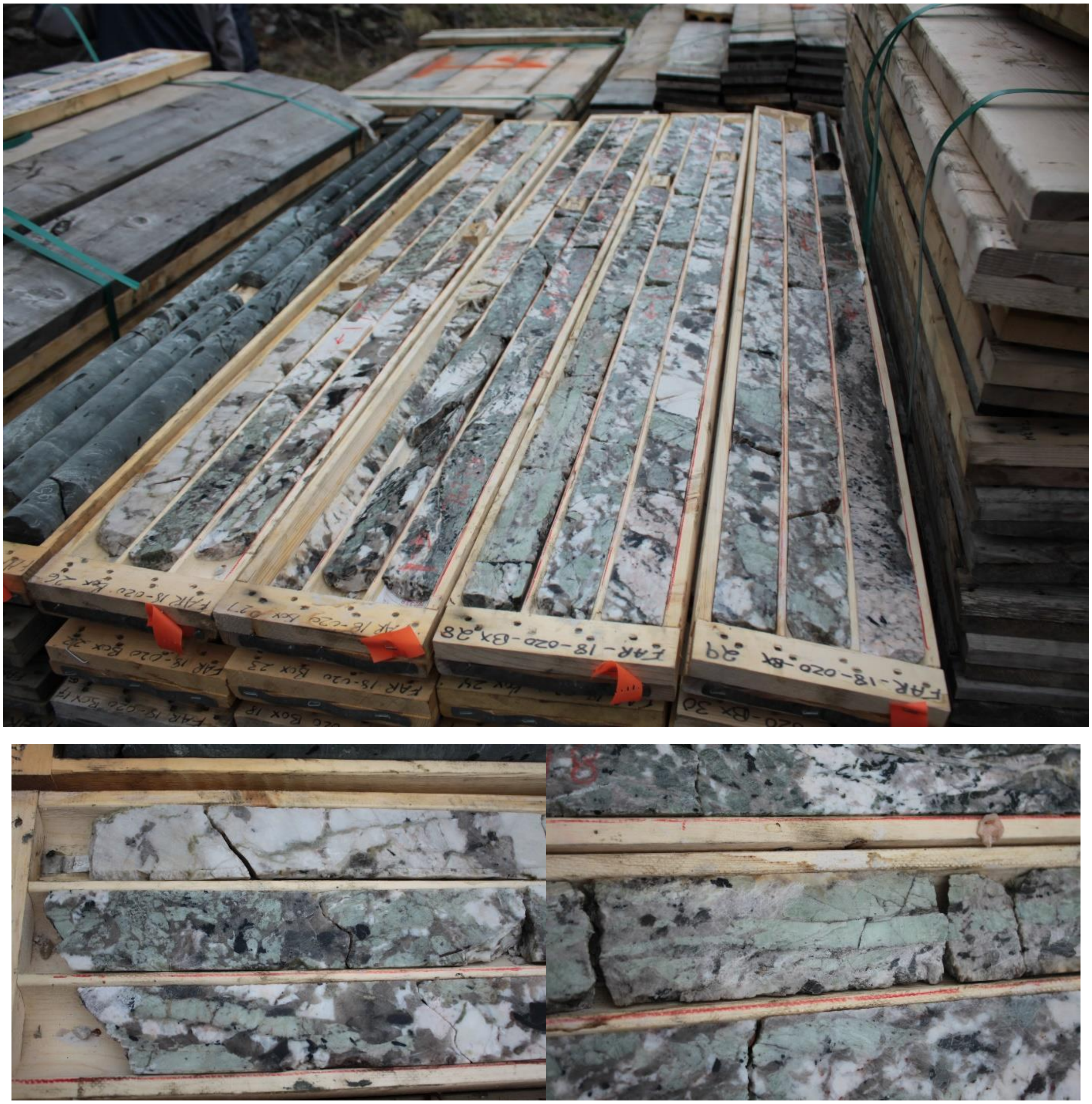


Figure 12.6.7: Interval reviewed in FAR18-020.

12.2.3 Qualified Persons Statements

The author had full access to the data and the required documentation for verification as requested, and no limitations were placed by FAR on that access.

In the opinion of the author, the data is adequate for the purposes of the inferred resource estimate calculated and presented in Section 14.

13.0 MINERAL PROCESSING AND METALLURGICAL TESTING

Not applicable.

14.0 MINERAL RESOURCE AND MINERAL RESERVE ESTIMATES

14.1 DYKE 1 INFERRED RESOURCE ESTIMATE

The following resource estimation was completed by Scott Zelligan, P.Geo, with an effective date of May 25, 2018. This resource covers only material within Dyke 1 at the Zoro Lithium Project.

Comparable properties are referenced within this section, and was obtained from: Richard & Pelletier, 2011; Dupere *et al*., 2017; Selway *et al*., 2012; and Boyko *et al*., 2018.

14.1.1 DATA

Drill hole sample data (.csv files and .dm files) and wireframes (.dm files) for this resource estimate were supplied by Orix Geoscience Inc., on behalf of Far Resources Ltd., and imported into GEOVIA Surpac™ software (version 6.3) and subsequently verified by standard internal Surpac™ processes. These .csv files contain collar, survey, lithological and assay data collated by Orix and confirmed by the author. Data includes logged and assayed diamond drill core. Lidar data (.las files) was supplied by Strider Resources Limited.

Orix supplied wireframes depicting the mineralized domains (interpreted with input from the author). These were imported and verified in Surpac™ software prior to implementation into the block model. These include:

- Low grade "pegmatite dyke" model
- Higher grade "FW" model
- Higher grade "HW" model
- Internal waste model

The author independently created an overburden/bedrock contact model, using data from the drillhole database in order to restrict estimation of the depth to the bedrock.

The main estimated commodity is Li_2O, however supplementary elements Be, Cs, Ga, Rb, and Ta were also estimated in order to best represent the value of the contained rock in Dyke 1.

The database contains 39 drillholes, 22 of which intersected Dyke 1 and were used in the estimate.

14.1.2 GEOLOGICAL INTERPRETATION

The deposit is hosted within a pegmatite dyke intruding andesite. The pegmatite is divisible and modelled from the pegmatite, however, the content of spodumene (and hence lithium) is not as clearly divisible within the dyke. For this reason, a statistical investigation was undertaken to determine the best course of action with regard to modelling the spodumene-rich volumes within the dyke. Figure 14.2.1 displays a log-histogram of the Li_2O% data, and indicates the presence of multiple populations within Dyke 1, including:

- “waste” population peaking at ~0.07 % Li_2O
- “low grade” population peaking at just above ~0.1 % Li_2O
- “high-grade” population(s) peaking imperfectly at or about 1.0-1.5 % Li_2O

A probability plot (Figure 14.2.2) of the data revealed a distinct population break at ~0.4 % Li_2O, indicating that this may be the division between the “high-grade” population(s) and the “low grade” population. This division was used as a basis to model the “high-grade” domains within Dyke 1, dubbed the “FW” and “HW” models.

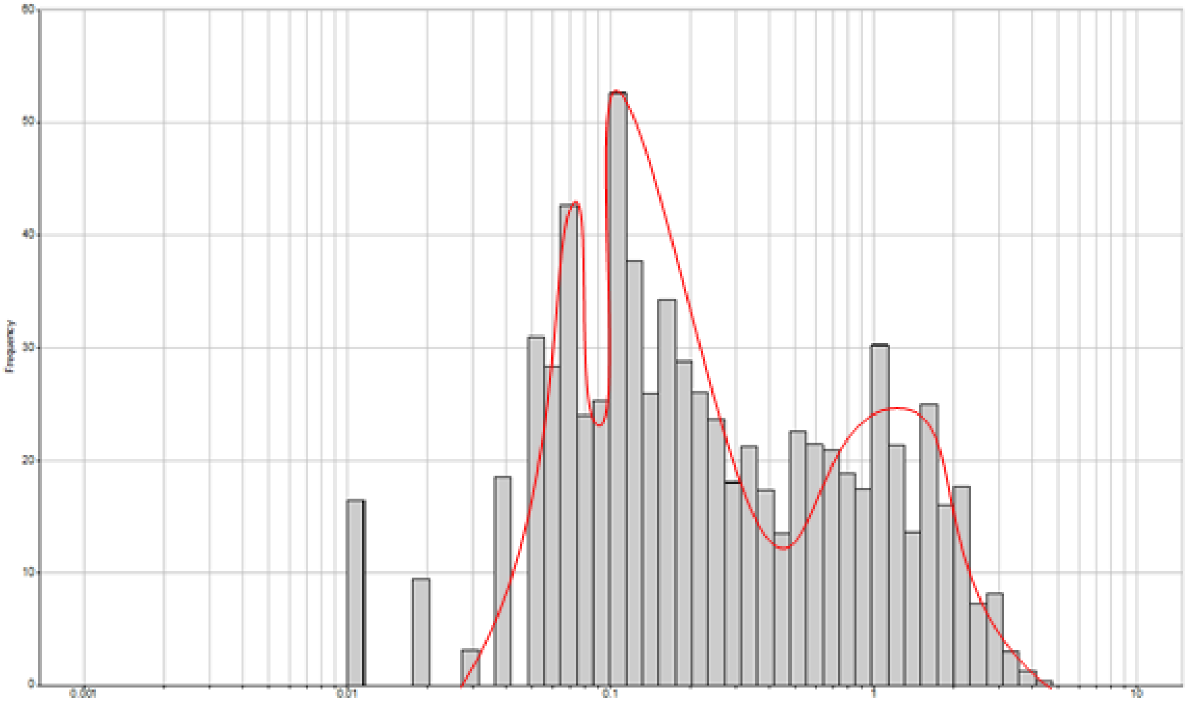
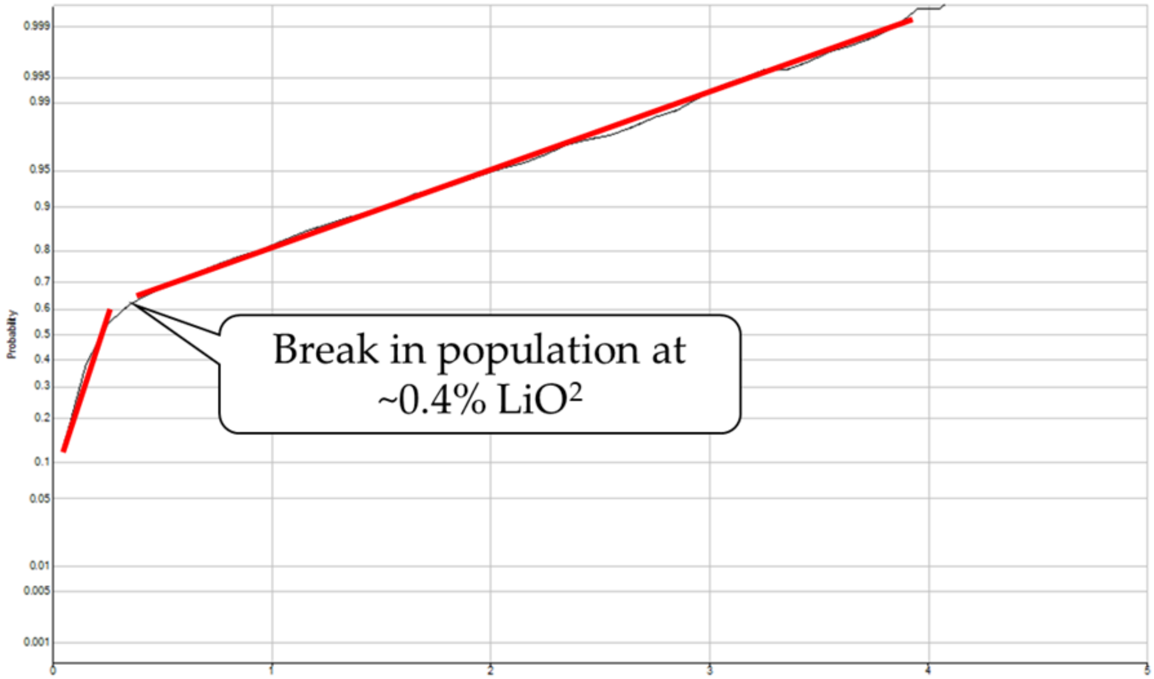


Figure 14.2.1: Log-Histogram of Li_2O% values within Dyke 1 results.



Break in population at ~0.4% LiO^2

Figure 14.2.2: Probability Plot of Li2O% values within Dyke 1 results.

14.1.3 WIREFRAMING

Wireframes based on pegmatite intersections and assay results, (Figure 14.2.3) were constructed by Orix Geoscience in Datamine (Studio EM) to represent the extents of Dyke 1. Wireframes include:

- High-grade “FW” model
- High-grade “HW” model
- Internal waste model
- Pegmatite model

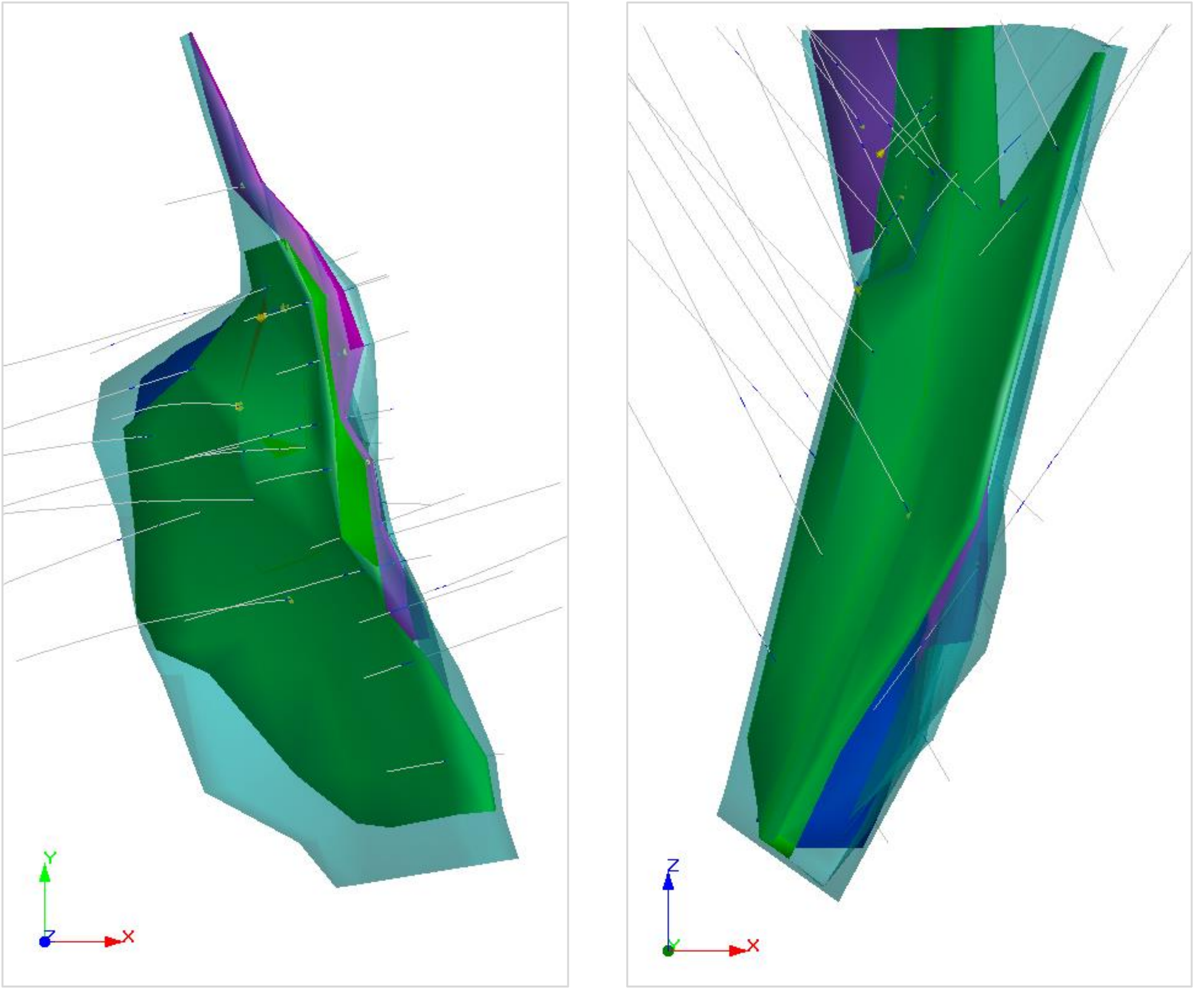


Figure 14.2.3: Domain wireframes. (Transparent light blue – Dyke 1; Green – "HW"; Pink – "FW"; Blue – internal waste)

Initial statistical investigation (Figure 14.2.4) shows that the wireframe domains appear to have captured (more or less) single log-normal grade populations. Due to the (relatively speaking) low population numbers perfect populations are statistically unlikely, however, for the purposes of this estimate these domains appear valid.

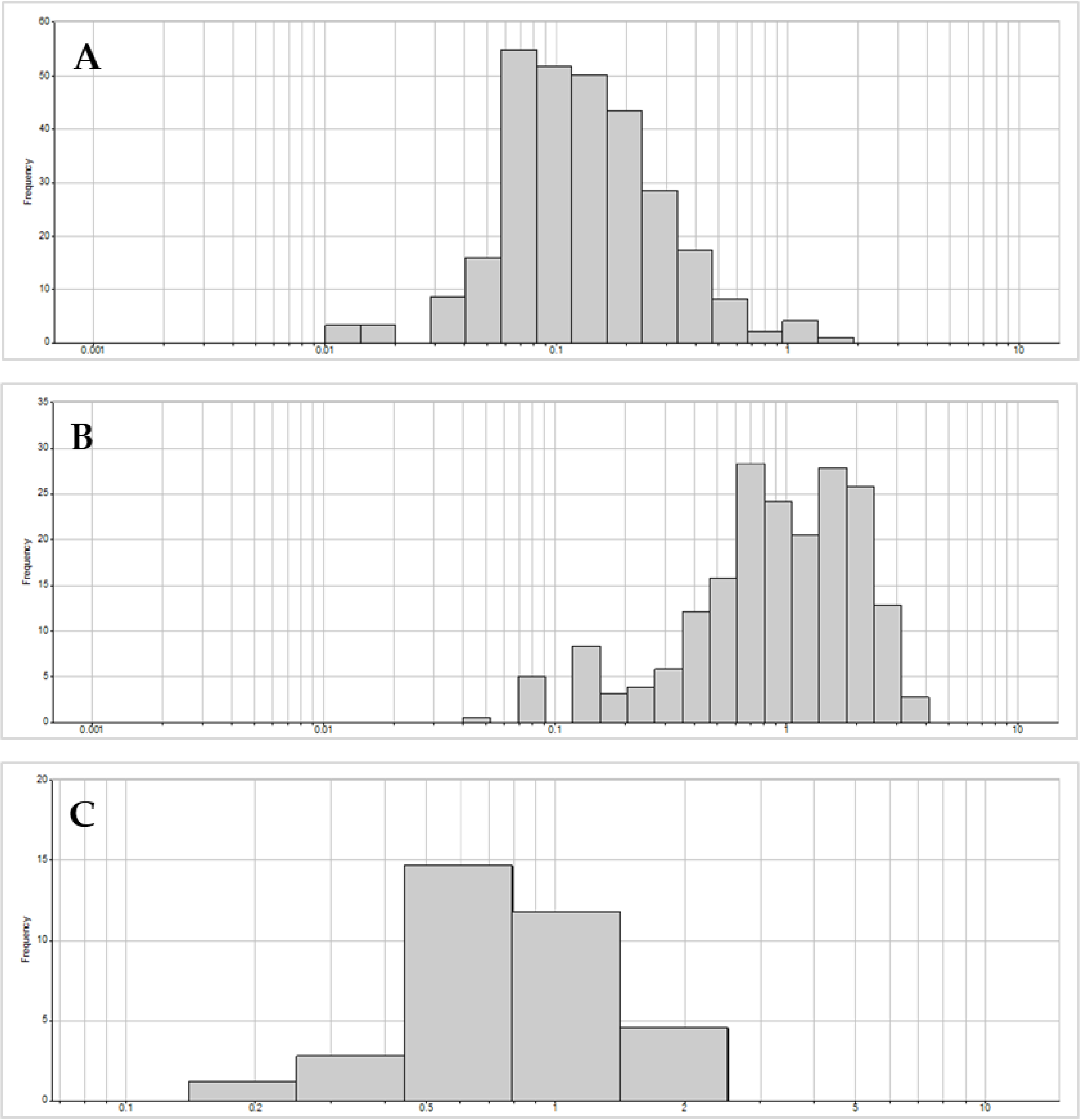


Figure 14.2.4: Log-Histogram plots of Li_2O% values for: A – Low grade, B – "HW" domain, C – "FW" domain

14.1.4 CONTACT PROFILES

Contact profiles were generated to test the validity of the wireframe models and to determine the ideal method for treating wireframe boundaries. Contact plots for Li_2O% were developed between the samples within the low-grade dyke domain and the waste, and between the "FW" and "HW" high-grade domains and the low-grade domain. These boundaries all appear to be hard/sharp. Figure 14.2.5 displays the plots.

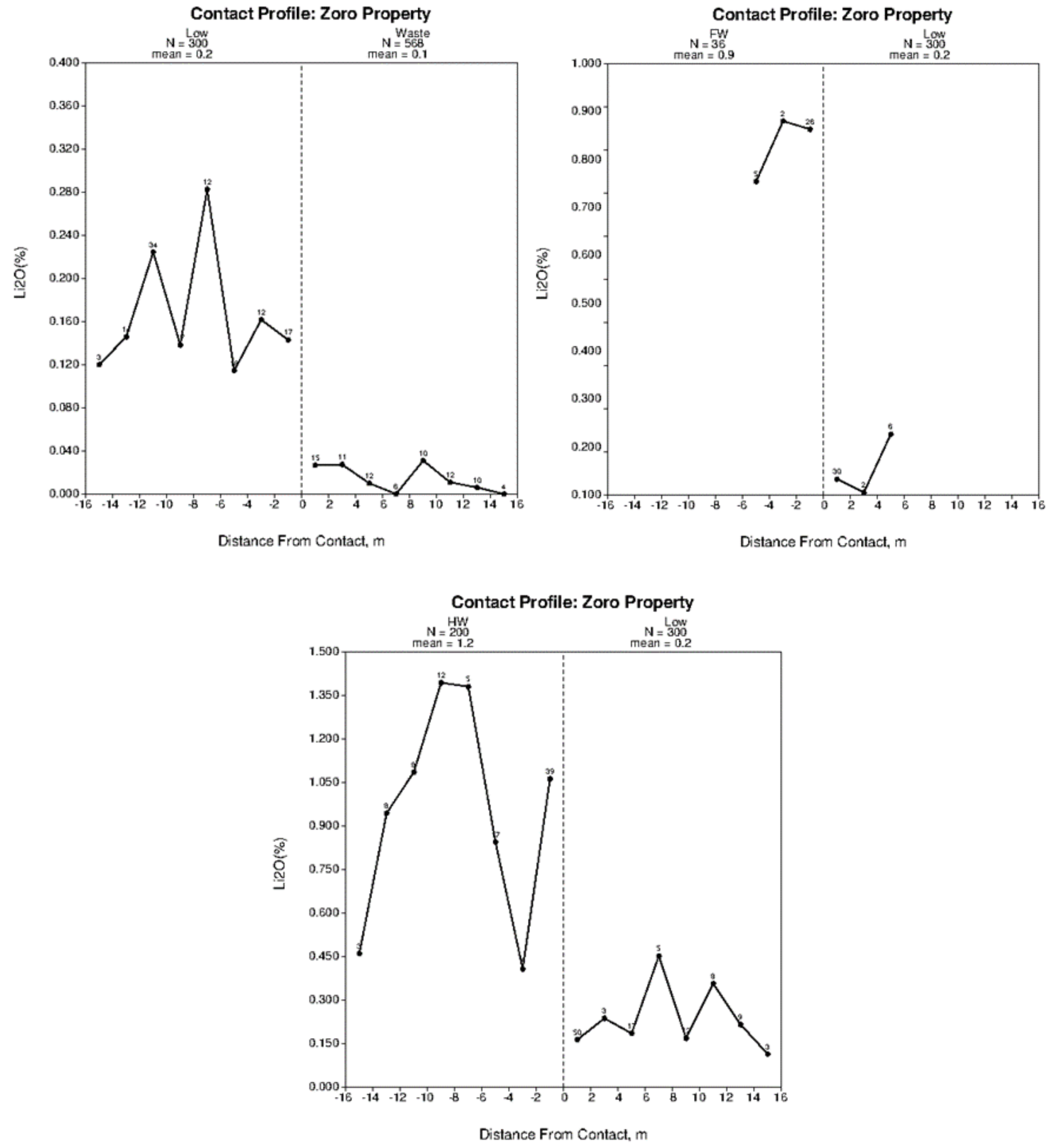


Figure 14.2.5: Contact plots for Li_2O%.

Contact profiles were also generated to test the validity of the wireframes for the other metals to be estimated. Based on statistical examination, there was potential that Cs, Rb, and Ta values in the “FW” domain represented separate populations. Based on the contact profiles (Figure 14.2.6) it was determined to treat only Cs as a separate population in the “FW” domain.

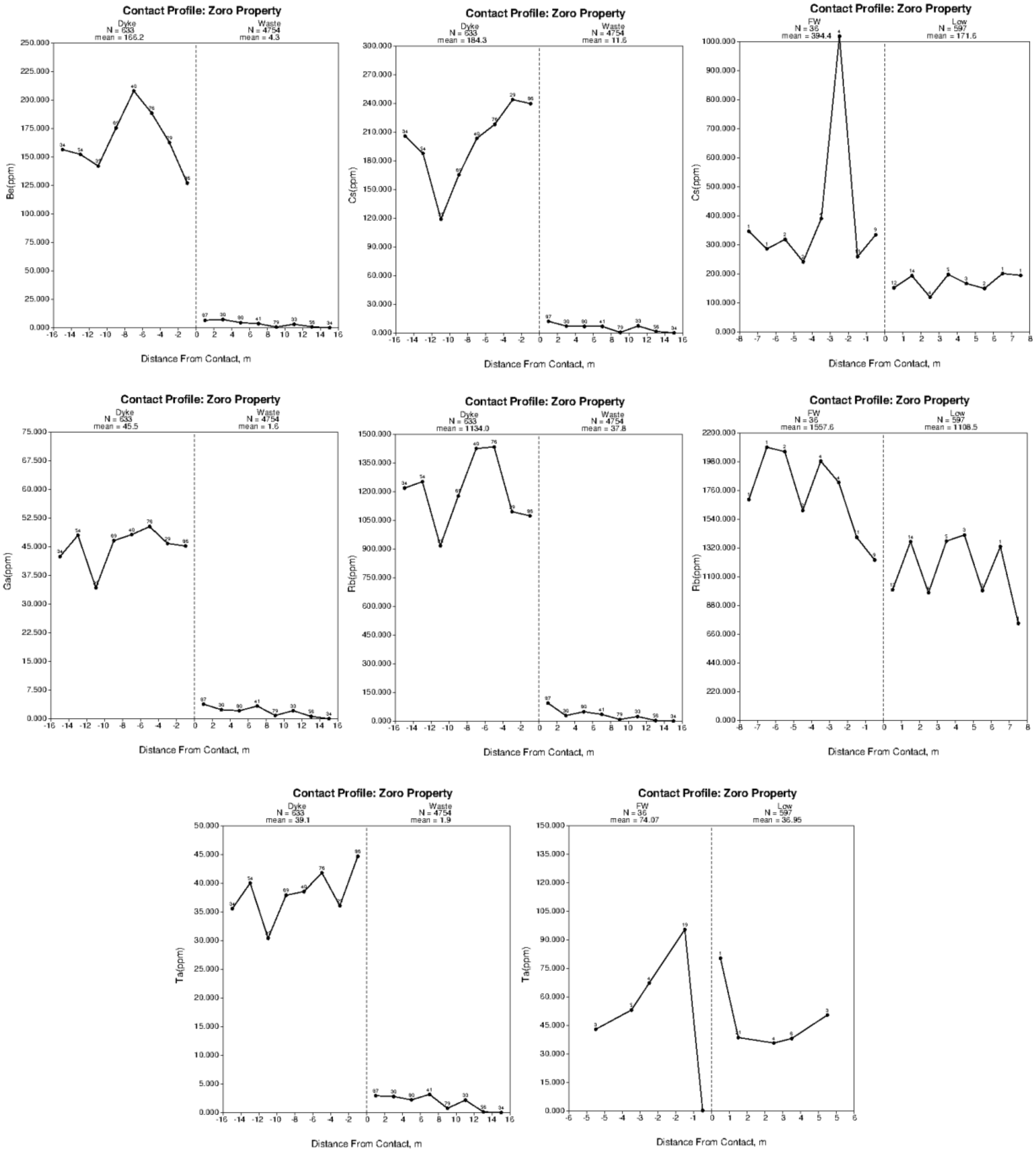


Figure 14.2.6: Contact Plots for other metals.

14.1.5 EXPLORATORY DATA ANALYSIS

RAW DATA ASSAYS AND STATISTICS

Summary statistics for the uncomposited Li_2O% assay data for each domain is shown in Table 14.2.1.

Table 14.2.1: Raw Li_2O% Sample Data by Mineralization Zone.

Li_2O%			
	Low	**FW**	**HW**
# of Samples	288	34	202
Minimum	0	0.14	0
Maximum	1.35	2.52	4.12
Mean	0.17	0.87	1.15
Variance	0.03	0.30	0.06
S.D.	0.18	0.55	0.85
Skewness	3.36	1.09	0.73
Kurtosis	14.90	1.14	0.56

Summary statistics for the uncomposited Be, Cs, Ga, Rb, and Ta assay data for the appropriate domains is shown in Table 14.2.2.

Table 14.2.2: Raw Sample Data for Be, Cs, Ga, Rb, and Ta. Units at ppm.

Dyke 1						
	Be	**Cs – Low**	**Cs – Fw**	**Ga**	**Rb**	**Ta**
# of Samples	524	490	34	524	524	524
Minimum	0	0	19.1	0	0	0
Maximum	1400	2180	3420	93.7	4100	927
Mean	185.05	197.88	394.97	51.50	1266	43.44
Variance	20745	22705	342125	291	519748	2157
S.D.	144.03	150.68	584.91	17.05	720.94	46.44
Skewness	2.18	6.03	4.60	-0.86	0.30	13.35
Kurtosis	12.60	66.50	23.08	0.92	0.15	250.93

14.1.6 COMPOSITING

Assay results from drilling were composited to 1 m, as the majority of samples were 1 m (see Figure 14.2.7) and therefore this resulted in the least amount of unnecessary sample blending. Rather than force samples to exactly 1 m, the compositing process approximated as closely to 1 m as possible within each domain and within each drillhole interval. Absent data within the raw data set was assumed to be 0 grade.

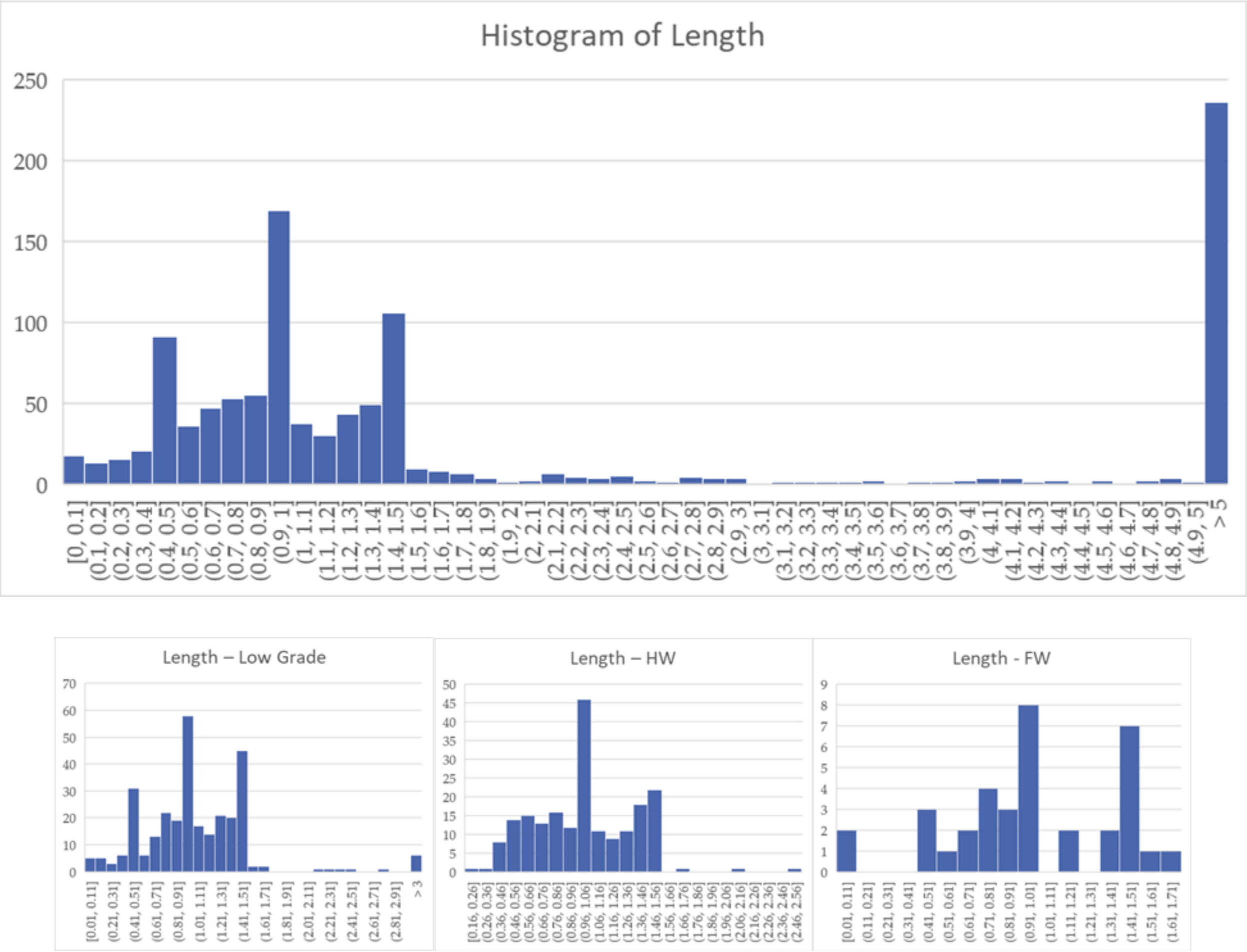


Figure 14.2.7: Histogram of Length by Domain

14.1.7 OUTLIER MANAGEMENT AND CAPPING STRATEGY

Li$_2$O%

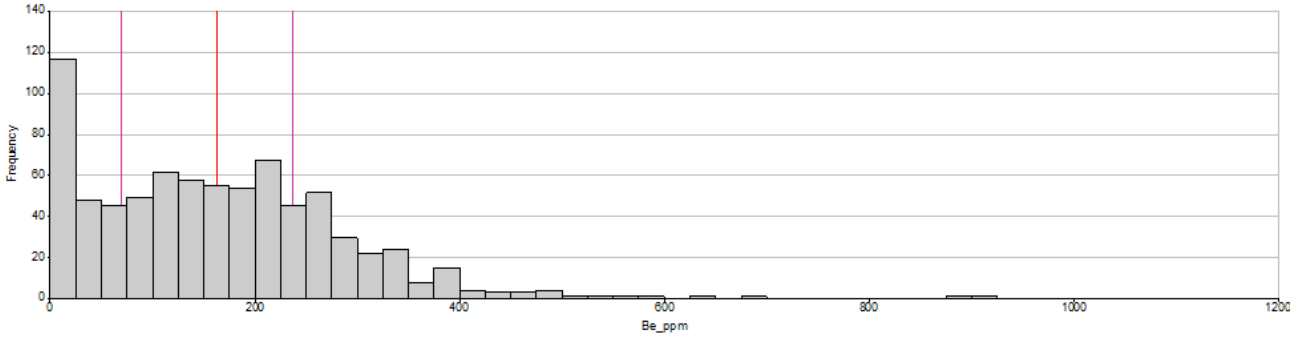
Li_2O% grades were not capped. Histograms as well as statistics (for instance, coefficient of variance is <1.2), indicate that there are no “extreme” grade values that would have an impact on the overall grade population. The maximum values, as is typical of this type of deposit, are not much greater than 2 standard deviations from the mean.

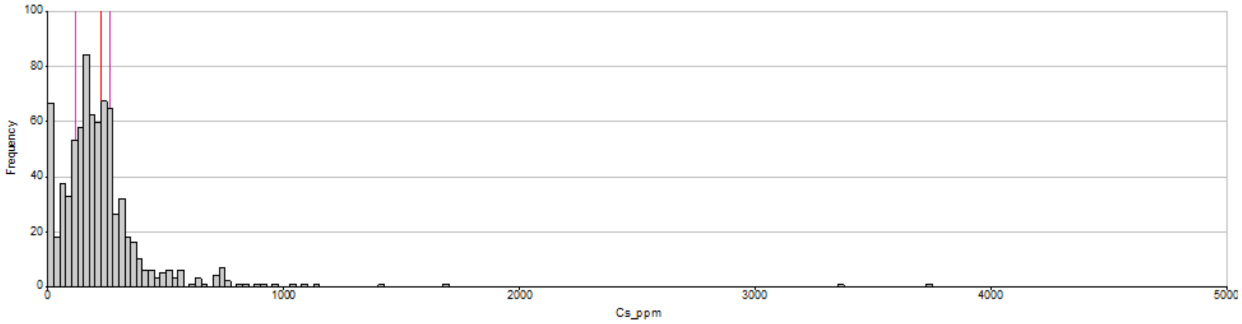
Be, Cs, Ga, Rb, and Ta

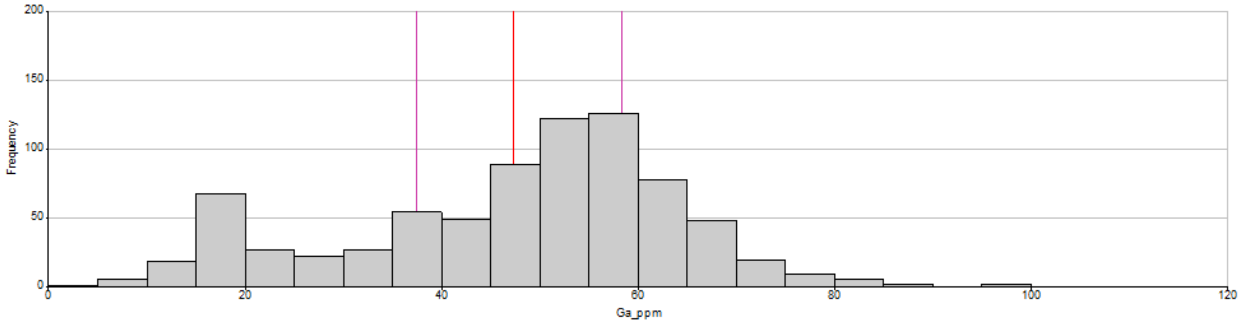
Statistics and histograms (Figure 14.2.8) for each metal were analysed to determine the best capping values.

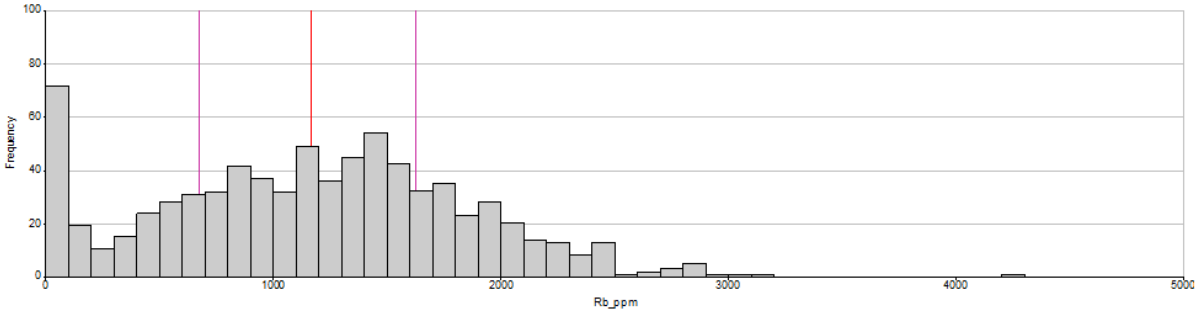
- Be was capped at 600 ppm, which resulted in 4 values being capped.
- Cs was capped at 1100 ppm, which resulted in 5 values being capped.

- Ga was not capped.
- Rb was capped at 3200 ppm, which resulted in 1 value being capped.
- Ta was capped at 200 ppm, which resulted in 1 value being capped.









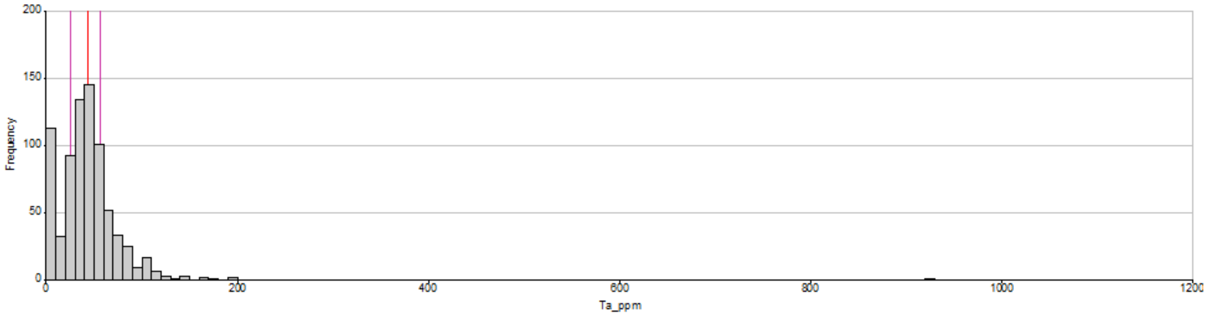


Figure 14.2.8: Histograms of Be, Cs, Ga, Rb and Ta.

14.1.8 DENSITY

A density of 2.75 t/m^3 was chosen for the tonnage estimate. This was based on values used for resource reports on comparable properties, as well as known values of pegmatite dykes.

14.1.9 INTERPOLATION PLAN

Inverse-distance-squared (ID^2) was chosen as the interpolation method. Nearest Neighbour (NN) and Inverse-distance-cubed (ID^3) were also run as a check for the results.

Variography was not performed as the sample populations were not large enough to support this method. Instead, the search ellipse anisotropy was designed to mimic the dominant orientation of Dyke 1 (as modelled).

As discussed above, contact profiles indicated a “hard” boundary between the low-grade “low” domain and the surrounding (and contained) “waste” domain, as well as the high-grade “FW” and “HW” domains and the “low” domain they are contained within, for the estimation of Li_2O%. For Be, Ga, Rb, and Ta, all material inside Dyke 1 was treated as one domain, with a “hard” boundary compared to the “waste” domain. For Cs, “low” and “HW” were treated as one domain with a “hard” boundary to the “waste” domain, and the “FW” domain was estimated as a “hard” boundary with the “low” domain.

14.1.10 BLOCK MODEL PARAMETERS

The Block Model was created with parent cells of 5 x 5 x 5 m, and a minimum sub-cell size of 1.25 x 1.25 x 1.25 m. Twenty-seven (27) interpolations were performed to populate the final grades for all metals into the block model. All domains and metals were estimated using three search ellipses, each with successively smaller search ellipses to better estimate volumes with higher sample density. Tables 14.2.3 and 14.2.4 display the search parameters and estimation parameters used in the estimation.

Table 14.2.3: Search Parameters.

	Search Ellipse 1	Search Ellipse 2	Search Ellipse 3
Minimum Samples	3	5	5
Maximum Samples	15	15	15
Maximum per Drillhole	3	3	3
Maximum Range	75	50	25
Samples Used	All samples within each domain		

Table 14.2.4: Estimation Parameters.

	Ellipse Anisotropy
Ellipsoid Plunge	0
Ellipsoid Bearing	250
Ellipsoid Dip	-75
Major: Semi-major Ratio	1.0
Major: Minor Ratio	5.0

14.1.11 RESOURCE BLOCK MODEL

CONFIGURATION

The geometrical configuration of the block model is summarized in Table 14.2.5.

Table 14.2.5: Block Model Configuration.

Origin (NAD83 14U UTM)			Block Size (m)			Min. Block Size (m)			Number of Blocks			Extent (m)		
X	Y	Z	X	Y	Z	X	Y	Z	X	Y	Z	X	Y	Z
458390	6078880	-50	5	5	5	1.25	1.25	1.25	32	63	67	160	315	335

14.1.12 CELL ATTRIBUTES

The cell attributes of the block model are summarized in Table 14.2.6.

Table 14.2.6: Block Model Attributes

Attribute	Type	Decimals	Description
be_id2	Real	2	Estimated Be grade (Be ppm)
cs_id2	Real	2	Estimated Cs grade (Cs ppm)
ga_id2	Real	2	Estimated Ga grade (Ga ppm)
li2o_id2	Real	2	Estimated Li_2O grade (Li_2O%)
rb_id2	Real	2	Estimated Rb grade (Rb ppm)
ta_id2	Real	2	Estimated Ta grade (Ta ppm)
zone	Character	-	WASTE; LOW; FW; HW

14.1.13 RESOURCE CATEGORIZATION

Mineral resource classification is the application of Measured, Indicated and Inferred categories, in order of decreasing geological confidence, to the resource block model. These are CIM definition standards (adopted by the CIM Council on May 10, 2014) for reporting on mineral resources and reserves, which are incorporated, by reference, in NI 43-101.

As per CIM (2014):

> ***Measured Resource***
>
> *A Measured Mineral Resource is that part of a Mineral Resource for which quantity, grade or quality, densities, shape, and physical characteristics are estimated with confidence sufficient to allow the application of Modifying Factors to support detailed mine planning and final evaluation of the economic viability of the deposit.*
>
> *Geological evidence is derived from detailed and reliable exploration, sampling and testing and is sufficient to confirm geological and grade or quality continuity between points of observation.*
>
> ***Indicated Resource***
>
> *An Indicated Mineral Resource is that part of a Mineral Resource for which quantity, grade or quality, densities, shape and physical characteristics are estimated with sufficient confidence to allow the application of Modifying Factors in sufficient detail to support mine planning and evaluation of the economic viability of the deposit.*
>
> *Geological evidence is derived from adequately detailed and reliable exploration, sampling and testing and is sufficient to assume geological and grade or quality continuity between points of observation.*
>
> ***Inferred Resource***
>
> *An Inferred Mineral Resource is that part of a Mineral Resource for which quantity and grade or quality are estimated on the basis of limited geological evidence and sampling. Geological evidence is sufficient to imply but not verify geological and grade or quality continuity.*

These categories are applied in consideration of, but not limited to, drill and sample spacing, QAQC, deposit-type and mineralization continuity, surface and/or underground mineralization exposure, and/or prior mining experience. With respect to resource classification of the Zoro Lithium deposit, due to the number of samples and spacing of the drillholes, the entire resource has been classified as inferred.

14.1.14 MODEL VALIDATION

STATISTICS

As in all estimates the grade average between the estimate and the originating samples has lowered. This is common in part because sampling is inevitably clustered around high-grade areas, or due to imperfections in the drilling process, creating a bias in the input which is rectified geometrically in the estimation process.

Tables 14.2.7 through 14.2.9 display the summary statistics for comparison between the raw samples, the composites, and the interpolated blocks.

Table 14.2.7: Li_2O% Summary Statistics.

	Raw Samples			Composites			Block Model		
	Low	FW	HW	Low	FW	HW	Low	FW	HW
# of Samples	288	34	202	390	36	207	14660	1590	5060
Minimum	0	0.14	0	0	0.14	0	0.00	0	0
Maximum	1.35	2.52	4.12	1.21	2.47	3.42	0.92	1.52	2.65
Mean	0.17	0.87	1.15	0.14	0.85	1.11	0.13	0.84	0.89
Variance	0.03	0.30	0.06	0.03	0.22	0.46	0.01	0.11	0.20
S.D.	0.18	0.55	0.85	0.16	0.47	0.68	0.10	0.33	0.45
Skewness	3.36	1.09	0.73	2.99	1.18	0.69	1.31	-1.09	0.09
Kurtosis	14.90	1.14	0.56	13.38	2.65	0.05	8.18	4.24	2.94

Table 14.2.8: Be, Ga, Rb, and Ta (ppm) Summary Statistics. Units are ppm.

	Raw Samples				Composites				Block Model			
	Be	Ga	Rb	Ta	Be	Ga	Rb	Ta	Be	Ga	Rb	Ta
# of Samples	524	524	524	524	633	633	633	633	15738	15738	15738	15738
Minimum	0	0	0	0	0	0	0	0	0	0	0	0
Maximum	1400	93.7	4100	927	923	87.7	3180	927	481.05	82.82	2741	136.20
Mean	185.05	51.50	1266	43.44	166.21	45.47	1134	39.07	134.42	38.10	939.56	31.71
Variance	20745	291	519748	2157	16195	486	509099	1900	7098	408	310701	379
S.D.	144.03	17.05	720.94	46.44	127.26	22.04	713.51	43.59	84.25	20.21	557.41	19.47
Skewness	2.18	-0.86	0.30	13.35	0.95	-1.05	-0.06	13.51	0.00	-0.89	-0.36	0.07
Kurtosis	12.60	0.92	0.15	250.93	3.13	0.04	-0.67	272.60	2.47	2.43	2.16	3.58

Table 14.2.9: Cs (ppm) Summary Statistics. Units are ppm.

	Raw Samples		Composites		Block Model	
	Cs – Low	Cs – Fw	Cs – Low	Cs – Fw	Cs – Low	Cs – Fw
# of Samples	490	34	597	36	15707	1590
Minimum	0	19.1	0	19.1	0	0
Maximum	2180	3420	1138.01	3371.22	628.94	1098.27
Mean	197.88	394.97	171.64	394.38	150.71	315.56
Variance	22705	342125	15792	314951	9257	40578
S.D.	150.68	584.91	125.67	561.20	96.21	201.44
Skewness	6.03	4.60	1.93	4.67	0.19	1.31
Kurtosis	66.50	23.08	11.03	23.98	3.20	5.40

14.1.15 Population Distribution

Histograms are used to determine whether the population distribution has been accurately maintained in the estimation process. This ensures that the data has not been unnecessarily smoothed.

Since this is an inferred only estimate, due to the number of samples, reproduction of grade trends will be less accurate by necessity. Considering this, the grade trends are relatively well maintained. Figures 14.2.9 through 14.2.11 display some representative grade histogram comparisons.

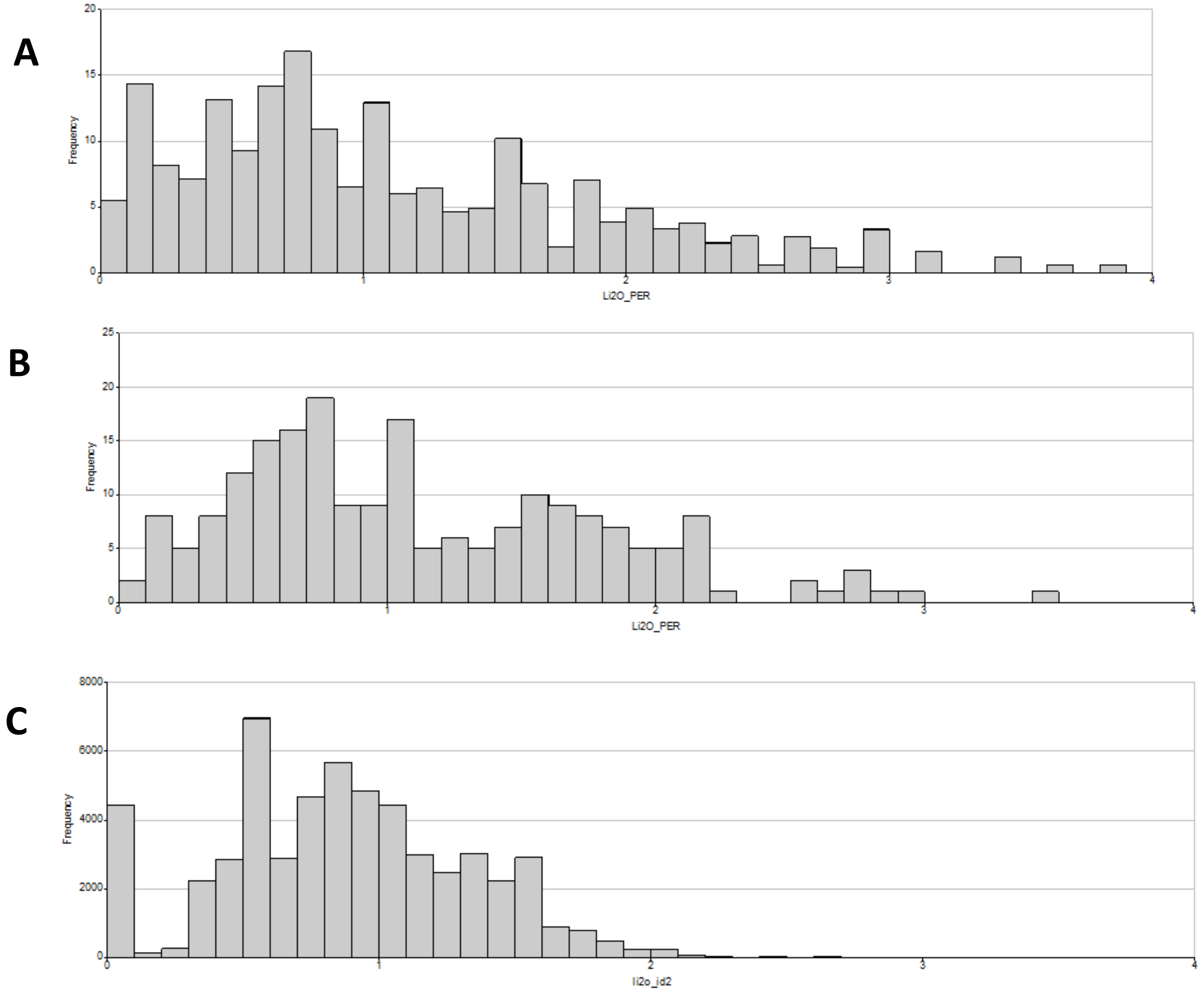


Figure 14.2.9: Li_2O% HW Histograms (A – Raw Samples; B – Composites; C – Block Model).

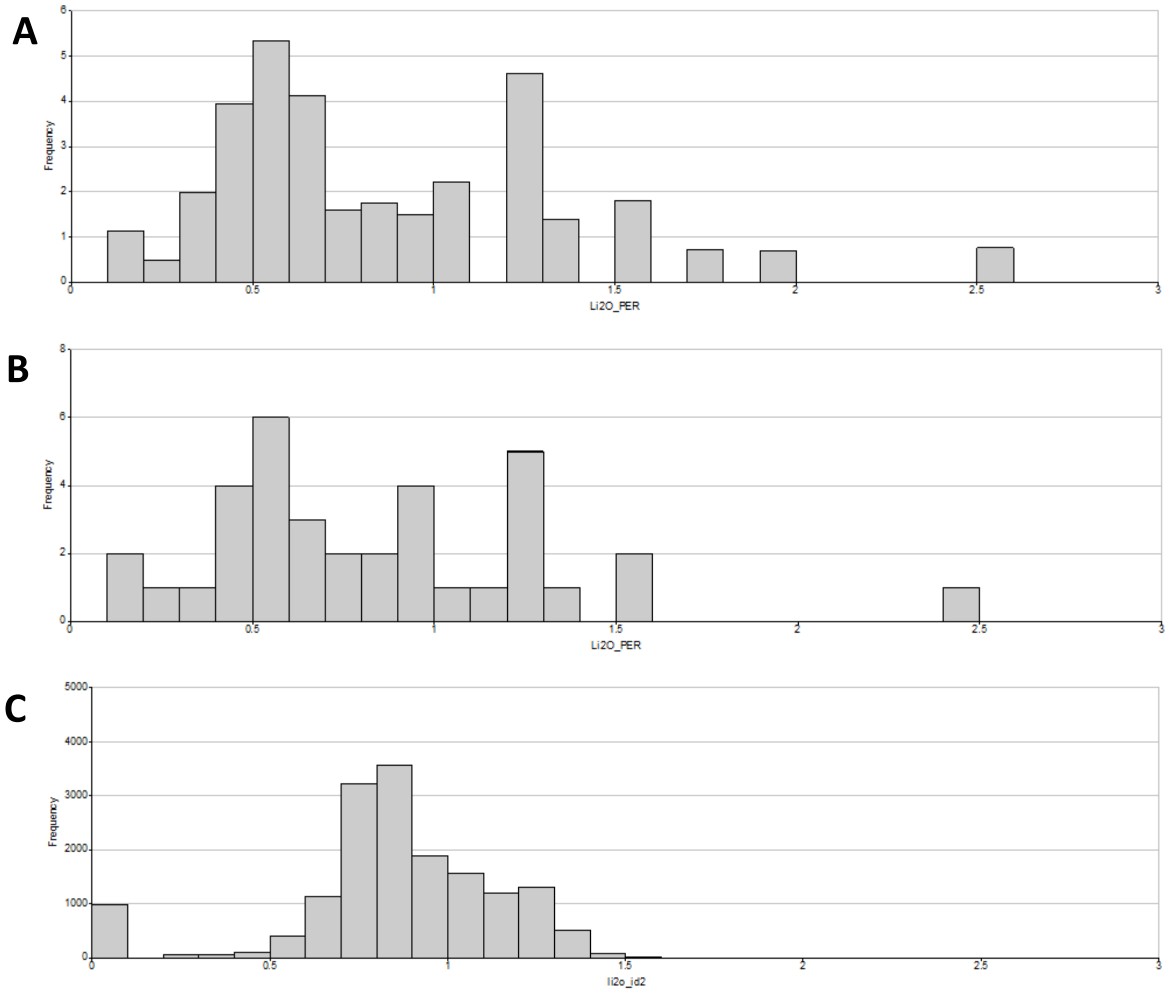


Figure 14.2.10: Li_2O% FW Histograms (A – Raw Samples; B – Composites; C – Block Model).

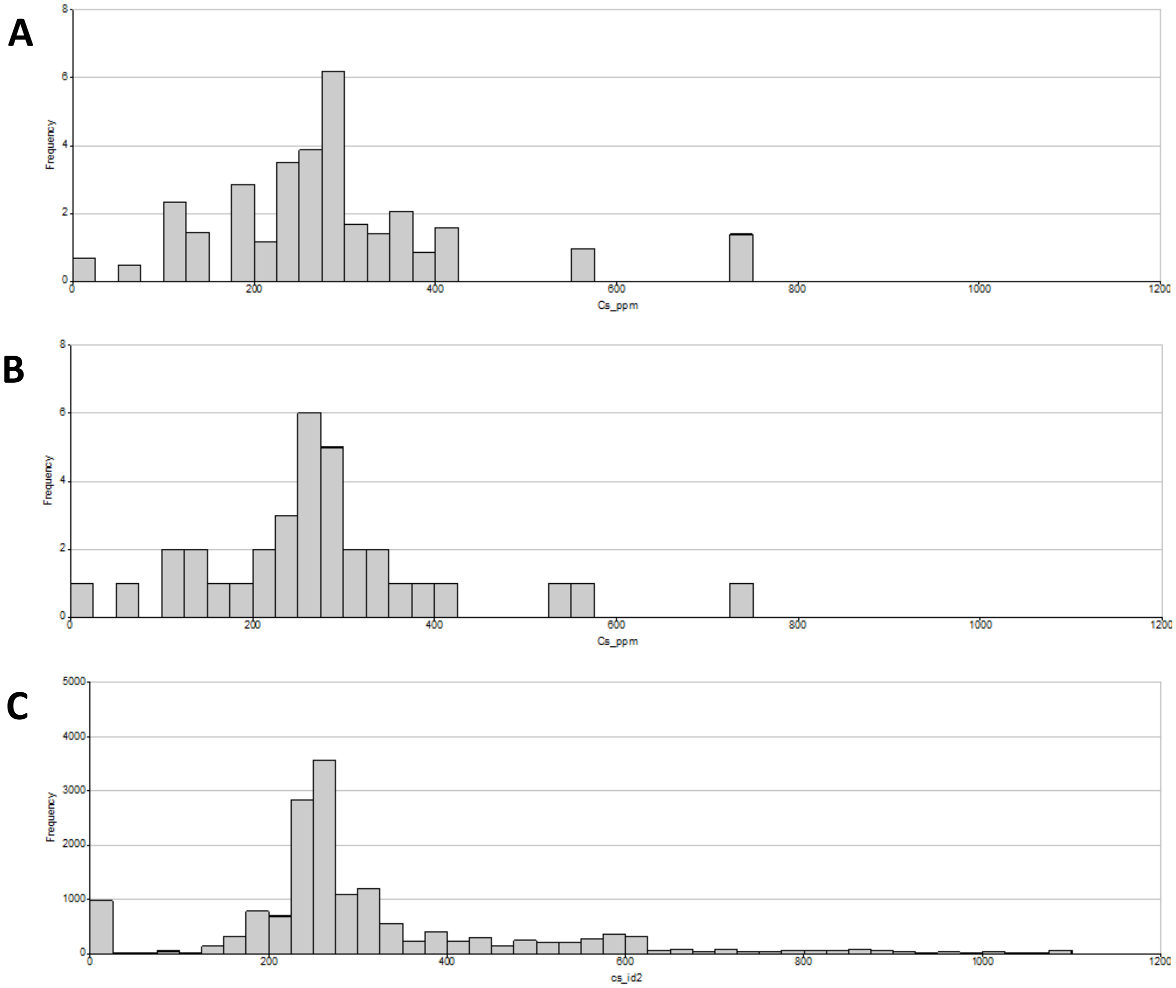


Figure 14.2.11: Cs (ppm) FW Histograms (A – Raw Samples; B – Composites; C – Block Model).

14.1.16 SECTIONS AND PLANS

Sections and Plans confirm the correlation between drill results and estimated grades. Continuity seems logical and there are no glaring mismatches between drillhole grades and block model grades. Figures 14.2.12 through 14.2.13 display representative plans and sections.

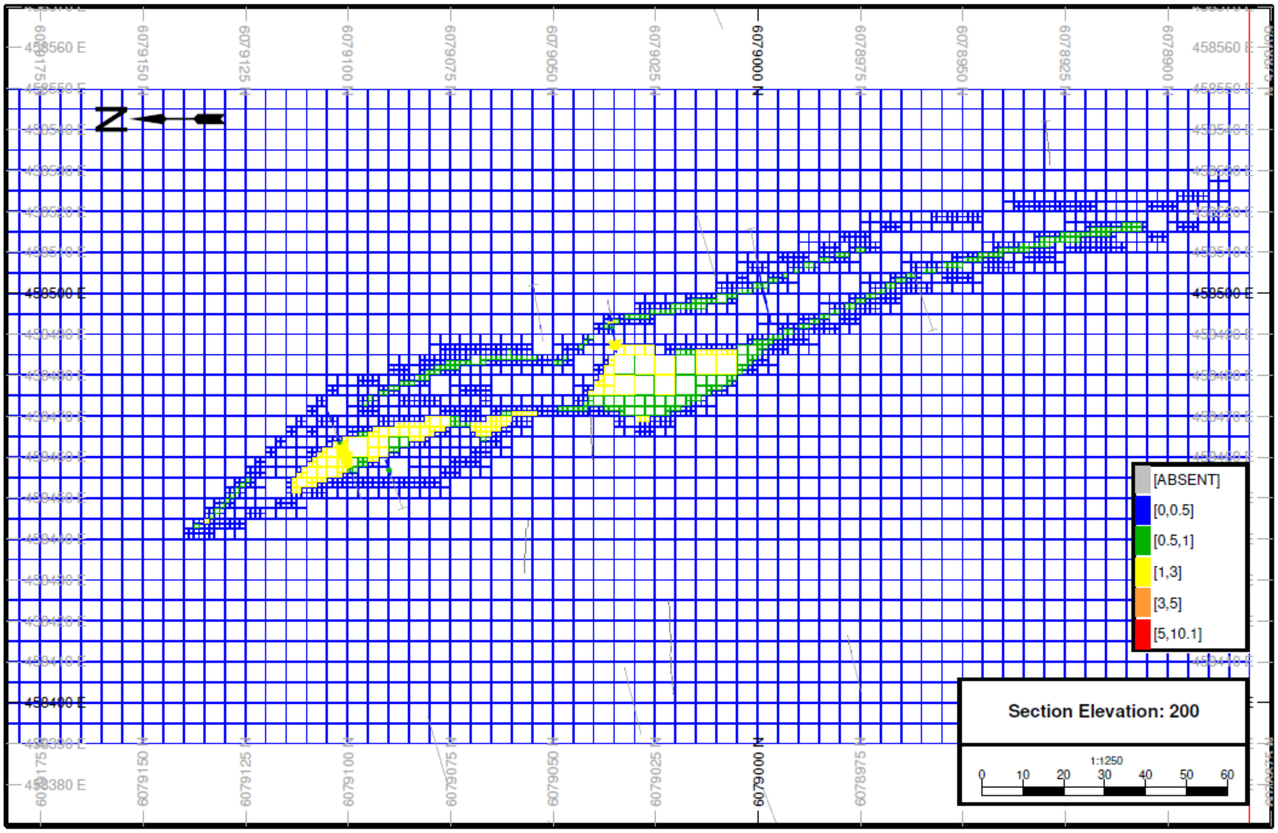


Figure 14.2.12: Plan 1, Li_2O% grades displayed, 25 m Section Width.

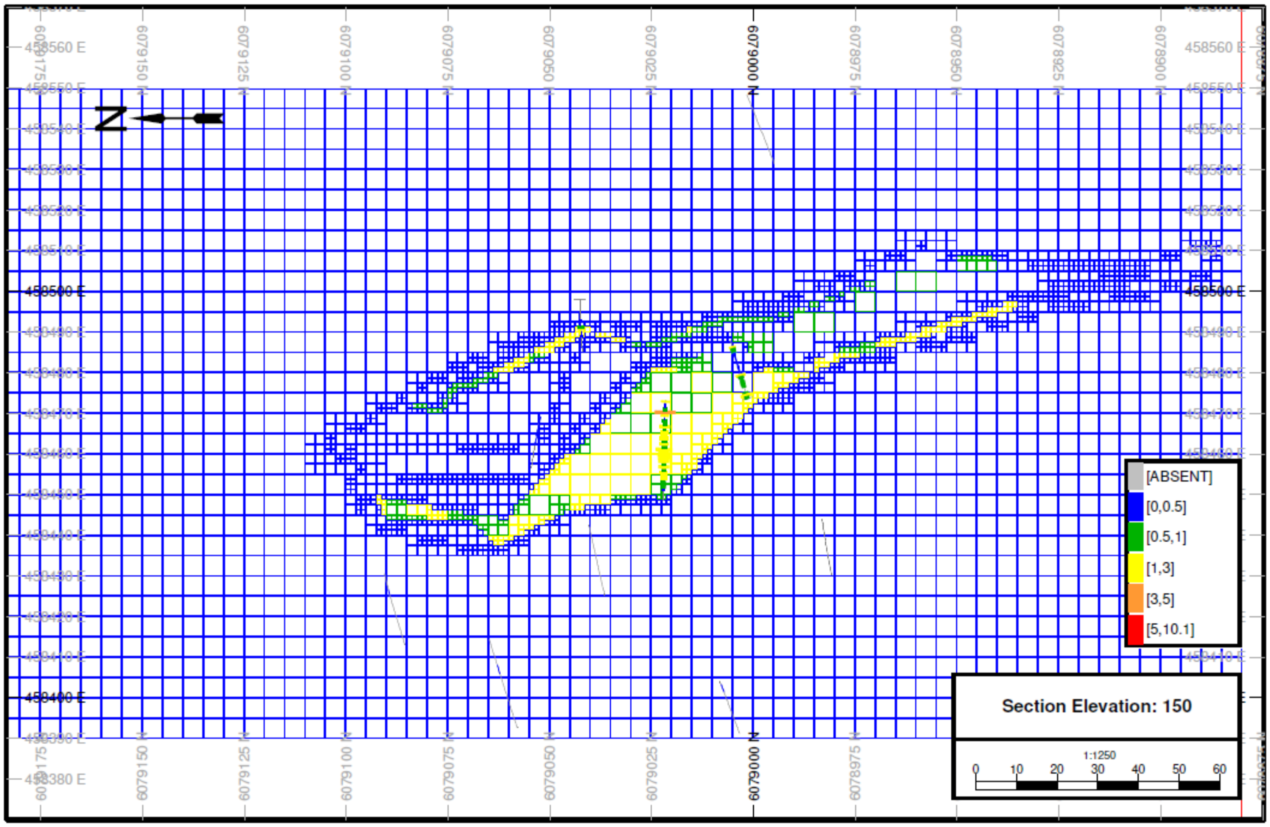


Figure 14.2.13: Plan 2, Li_2O% grades displayed, 25 m Section Width.

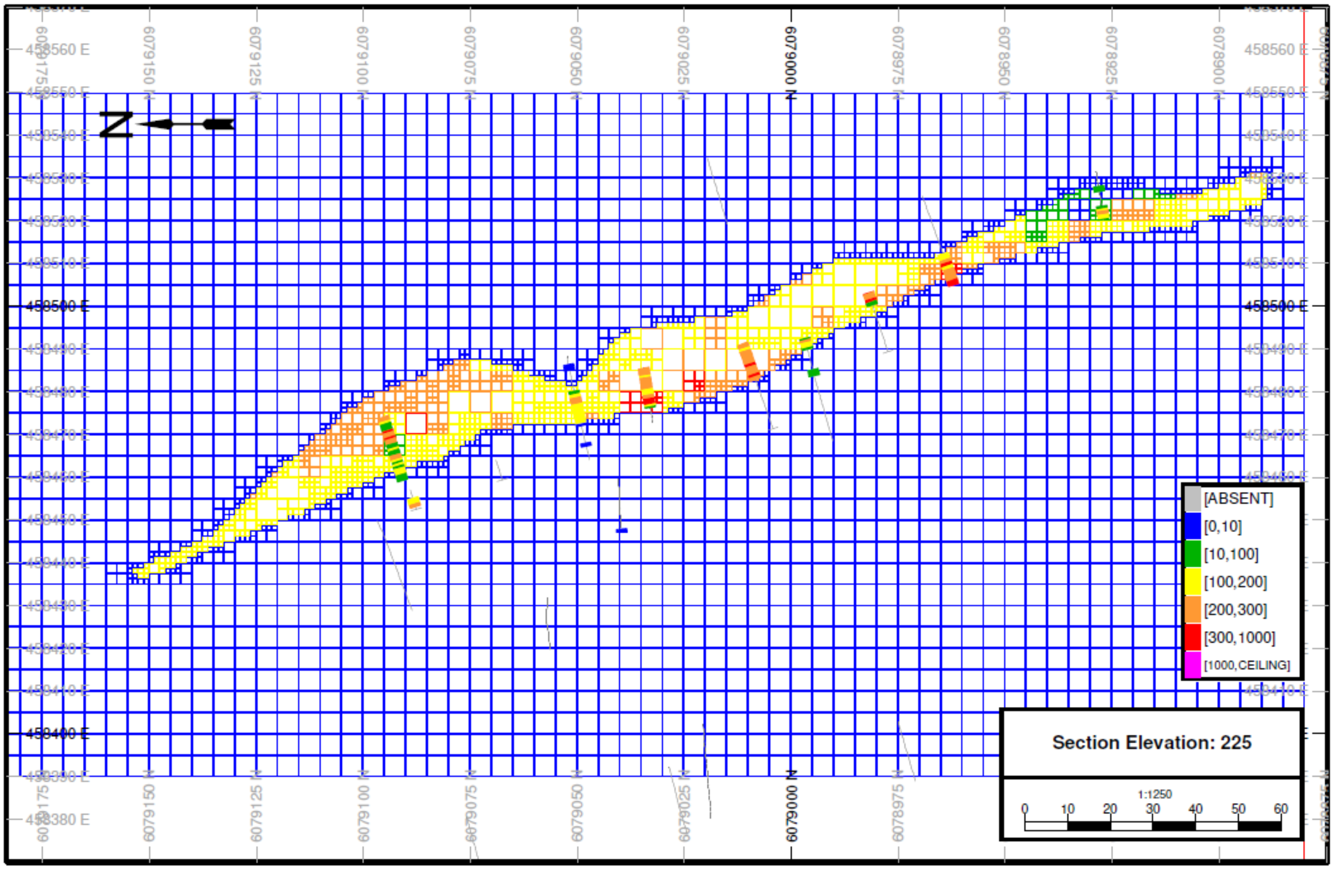


Figure 14.2.14: Plan 3, Cs ppm grades displayed, 25 m Section Width.

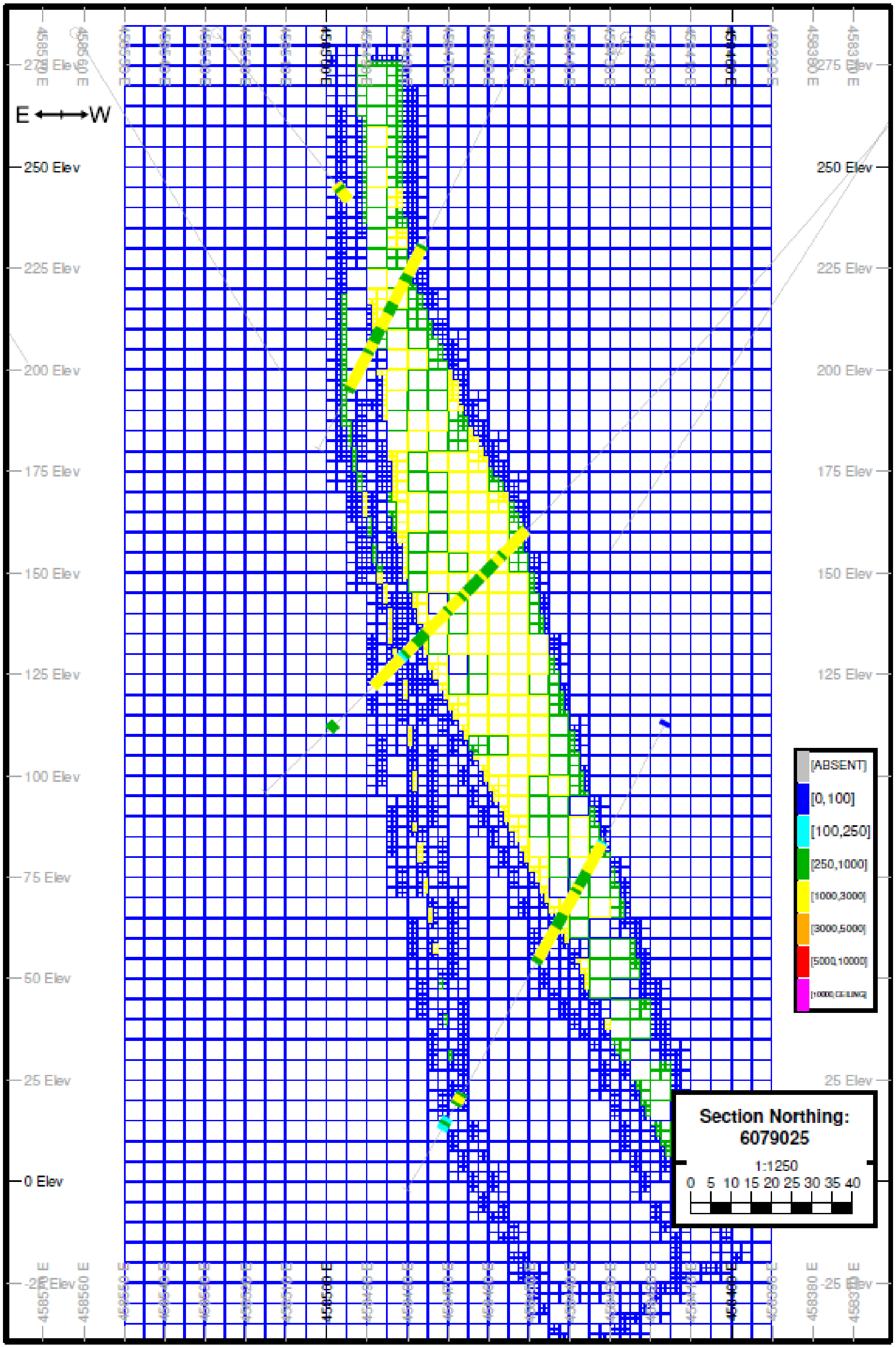


Figure 14.2.15: Section 1, Li_2O% grades displayed, 25 m Section Width.

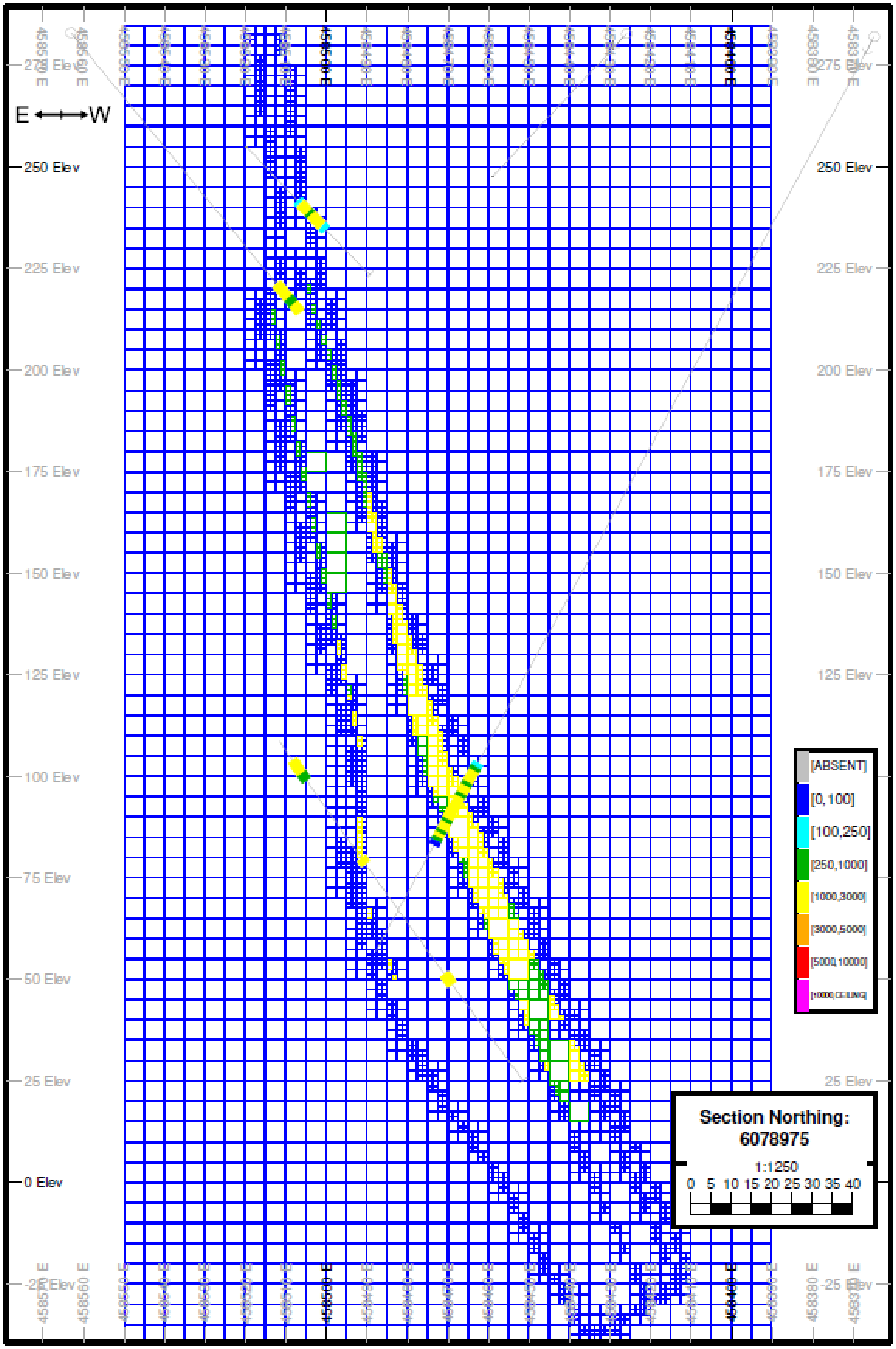


Figure 14.2.16: Section 2, Li2O% grades displayed, 25 m Section Width.

14.1.17 MINERAL RESOURCE TABULATION

For reporting purposes, the Zoro inferred mineral resource is tabulated at Li_2O (%) cut-offs. A cut-off of 0.3 % was chosen as the base case and is deemed a reasonable prospect for economic extraction based on similar reporting on other comparable properties, as well as the relevant factors discussed in Sections 4 through 8 of this report. The author is unaware of any known environmental, permitting, legal, title, taxation, socio-economic, marketing, political, or other relevant factors that may negatively affect the economic extraction of the inferred resource.

TABLES

Table 14.2.10 displays the grade-tonnage summaries by Li_2O (%) cut-off.

Table 14.2.10: Inferred Resource Grade-Tonnage Table.

Li_2O (%) Cut-off	Tonnes (t)	Li_2O (%)	Be (ppm)	Cs (ppm)	Ga (ppm)	Rb (ppm)	Ta (ppm)
0.3	**1,074,567**	**0.91**	**182**	**198**	**51**	**1212**	**43**
0.4	946,402	0.99	180	201	51	1203	43
0.5	881,815	1.03	179	203	51	1197	43
0.6	780,350	1.09	180	207	52	1196	42
0.7	721,660	1.13	179	208	52	1190	42
0.8	629,578	1.18	181	210	52	1174	42
0.9	515,652	1.26	183	211	53	1152	43
1.0	419,961	1.33	188	212	54	1135	43

***Base Case in bold text**

15.0 PROJECT INFRASTRUCTURE

16.0 MARKET STUDIES AND CONTRACTS

17.0 ENVIRONMENTAL STUDIES, PERMITTING AND SOCIAL AND COMMUNITY IMPACT

Work and drill permits for exploration are received within two weeks from the local Snow Lake office of Sustainable Development. The Zoro Lithium Project occurs within the immediate area of the historic mining town of Snow Lake and as such very little negative community and social impact is evident. Exploration and mine development has been a part of the local community for 80 years. Currently the project is an early stage exploration project. Accordingly, environmental studies have not been undertaken.

18.0 CAPITAL AND OPERATING COSTS

Not Applicable.

19.0 ECONOMIC ANALYSIS

Not Applicable.

20.0 MINERAL PROCESSING AND METALLURGICAL TESTING

No mineral processing or metallurgical testing has been undertaken as of the effective date of this report.

21.0 ADJACENT PROPERTIES

The Thompson Brothers Lithium property is located on the east shore of Crowduck Bay at the NE end of Wekusko Lake and comprises spodumene bearing pegmatite dykes. The property is 20 km east of the mining community of Snow Lake, Manitoba and approximately 5 km due west of the Zoro 1 claim of Far Resources Ltd. Highway 39 provides access from Flin Flon to Thompson and the railway going from Winnipeg to the seaport of Churchill passes approximately 30 km to the south. The main power line to Snow Lake is about 2 km south of the southern Property boundary. In the summer the Property can be accessed by boat from Wekusko Lake and a winter road can be opened on the eastern side of Wekusko Lake to provide access. Gogal Air Services based in Snow Lake offers helicopter and float plane transportation to the Property.

The lithium deposit is a spodumene-rich pegmatite dyke with a near vertical dip and a strike length of greater than 800 m. Combined Developments Limited (CDL) were the original drillers of the dyke in 1955 and 1956 and tested the pegmatite with 26 diamond drill holes for 2536.2 m commencing in 1955-56. At that time the property was known as the “Violet Property”. Prior to 1955-56, records are unavailable documenting the timing of the original discovery and when the first claims were staked. Carta Resources reported in a 1998 Business Plan that CDL had estimated a tonnage of 5,260,000 tonnes grading 1.2% Li_2O based on their drilling results. The parameters used to calculate this resource are unknown and there are no references to tonnage categories. Accordingly, this resource cannot be verified. The Thompson Brothers completed 2 diamond drillholes between 1978 and 1981, and in 1997 Carta Resources Limited financed a three-hole drill program bringing the total drilling on the property to 31 drillholes for 3,586 meters. A non-verified resource calculation made by Mr. B. Ainsworth, P.Eng. in 1998, resulted in an "undiluted drill indicated mineral resource" of 3,968,000 tonnes with a weighted average value grade of 1.29% Li_2O to the 130-metre level, and an average width of 10 meters. A further 337,000 tonnes is indicated by the deepest hole (D.H. Car-97-3), to the 380-metre level, which is over 200 meters from the nearest hole, and which cuts a horizontal width of 8 m of 1.3% Li_2O. Ainsworth suggests a total drill indicated and possible total resource of 4,305,000 tonnes of 1.3% Li_2O for the deposit. The deposit is open to depth and along strike. This 1998 resource calculation is historical and not National Instrument 43-101 compliant as it was completed prior to the implementation of these requirements. In addition, a qualified person has not done sufficient work to classify the historical estimate as a current mineral resource and the Author is not treating the historical estimate as current. Hence, the historical estimate should not be relied upon.

A metallurgical evaluation of the spodumene deposit was completed in 1997 by Dr. W. Dressler of Laurentian University. This study indicated that simple floatation would recover 92% of the spodumene and produce a concentrate grading 6.6% Li_2O or 89.2% spodumene. Further processing of the concentrate yielded a high purity lithium carbonate, Li_2CO_3, with a concentration of 98%. Carta Resources completed a business plan to develop the Strider Lithium Deposit for the production of lithium carbonate, but a decline in lithium carbonate prices stalled the project in 1998. Since then, no other work has been carried out on the Property. A NI43-101 Technical Report has been prepared by Dufresne, R. (2009) for Rodinia Minerals Inc.

A recent 43-101 report has been completed and reported (24/7/2018) by Nova Minerals Limited with an inferred resource of 6.3 Mt @ 1.38% Li2O containing 86,940 tonnes of Li2O using a 0.6% Li2O reporting cutoff. The report also identifies a remaining exploration target of 3 to 7Mt grading between 1.3 and 1.5% Li_2O in the immediate area of the resource. This resource is based entirely on a single high grade lithium bearing pegmatite dyke although a second pegmatite dyke (27/8/2018) has been reported in outcrop 300 m south of the resource.

The Thompson Brothers and Zoro properties are viewed as distinct mineralized zones without any important bearing on one another. However insufficient geologic information attributable to extensive overburden cover makes it unequivocal to determine whether the properties are geologically related and hence have a direct bearing on the relevant importance to one another.

22.0 OTHER RELEVANT DATA AND INFORMATION

The target of the exploration undertaken by Far Resources is bedrock-hosted lithium-bearing pegmatite.

Lithium is the third element in the periodic table and is the lightest of all the metals, with an atomic weight of 6.94. It tends to be concentrated in residual magmas, hence its enrichment in silicic rocks and pegmatite. Lithium occurs in some 145 different minerals, but spodumene, lepidolite, petalite, amblygonite and eucryptite are the main minerals that have been exploited commercially (Kunasz, 1983). Lithium is also produced from saline brines in desert areas, a source of important production at present.

Lithium is the most commercially important of the rare alkali metals and finds application in a wide range of industrial processes. About 10% of lithium ores and concentrates are consumed directly in the glass, ceramic, and porcelain enamel industries. Lithium is useful in these applications because it creates favorable internal nucleation conditions and imparts high mechanical strength, thermal shock resistance, as well as good chemical resistance to the product (Kunasz, 1983; Ferrell, 1983).

Most lithium is used in the metallic form or as lithium-bearing compounds and chemicals. The current use of lithium for batteries used in the electrification of transport is very significant. The most widely used compound is lithium carbonate (Li2CO3), which is added during aluminum smelting to reduce electricity consumption and fluorine emissions. Lithium carbonate is also used in the ceramics industry as a flux to lower firing temperatures and to reduce thermal expansion of enamel coatings. Lithium hydroxide (LiOH•H2O) has found an important application in lithium-based greases which maintain their viscosity over wide temperature ranges and remain stable in the presence of water. Other compounds such as lithium chloride, lithium fluoride, lithium bromide, and butyllithium have a variety of industrial uses.

Most lithium production is presently from pegmatitic minerals, mainly spodumene, though significant quantities of petalite, lepidolite, and amblygonite are shipped to Europe from several African countries. Spodumene has the highest theoretical lithium content of any mineral at nearly 8% Li2O, but most concentrates grade between 4 and 7.5% Li2O. Spodumene has high iron and low iron varieties, depending on the type of pegmatite which it is derived. High iron spodumene (about 0.6 to 0.9% Fe2O3) is generally greenish in color. Low iron spodumene (less than 0.05% Fe2O3) is white in colour and often occurs in complex, zoned pegmatites. Low iron spodumene, such as that historically produced by the Tantalum Mining Corporation of Canada (TANCO) at Bernic Lake, Manitoba, is used in glass and ceramic manufacturing. High iron spodumene generally goes into lithium chemical production, but can be processed with high-intensity magnetic separation and chlorine leach to produce low iron, ceramic grade spodumene (Harben and Bates, 1984; Buckley, 1983; Kunasz, 1982).

Lithium still finds extensive use in the glass and ceramic industry, but the new growth area for lithium is the batteries industry. Lithium, the lightest and least-dense metal in existence, provides a number of advantages over nickel and alkaline batteries. It is used to produce batteries which now have a lifespan of 15 years that can serve as the energy source for digital cameras, cell phones, clocks, watches, and toys. Lithium batteries are far lighter than their alkaline counterparts yet can last up to eight times as long. They can also withstand very harsh conditions and temperatures that would cause alkaline batteries to malfunction. Currently, nickel batteries are more affordable than lithium, but as the supply of lithium rises the demand is also expected to rise with the result being a comparably priced better-quality energy source. This has significant implications for the automobile industry and the production of lithium-powered vehicles.

The lithium market can be divided between lithium chemicals (sourced from brines or minerals), which account for some 80% of total consumption, and lithium minerals consumed directly. Reflecting growth in demand, world lithium production is estimated to have increased by some 4% per year between 2002 and 2005. The industry is characterized by a high degree of concentration of production, with two countries, Chile and Australia, together accounting for nearly two-thirds of world output and for most of the growth in production in the mid-2000s.

An interesting feature of world lithium production is the potential emergence of China as a leading supplier. The development of technology to extract lithium from high-magnesium brines has led to the start of lithium carbonate production from salt lakes in Qinghai and Tibet provinces. In late 2005, CITIC Guorun began construction of a 35,000t per year lithium carbonate plant to exploit lithium reserves in the Xitai Ginar salt lake in Qinghai province.

The long-range market picture however, remains bright because new and large uses for lithium, electric vehicle batteries and lithium alloys for aircraft, will impact demand within the near term. It is expected that demand will be strong and sustained over the long term past 2020.

23.0 INTERPRETATION AND CONCLUSIONS

The Zoro Lithium Project is located near the east shore of Wekusko Lake in west-central Manitoba, approximately 249 km southeast of Thompson and 571 km north-northeast of Winnipeg. The property consists of fourteen claims covering an area of 3603 hectares. All claims are in good standing. The pegmatite dykes were initially staked in 1953. Access is reasonable with historic bush and drill roads and trails. The property hosts a number of LCT pegmatite dykes containing high-grade lithium mineralization as spodumene.

23.1 MINERALOGY

In general, the outer zones of the pegmatite dykes contain pink aplite and coarse feldspar, muscovite, tourmaline, and beryl. Spodumene, quartz, cleavelandite, and tourmaline form core zones with interstitial coarse feldspar. Spodumene is usually coarse-grained and is sometimes altered. It is most prevalent in the central 9 m (30 ft.) of Dyke 1. In this dyke, spodumene crystals (up to 35 cm long) occur either in clusters, over widths of 6 m or more, or associated with coarse tourmaline and perthite megacrysts; some spodumene crystals show a preferred orientation of 45° to 55°. One of two parallel dykes south of the main outcrop, is 5 m wide, and contains spodumene crystals in pods (up to 33 cm across). In other dykes, coarse grained spodumene is abundant in lenticular bands and fine-grained spodumene is distributed through aplitic patches. Beryl occurs as white, anhedral to subhedral crystals less than 1 inch (2.5 cm) in diameter in three of the seven dykes. Columbite-tantalite and sparse minute grains of pyrite and chalcopyrite were found in thin sections and gold mineralization is present in quartz-rich veins and laminae in association with rare arsenopyrite, pyrrhotite and chalcopyrite. Tourmaline and muscovite are common constituents in the pegmatite dykes.

23.2 RESOURCE ESTIMATES

Based on historic drilling a grade and tonnage calculation was derived for Dyke 1 and multiple unsubstantiated resource estimates for the property were presented in the historic literature. However, these data have been acquired prior to the implementation of NI 43-101 and as such cannot be relied upon for reserve calculation, and are not being treated by FAR as current mineral resources. At best the reported grade and tonnage should be considered as a historic reserve estimate. The database resulting from historic assays was not accompanied by a quality assurance and quality control program and sampling and analytical specifics are also not reported. A program of mucking out trenches and channel sampling of these trenches with assays confirmed the presence of significant lithium-mineralized zones in Dyke 1.

Utilizing newly acquired diamond drill information and assay results from ISO-certified laboratories that have been presented in this report an inferred resource estimate of grade and tonnage for Dyke 1 has been derived. The estimates are presented in Table 23.2.1 at various cut-off levels.

Table 23.2.1: Inferred Resource Estimates of grade and tonnage for Dyke 1 at various cut-off levels.

Li_2O (%) Cut-off	Tonnes (t)	Li_2O (%)	Be (ppm)	Cs (ppm)	Ga (ppm)	Rb (ppm)	Ta (ppm)
0.3	**1,074,567**	**0.91**	**182**	**198**	**51**	**1212**	**43**
0.4	946,402	0.99	180	201	51	1203	43
0.5	881,815	1.03	179	203	51	1197	43
0.6	780,350	1.09	180	207	52	1196	42
0.7	721,660	1.13	179	208	52	1190	42
0.8	629,578	1.18	181	210	52	1174	42
0.9	515,652	1.26	183	211	53	1152	43
1.0	419,961	1.33	188	212	54	1135	43

***Base Case in bold text**

23.3 GEOCHEMICAL STUDIES

23.3.1 *MINERAL CHEMISTRY*

The results from mineral chemical studies described in this report corroborate conclusions from other studies (e.g., Halden *et al*., 1989; Linnen *et al*., 2009, 2015) based on lithogeochemistry of country rocks as a viable and relatively inexpensive tool to explore for rare-element pegmatites. This would appear to be substantiated for the metavolcanic country rocks in the Zoro study area. Linnen *et al.* (2009, 2015) found that a major drawback of using lithogeochemistry of country rocks is the occurrence of Li-Rb-Cs–bearing minerals along fractures, which complicates the interpretation of results and also suggested that indicator minerals (such as biotite) are potentially more reliable than lithogeochemistry in pegmatite exploration. Despite these potential complications exploration successes have been documented from the Dibs pegmatite (Tanco area, southeast Manitoba; Linnen *et al*., 2009). The presence of holmquistite-bearing assemblages in the amphibolitic country rock to the Dyke 1 pegmatite indicates interaction of Li-enriched fluid sourced from the Li-bearing pegmatite. Accordingly, identification of these assemblages could also be a very useful and inexpensive tool in exploration for Li-bearing pegmatite because they can occur up to 20 m away from pegmatite contacts (Cerny *et al.,* 1981). Mineral-chemistry results for muscovite and K-feldspar indicate that Dyke 1 is a moderately fractionated pegmatite (Figure GS2017-5-7a, b in Martins and Linnen, 2017). This information could be a useful tool for understanding fractionation and vectoring ongoing exploration.

23.3.2 SOIL CHEMISTRY

The discovery of Dyke 8 on the Zoro Lithium Project is attributed to the drill testing of a Mobile Metal Ions soil geochemical anomaly. The application of this technology has provided good results for a commodity element that has not been the focus of MMI applications, historically. The method, if utilized in the preferred manner provides well-defined high-contrast geochemical responses to buried sources of lithium and accordingly drill targets.

23.3.3 Uncertainties

The presence of significant overburden on the Zoro property reduces significantly the potential for exposed lithium-bearing pegmatite in the general exploration area currently under assessment by Far Resources. Wet swamp also complicates the exploration process. To overcome these issues the collection of soil samples during winter months has been undertaken resulting in the ability to effectively collect soil samples from nearly 100% of the property except where organic overburden or water preclude the practicalities of sample collection. The impact of hostile overburden conditions can reasonably be expected to slow progress of property evaluation and the search for additional lithium-bearing pegmatite dykes.

24.0 RECOMMENDATIONS

The following recommendations flow from the review of all databases derived from exploration on the Zoro Lithium Project (Table 24.1).

- Complete a detailed geologic map for the property at a scale appropriate to document lithologies exposed on the property and assess any structural characteristics relevant to an improved understanding of the emplacement and possible repetitions of Zoro LCT pegmatites.
- Diamond drilling should target the remaining pegmatite dykes exposed on the property with the aim of ascertaining the physical size and extent of the main or historic dyke in three dimensions. The deeper sections of Dyke 1 may warrant additional drilling and additional holes are required to assess the three high-grade intersections in newly discovered Dyke 8 by drill hole DDHFar18-035. To this end a program of 2000 m of core drilling is recommended.
- A mineralogical and metallurgical program for Zoro dyke 1 is strongly recommended.
- Ongoing soil geochemical surveys based on the use of Mobile Metal Ions Technology is strongly recommended given the success of drill testing an MMI anomaly by drill hole FAR18-035. This method should be applied where extensions of lithium-bearing pegmatite below overburden are sought and routinely in areas deemed to be highly prospective for lithium-bearing pegmatite but where no surface outcrop exposure is available. Drill testing of defined MMI anomalies is mandatory based on results to date.
- Option agreements for surrounding property should be considered based on the results of ongoing prospecting.

Table 24.1: Recommended ongoing exploration program and budget for the Zoro Lithium Project.

ITEM	COST
1. Diamond Drill Program (2000 m):	$375,000.00
2. Drill Geologist	$35,000.00
3. Core Technologist	$15,000.00
4. Helicopter and Drill Pad Cutting	$15,000.00
5. Assays-Drill Core (includes shipping)	$15,000.00
6. MMI Surveys including analyses @$30.00/sample for 800 soil samples, helicopter , collection (4 crew), interpretation	$110,000.00
7. Prospecting	$30,000.00
8. Assays @$50.00/sample for 100 rock samples:	$5,000.00
9. Office Support	$40,000.00
10. Project Management	$25,000.00
11. Applied Research	$30,000.00
12.Metallurgical Studies	$75,000.00
Sub-total:	$770,000.00
Contingency @ 15%:	$115,500.00
Total:	CAD$**885,500.00**

25.0 REFERENCES

Alcock, F.J., 1920: The Reed-Wekusko map-area, northern Manitoba; Geological Survey of Canada, Memoir No. 119, 47 p.

Ansdell, K.M., 1993: U-Pb zircon constraints on the timing and provenance of fluvial sedimentary rocks in the Flin Flon and Athapapuskow basins, Flin Flon Domain, Trans-Hudson Orogen, Manitoba and Saskatchewan: Geological Survey of Canada, Paper 93-2, p. 49-57.

Ansdell, K.M., and Norman, A.R., 1995: U-Pb geochronology and tectonic development of the southern flank of the Kisseynew Domain, Trans Hudson Orogen, Canada: Precambrian Research, v. 72, p. 147-167.

Ashton, K.E., and Lewry, J.F. 1994: Vergence of the "Pelican Slide" and Sturgeon-weir Shear zone: LITHOPROBE, Report 38, p. 12-17.

Ashton, K.E., Wilcox, K.H., Wheatley, K.J., Paul, D., and de Tombe, J., 1987: The boundary zone between the Flin Flon Domain, Kisseynew gneisses, and Hanson Lake Block, northern Saskatchewan: Saskatchewan Geological Survey, Saskatchewan Energy and Mines, Miscellaneous Report 87-4, p. 131-134.

Baadsgaard, H. and Cerny, P., 1993. Geochronological studies in the Winnipeg River pegmatite populations,southeastern Manitoba. Geological Association of Canada—Mineralogical Association of Canada, Annual Meeting, Edmonton 1993, Program with Abstracts, 18, p. A5.

Bailes, A.H., 1971: Preliminary compilation of the geology of the Snow Lake - Flin Flon - Sherridon area; Manitoba Department of Mines and Natural Resources, Mines Branch, Geological Paper 1/71, 27 p.

Bailes, A.H., and Syme, E.C., 1989: Geology of the Flin Flon-White Lake area: Manitoba Energy and Mines, Geological Report GR87-1, 313 p.

Bailes, A.H., and Galley, A.G., 1999: Evolution of the Paleoproterozoic Snow Lake arc assemblage and geodynamic setting for associated volcanic- hosted massive sulphide deposits, Flin Flon Belt, Manitoba, Canada: Canadian Journal of Earth Science, v. 36, p. 1789-1805.

Baker, D.R. 1998: The escape ofpegmatite dykes from granitic plutons-Constraints of new models of viscosity and propagation, Canadian Mineralogist, v. 36, p255-263.

Bannatyne, B.B., 1973: Pegmatite project; in Summary of Geological Fieldwork 1973; Manitoba Department of Mines, Resources and Environmental Management, Mines Branch, Geological Paper 2/73, p. 29, 30, 33.

Bannatyne, B.B., 1985: Industrial minerals in rare-element pegmatites of Manitoba; Manitoba Energy and Mines, Geological Services, Economic Geology Report ER84-1, p. 4, 73-75, 79.

Beus, A.A. 1960: Geochemistry of beryllium and genetic types of beryllium deposits; Academy of Science, USSR, Moscow (in Russian; translation, Freeman & Co, San Francisco, 1966), 329 p.

Bell, C.K., 1978: Geology, Wekusko Lake map area, Manitoba; Geological Survey of Canada, Memoir No. 384, p. 74-75.

Boyko, K., Canosa, J., Wilson, S., Levesque, R., and McCracken, T., 2018: NI 43-101 Technical Report Preliminary Feasibility Study For The PAK Lithium Project: Prepared by Nordmin Resource & Industrial Engineering on behalf of Frontier Lithium. Effective Date Mar 2, 2018.

Bradley, D.B., Munk, L.A., Jochens, H., Hynek, S.A. and Labay, K. 2013: A preliminary deposit model for lithium brines; U.S. Geological Survey Open File Report 2013-1006, 6 p.

Breaks, F.W. and Tindle, A.G., 1997b. Rare element exploration potential of the Separation Lake area: An emerging target for Bikita-type mineralization in the Superior province of northwest Ontario. Ontario Geological Survey, Open File Report 5966, 27 p.

Breaks, F.W., Tindle, A.G. and Smith, S.R., 1999. Geology, mineralogy, and exploration potential of the Big Mack pegmatite system: A newly discovered western extension of the Separation Rapids pegmatite group, NW Ontario. In Summary of Field Work and Other Activities 1999. Ontario Geological Survey, Open File report 6000, p. 25-1 to 25-22.

Breaks, F.W., Selway, J.B. and Tindle, A.G., 2003. Fertile peraluminous granites and related rare element mineralization in pegmatites, Superior Province, northwest and northeast Ontario: Operation Treasure Hunt. Ontario Geological Survey, Open File Report 6099, 179 p.

Breaks, F.W., Selway, J.B. and Tindle, A.G., 2005. Fertile peraluminous granites and related rare-element pegmatite mineralization, Superior Province of Ontario: Rare-Element Geochemistry and Mineral Deposits (R.L. Linnen and I.M. Samson, eds.), Geological Association of Canada, Short Course Notes 17, p. 87-125.

Brisbin, W.C., 1986, Mechanics of pegmatite intrusion: American Mineralogist, v. 71, p. 644–651.

Bruce, E.L., 1918: Amisk-Athapapuskow Lake district: Geological Survey of Canada, Memoir 105, 91 p.

Cameron, E.N., Jahns, R.H., McNai R, A.H. and Page, L.R. (1949): Internal structure of granitic pegmatites. Econ.Geol., Monogr. 2.

Cerny, P., Trueman, D.L., Ziehlke, D.V., Goad, B.E. and Paul, B.J., 1981. The Cat Lake-Winnipeg River and the Wekusko Lake pegmatite fields, Manitoba. Manitoba Department of Energy and Mines, Mineral Resources Division, Economic Geology Report ER80-1, 216 p.

Cerny, P. and Ercit, T.S., 1985. Some recent advances in the mineralogy and geochemistry of Nb and Ta in rareelement granitic pegmatites. Bull. Minéral. 108, 499-532.

Cerny, P., 1987: Petrogenesis of pegmatites in the Cat Lake - Winnipeg River and Wekusko Lake fields, Manitoba; in Centre for Precambrian Studies; University of Manitoba, Annual Report, 1987, p. 13.

Cerny, P., 1989a. Exploration strategy and methods for pegmatite deposits of tantalum. In Lanthanides, Tantalum, and Niobium. Edited by P. Moller, P. Černý and F. Saupe. Springer-Verlag, New York, p. 274-302.

Cerny, P., 1989b. Characteristics of pegmatite deposits of tantalum. In Lanthanides, Tantalum, and Niobium. Edited by P. Moller, P. Černý and F. Saupe. Springer-Verlag, New York, p. 195-239.

Cerny, P., 1989: Contrasting geochemistry of two pegmatite fields in Manitoba: products of juvenile Aphebian crust and polycyclic Archean evolution; Precambrian Research, v. 45, November 1989, p. 215-234.

Cerny, P., 1990. Distribution, affiliation and derivation of rare-element granitic pegmatites in the Canadian Shield. Geol. Rundschau 79, 183-226.

Cerny, P., 1991a. Rare-element granitic pegmatites. Part I: anatomy and internal evolution of pegmatite deposits. Geoscience Canada; volume 18, pages 49-67.

Cerny, P., 1991b. Rare-element granitic pegmatites. II. Regional to global environments and petrogenesis. Geoscience Canada 18, 68-81.

Cerny, P., 1991c. Rare-element granitic pegmatites. Part II: regional to global environments and petrogenesis. Geoscience Canada; volume 18, pages 68-81.

Cerny, P., 1992. Geochemical and petrogenetic features of mineralization in rare-element granitic pegmatites in the light of current research. Appl. Geochem. 7, 393-416.

Cerny, P., Ercit, T.S. and Vanstone, P.J., 1996. Petrology and mineralization of the Tanco rare-element pegmatite, southeastern Manitoba, Field Trip A3, Geological Association of Canada-Mineralogical Association of Canada, 63 p.

Cerny, P., Ercit, T.S. and Vanstone, P.J., 1998. Mineralogy and petrology of the Tanco rare-element pegmatite deposit, southeastern Manitoba. International Mineralogical Association, 17th General Meeting, Field Trip Guidebook B6, 74 p.

Cerny, P., 2005. The Tanco rare-element pegmatite deposit, Manitoba: regional context, internal anatomy, and global comparisons. In Rare-Element Geochemistry and Mineral Deposits (R.L. Linnen and I.M. Samson, eds.). Geol.Assoc. Can., Short Course Notes 17, 127-158.

Cerny, P., Blevin, P.L., Cuney, M. and London, D. 2005: Granite-related ore deposits. Econ. Geol., 100th Anniversary Volume, 337-370.

Cerny, P. and Ercit, T.S., 2005. The classification of granitic pegmatites revisited. In The Canadian Mineralogist, Vol. 43, pp. 2005-2026.

CIM; 2000: Exploration Best Practice Guidelines; August 20, 2000 (http://web.cim.org/standards/documents/Block465_Doc21.pdf)

CIM, 2014: CIM Definition Standards for Mineral Resources and Mineral Reserves; Prepared by the CIM Standing Committee on Reserve Definitions; Adopted by CIM Council on May 10, 2014.(http://www.cim.org/~/media/Files/PDF/Subsites/CIM_DEFINITION_STANDARDS_20142)

Connors, K.A., 1996: Unravelling the boundary between turbidites of the Kisseynew Domain and volcano-plutonic rocks of the Flin Flon domain in the eastern Trans-Hudson Orogen, Canada: Canadian Journal of Earth Sciences, v. 33, p. 811-829.

Connors, K.A., Ansdell, K.M., and Lucas, S.B, 1999: Coeval sedimentation, magmatism, and fold-thrust development in the Trans-Hudson Orogen: Propagation of deformation into an active continental arc setting, Wekusko Lake area, Manitoba: Canadian Journal of Earth Sciences, v. 36, p. 275-291.

Corrigan, D., Galley, A., and Pehrsson, S., 2007: Tectonic evolution and metallogeny of the southwestern Trans-Hudson Orogen: Geological Association of Canada, Mineral Deposits Division, Special Publication 5, p. 881-902.

David, J., Machado, N., Bailes, A., and Syme, E., 1993: U-Pb geochronology of the Proterozoic Flin Flon-Snow Lake Belt: New results: LITHOPROBE, Report 34, p. 84-87.

David, J., Bailes, A.H., and Machado, N., 1996: Evolution of the Snow Lake portion of the Paleoproterozoic Flin Flon and Kisseynew belts, Trans-Hudson Orogen, Manitoba, Canada: Precambrian Research, v. 79, p. 107-124.

Dufresne, R. 2009: Strider Lithium Property, Snow Lake Area, Herb lake Mining Division, Manitoba, Canada; NI43-101 Independent Technical Report, 65p.

Dupere, M., Perez, P., Boyd, A., Anson, J., Girard, P., Tremblay, D., Stapinsky, M., Michaud, A., Tremblay, M., and Gagnon, R., 2017: Ni 43-101 Technical Report Feasilibity Study on the Whabouchi Lithium Mine and Shawinigan Electrochemical Plant: Prepared by Met-Chem DRA Americas Inc. on behalf of Nemaska Lithium Inc. Effective Date Nov 7, 2017.

Energy, Mines and Resources Canada. 1987: Canadian mineral deposits not being mined in 1986; Mineral Bulletin MR 213, p. 211.

Fedikow, M.A.F., Roney, C.T., Schmidt, G.J. and Robbie, T.J., 1986: Mineral deposit studies in the Snow Lake area; in Manitoba Energy and Mines, Minerals Division, Report of Field Activities, 1986, p. 77, 80, 84.

Fedikow, M.A.F., Athayde, P. and Galley, A.G. 1993: Mineral deposits and occurrences in the Wekusko Lake area, 63J/13; Mineral Deposit Series Report No.14, Manitoba Energy and Mines, Geological Services Branch, 460p. includes M.D.S. Map No.14, 1:50 000.

Fedikow, M.A.F. 2009: NI 43-101 Technical Report on the Zoro 1 Property, Snow Lake Area, Manitoba (NTS 63J13SE); 67 pages, For Top Notch Marketing Ltd. 201-14881 Marine Drive, White Rock, British Columbia.

Fedikow, M.A.F. 2011: Assessment Report on Rock Geochemical Survey Results on the Zoro 1 Property (Claim P1993F), Snow Lake, Manitoba (NTS 63J13SE); 30 pages, For Top Notch Marketing Ltd. 201-14881 Marine Drive, White Rock, British Columbia.

Fedikow, M.A.F. 2012: Geology and Mineralization on the Zoro 1 Lithium Property, Snow Lake Area, Manitoba, CANADA (NTS MAP: 63J13SE), NI43-101 Report Prepared for Force Energy Ltd., 97p.

Frarey, M.J., 1948: Crowduck Bay (Descriptive Notes); Geological Survey of Canada, Paper 48-22.

Frarey, M.J., 1950: Crowduck Bay (Descriptive Notes); Geological Survey of Canada, Map 987A.

Galley, A.G., Syme, E.C., and Bailes, A.H., 2007a, Metallogeny of the Paleoproterozoic Flin Flon Belt, Manitoba and Saskatchewan, in Goodfellow, W.D., ed., Mineral Deposits of Canada: A Synthesis of Major Deposit Types, District Metallogeny, the Evolution of Geological Provinces, and Exploration: Mineral Deposits, Special Publication No.5, p.509-531.

Ginsburg, A.I., Timofeyev, I.N. and Feldman, L.G., 1979. Principles of Geology of the Granitic Pegmatites. Nedra, Moscow, USSR (in Russ.).

Harrison, J.M., 1951: Precambrian correlation and nomenclature, and problems of the Kisseynew gneisses in Manitoba: Geological Survey of Canada, Bulletin 20, 53 p.

Hoffman, P.F., 1989: Precambrian geology and tectonic history of North America, in A.W. Bailey, and Palmer, A.R., eds., The Geology of North America-An overview: Geological Society of America, Boulder, Co., p. 447-512.

Holland, H.D., Feakes, C.R., and Zbinden, E.A., 1989: The Flin Flon paleosol and the composition of the atmosphere 1.8 BYBP: American Journal of Science, v. 289, p. 362-389.

Huston, C.C. and Associates 1956: Progress report on lithium property, Herb Lake, Manitoba, 16p.

London, D., 2005. Granitic pegmatites: an assessment of current concepts and directions for the future. Lithos 80,281-303.

Lucas, S.B., White, D., Hajnal, Z., Lewry, J., Green, A., Clowes, R., Zwanzig, H., Ashton, K., Schledewitz, D., Stauffer, M., Norman, A., Williams, P.F., and Spence, G., 1994: Three-dimensional collisional structure of the Trans-Hudson orogen, Canada: Tectonophysics, v. 232, p. 161-178.

Lucas, S.B., Stern, R.A., Syme, E.C., Reilly, B.A., and Thomas, D.J., 1996: Intraoceanic tectonics and the development of continental crust; 1.92- 1.84 Ga evolution of the Flin Flon Belt, Canada: Bulletin of the Geological Association of America, v. 108, no. 5, p. 602-629.

Lucas, S.B., Syme, E.C., and Ashton, K.E., 1999: New perspectives on the Flin Flon Belt, Trans-Hudson Orogen, Manitoba and Saskatchewan: An introduction to the special issue on the NATMAP Shield Margin Project, Part 1: Canadian Journal of Earth Sciences, v. 36, p. 135-140.

Machado, N., and Zwanzig, H., 1995: U-Pb geochronology of the Kisseynew Domain in Manitoba: Provenance ages for the metasediments and timing of magmatism: LITHOPROBE, Report 38, p. 133-142.

Manitoba Mines Branch: a. Corporation Files; Green Bay Mining & Exploration Ltd.; b. Non-confidential Assessment Files; File Nos. 91564, 91614, 91624, 91650, 93562, 73164 and 73601.

Martins, T., Kremer, P., and Vanstone, P. 2013: The Tanco Mine: Geological setting, internal zonation and mineralogy of a world class rare element pegmatite deposit, Field trip Guidebook FT-C1/OF2013-8; Geological Association of Canada-Mineralogical Association of Canada Joint Annual Meeting, May 22-24, 2013, 16p.

Matile, G.L.D. and Keller, G.H. 2006: Surficial Geology of the Wekusko Lake Map Sheet (63J), Manitoba; Manitoba Science, Technology, Energy and Mines, Manitoba Geological Survey, Surficial Geology Compilation Map Series, SG-63J, Scale 1:250 000.

Maxeiner, R.O., Sibbald, T.I.I., and Watters, B.R., 1993: Geology of the Hanson Lake area (part of NTS 63L/10): Saskatchewan Geological Survey, Saskatchewan Energy and Mines, Miscellaneous Report 93-4, p. 40-49.

Maxeiner, R.O, Sibbald, T.I, Slimmon, W.L., Heaman, L.M., and Watters, B.R., 1999: Lithogeochemistry of volcano-plutonic assemblages of the southern Hanson Lake Block and southeastern Glennie Domain, Trans- Hudson Orogen: Evidence for a single island arc complex: Canadian Journal of Earth Science, v. 36, p. 209-225.

Mulligan, R., 1957a: Lithium in Canada - recent developments and geological features; Canadian Mining Journal, v. 78, No. 4, p. 125.

Mulligan, R., 1957b: Lithium deposits of Manitoba, Ontario, and Quebec 1956; Geological Survey of Canada, Paper 57-3, p. 23-24.

Mulligan, R., 1960: Beryllium occurrences in Canada; Geological Survey of Canada, Paper 60-21, p. 18.

Mulligan, R., 1965: Geology of Canadian lithium deposits; Geological Survey of Canada, Economic Geology Report No. 21, p. 80-82.

Mulligan, R., 1968: Geology of Canadian beryllium deposits; Geological Survey of Canada, Economic Geology Report No. 23, p. 70-71.

Munk, L.A., Hynek, S.A., Bradley, D.C., Boutt, D., Labay, K. and Jochens, H. 2016: Lithium brines: a global perspective; Review in Economic Geology, v. 18, p. 339–365.

Newman, W.R. 1955: Geological report on Lit claims, Johnson Lake, Herb Lake P.O., Manitoba, 9p.

Nicolas, M.P.B. 2017: Preliminary investigation of the potential for lithium in groundwater in sedimentary rocks in southwestern Manitoba; in Report of Activities 2017, Manitoba Growth, Enterprise and Trade, Manitoba Geological Survey, p. 183–190.

Partington, G.A., McNaughton, N.J. and Williams, I.S., 1995. A review of the geology, Mineralization and geochronology of the Greenbushes pegmatite, Western Australia. Economic Geology, 90, p. 616-635.

Pezzotta, F. (2000): Internal structures, paragenesis and classification of the miarolitic Li-bearing complex pegmatites of Elba Island (Italy). Memorie Soc. It. Sci. Nat. Museo.

Phillips, K.A., 1978, Minerals of Manitoba, Volume 1: Non-metallic and pegmatitic; Manitoba Department of Mines, Resources and Environmental Management, Mineral Resources Division, MRD Educational Series 78/1, p. 75.

Richard, Pierre-Luc and Pelletier, C., 2011: 43-101 Technical Report and Resource Estimate on the Pivert-Rose Property; Prepared by InnovExplo on behalf of Critical Elements Corporation. Effective Date Sept 7, 2011.

Rudenko, S.A., Romanov, V.A., Morakhovskyi, V.N., Tarasov, E.B., Galkin, G.A. and Dorokhin, V.A., 1975: Conditions of formation and controls of distribution of muscovite objects of the north Baikal muscovite province and some general problems of pegmatite consolidation. In: Muscovite Pegmatite of the U.S.S.R., Yu. M. Sokolov (ed.) Mauka, Leningrad, 174-182.

Russell, G.A., 1957: Structural studies of the Snow Lake - Herb Lake area; Manitoba Mines Branch, Publication 55-3, p. 29.

Sabina, A.P., 1972: Rocks and minerals for the collector, La Ronge-Creighton, Saskatchewan; Flin Flon-Thompson, Manitoba; Geological Survey of Canada, Paper 71-27, p. 60-61.

Sabina, A.P., 1987: Rocks and minerals for the collector, La Ronge-Creighton, Saskatchewan; Flin Flon-Thompson, Manitoba; Geological Survey of Canada, Miscellaneous Report 42, p. 50.

Schaller, R.J.R. 1955: Report to Green Bay Uranium Limited, 12p.

Selway, J.B., Breaks, F.W., and Tindle, A.G., 2005. A Review of Rare-Element (Li-Cs-Ta) Pegmatite Exploration Techniques for the Superior Province, Canada, and Large Worldwide Tantalum Deposits. Exploration and Mining Geology, Vol. 14, Nos. 1-4, pp. 1-30.

Selway, J., Baker, J., Magyarosi, Z., Dixon, A., and Sanders, R., 2012: Independent Technical Report and Updated Resource for Nama Creek Main Zone North Pegmatite: Prepared by Caracle Creek International Consulting Inc. on behalf of Rock Tech Lithium Inc. Effective Date Aug 29, 2012.

Solodov, N.A. (1962): Internal Structure and Geochemistry of Rare-Element Granitic Pegmatites. Academy Sci. USSR, Moscow, USSR (in Russian).

Stahl, C.R. 1956: Progress report on Green Bay lithium property, 17p.

Stern, R.A., and Lucas, S.B., 1994: U-Pb zircon constraints on the early tectonic history of the Flin Flon accretionary collage, Saskatchewan: Geological Survey of Canada, Current Research 1994-F, Report 8, p. 75-86.

Stern, R.A., Syme, E.C., Bailes, A.H., and Lucas, S.B., 1995a: Paleoproterozoic (1.86-1.90 Ga) arc volcanism in the Flin Flon Belt, Trans-Hudson Orogen, Canada: Contributions to Mineralogy and Petrology, v. 119, p. 117-141.

Stern, R.A., Syme, E.C., and Lucas, S.B., 1995b: Geochemistry of 1.9 Ga MORB- and OIB-like basalts from the Amisk collage, Flin Flon Belt, Canada: Evidence for an intra-oceanic origin: Geochimica et Cosmochimica Acta, v. 59, p. 3131-3154.

Stern R.A., David, J., Lucas, S.B., and Syme, E.C., 1996: Chronology of crustal growth: 1.9-1.8 Ga magmatism and accretion tectonics in the Flin Flon Belt, Trans-Hudson Orogen (Canada): Precambrian '95 International Conference on Tectonics and Metallogeny of Early/Mid Precambrian Orogenic Belts, August 28 to September 1, 1995, Montreal, Quebec, Special Issue, p. 21-27.

Stern, R.A., Machado, N., Syme, E.C., Lucas, S.B., and David, J., 1999: Chronology of crustal growth and recycling in the Paleoproterozoic Amisk collage (Flin Flon Belt), Trans-Hudson Orogen, Canada: Canadian Journal of Earth Science, p. 1807-1827.

Stilling, A., 1998. Bulk composition of the Tanco pegmatite at Bernic Lake, Manitoba, Canada. M.Sc. thesis, University of Manitoba, Winnipeg, Manitoba, 76p.

Stilling, A., Cerny, P. and Vanstone, P.J. 2006: The Tanco pegmatite at Bernic Lake, Manitoba: zonal and bulk composition and their petrogenetic significance, The Canadian Mineralogist, 44 (3).

Syme, E.C., 1988: Athapapuskow Lake Project: Manitoba Energy and Mines, Minerals Division, Report of Field Activities, 1988, p. 20-34.

Syme, E.C., 1995: 1.9 Ga arc and ocean floor assemblages and their bounding structures in the central Flin Flon Belt: LITHOPROBE, Report No. 48, p. 261-272.

Syme, E.C., and Bailes, A.H., 1993: Stratigraphic and tectonic setting of volcanogenic massive sulphide deposits, Flin Flon, Manitoba: Economic Geology, v. 88, p. 566-589.

Syme, E.C., Bailes, A.H., and Lucas, S.B., 1996: Tectonic assembly of the Paleoproterozoic Flin Flon Belt and setting of VMS deposits: Geological Association of Canada/Mineralogical Association of Canada Joint Annual Meeting, Winnipeg, Manitoba, May 27-29, 1996, Field Trip Guidebook B1, 131 p.

Syme, E.C., Lucas, S.B., Bailes, A.H., and Stern, R.A., 1999: Contrasting arc and MORB-like assemblages in the Paleoproterozoic Flin Flon Belt, Manitoba, and the role of intra-arc extension in localizing volcanic- hosted massive sulphide deposits: Canadian Journal of Earth Sciences, v. 36, p. 1767-1788.

Teertstra, D.K. and Cerny, P., 1995. First Natural Occurrences of End-Member Pollucite: A Product of Low temperature Re-Equilibration. European Journal of Mineralogy, 7, p. 1137-1148.

Thoreau, J. (1950): La pegmatite stannifère de Manono, Katanga. C.R. de Travaux, Congrès Scientifique Elisabethville 41, 1-33.

Van Lichtervelde, M., Salvi, S., Beziat, D. and Linnen, R.L., 2007: Textural features and chemical evolution in tantalum oxides: Magmatic versus hydrothermal origins for tantalum mineralization in the Tanco lower pegmatite, Manitoba, Canada; Economic Geology, v. 102, p. 257-276.

Wanless, R.K., Stevens, R.D., Lachance, G.R. and Edmonds, C.M., 1967: Age determinations and geological studies, K-Ar isotopic ages, Report 7; Geological Survey of Canada, Paper 66-17, p. 78-79.

Whalen, J.B., Syme, E.C., and Stern, R.A., 1999: Geochemical and Nd isotopic evolution of Paleoproterozoic arc-type granitoid magmatism in the Flin Flon Belt, Trans-Hudson orogen, Canada: Canadian Journal of Earth Science, v. 36, p. 227-250.

Whalen, J.B., and Hunt, P.A., 1994: Geochronological study of granitoid rocks in the Elbow Lake map area, Manitoba: A portion of the Flin Flon Domain of the Trans-Hudson Orogen: Geological Survey of Canada, Current Research 1994-F, Report 8, p. 1-10.

Wright, J.F., 1934: The Pas sheet, Manitoba and Saskatchewan (Marginal Notes); Geological survey of Canada.

Ziehlke, D.V., 1977: Petrogenesis of pegmatites in the Herb Lake area; in Centre for Precambrian Studies; University of Manitoba, Annual Report, 1977, p. 25, 30-31.

Zwanzig, H.V., 1990: Kisseynew gneiss belt in Manitoba: stratigraphy, structure, and tectonic evolution: Geological Association of Canada, Special Paper 37, p. 95-120.

Zwanzig, H.V.1999: Structure and stratigraphy of the south flank of the Kisseynew Domain in the Trans-Hudson Orogen, Manitoba: Implications for 1.845-1.77 Ga collision tectonics: Canadian Journal of Earth Science, v. 36, p. 1859-1880.

MAP REFERENCES

Map 63J/13, Herb Lake, (Topographic), Scale 1:50 000 - Mines and Technical Surveys, Canada.

Map 2566 G, Herb Lake, (Aeromagnetic), Scale 1:63 360 - Manitoba Mines Branch and Geological Survey of Canada.

Map 25019G (Airborne electromagnetic and AFMAG), Scale 1:50 000, Geological Survey of Canada (1970).

Map 36063G, Wekusko Lake, (Airborne gamma ray spectrometry), Scale 1:250 000, Province of Manitoba, Geological Survey of Canada (1979).

Maps, Snow Lake area, (Airborne geophysical surveys: airborne gamma ray spectrometry, VLF electromagnetic and Total Field magnetic), Scales 1:250 000, colour maps, 1:150 000, stacked profiles; GSC Open File 2300, Geological Survey of Canada (1990).

Maps C 21507G, C 41507G, Herb Lake, (Aeromagnetic Total Field and Vertical Gradient, colour compilation maps), Scale 1:50 000, Manitoba Energy and Mines and Geological Survey of Canada (1990).

Maps 63J/13 a, b; g, h, (Aeromagnetic Vertical Gradient/Total Field and VLF EM surveys), Scale 1:20 000; GSC Open File 2219, Geological Survey of Canada (1990).

Map 1801, Reed and Wekusko Lakes region, (Geology), Scale 1:126 720; accompanying Memoir by Alcock (1920), Geological survey of Canada.

Map 268A, The Pas sheet, Manitoba and Saskatchewan, (Geology), Scale 1:506 880; accompanying Marginal Notes by Wright (1934), Geological Survey of Canada.

Preliminary Map 48-22, Crowduck Bay, (Geology), Scale 1:31 680, by M.J. Frarey, 1946-1947; accompanying Descriptive Notes by Frarey (1948), Geological Survey of Canada.

Map 987A, Crowduck Bay, (Geology), Scale 1:63 360, by M.J. Frarey, 1946-1948; accompanying Descriptive Notes by Frarey (1950), Geological Survey of Canada.

Map 1423A, Wekusko Lake, (Geology), Scale 1:250 000, by J.E. Armstrong, 1939, M.J. Frarey, 1946-1948, G.A. Russell, 1954-1955, C.K. Bell, 1962-1963, compiled by W.L. Davison; accompanying Memoir by Bell (1978), Geological Survey of Canada.

Map, Snow Lake - Flin Flon - Sherridon area, (Geological compilation), Scale 1:253 440, by A.H. Bailes, 1970; accompanying Paper by Bailes (1971), Manitoba Mines Branch.

Map ER80-1-4, Wekusko Lake pegmatite field, (Geology), Scale 1:100 000 - Accompanying Report by Cerny *et al.* (1981), Manitoba Energy and Mines.

Claim Map Series, 63J/13 SE, Scale 1:31 680, circa 1976, 1980, 1982, Mining Recording, Manitoba Mines Branch.

26.0 CERTIFICATES OF QUALIFIED PERSONS

I, **Mark A.F. Fedikow**, B.Sc. (Honours) M.Sc. Ph.D. P.Eng. P.Geo. C.P.G., do hereby certify that:

1. I am currently a self-employed Consulting Geologist/Geochemist with a field office at:

   1207 Sunset Drive, British Columbia, Canada V8K 1E3.

   I am the senior author of the 43-101 Technical Report "Zoro Lithium Project" with an effective date of July 6, 2018. I am responsible for the contents of this Technical Report and I have reviewed and edited all sections of it.

2. I graduated with a degree in Honours Geology (B.Sc.) from the University of Windsor (Windsor, Ontario) in 1975. In addition, I earned a M.Sc. in geophysics and geochemistry from the University of Windsor and a Doctor of Philosophy (Ph.D.) in exploration geochemistry from the School of Applied Geology; University of New South Wales (Sydney, Australia) in 1982. My Ph.D. thesis concerned the development of exploration geochemical techniques to aid mineral exploration in covered terrain.
3. I am a member in good standing of the Association of Professional Engineers and Geoscientists of Manitoba as a Professional Geoscientist (P. Geo.) and a Professional Engineer (P. Eng.) being registered as a Professional Engineer since 1983. I am also registered as a professional geoscientist (P.Geo.) for the Northwest Territories and Nunavut (2017). I am a Fellow of the Association of Applied Geochemists (A.A.G.), a Member of the Prospectors and Developers Association of Canada and a Certified Professional Geologist (C.P.G.) registered with the American Institute of Professional Geologists (Westminster, Colorado, U.S.A.).
4. I have practiced my profession as a geologist for a total of forty-three years since my graduation from university; as a graduate student, as an employee of major and junior mining companies, the Manitoba Geological Survey and as an independent consultant.
5. As an exploration geologist I have been involved with all aspects of geoscientific investigation and exploration of base and precious metal mineral deposits, diamonds and industrial minerals. I have worked globally including North and South America, China, Africa, Mexico, Portugal and Israel. Specifically, I have prepared geological maps at a wide range in scale, designed geochemical exploration programs, undertook bulk sampling programs for kimberlite and diamonds, logged, sampled and analyzed drill core

1. 30 years.
2. I have read the definition of "qualified person" set out in National Instrument 43-101 ("NI 43-101") and certify that by reason of my education, affiliation with a professional association (as defined in NI 43-101) and past relevant work experience, I fulfil the requirements to be a "qualified person" for the purposes of NI 43-101.
3. I am responsible for the preparation of sections one through 11, 13 and 15 through 26 of the technical report titled "Zoro Lithium Project". Sections 12 and 14 were prepared by co-author Scott Zelligan. All sections have been prepared according to NI43-101. The date and duration of the most recent inspection of the property was May 25, 2018.
4. I am not aware of any material fact or material change with respect to the subject matter of the Technical Report that is not reflected in the Technical Report, the omission to disclose which makes the Technical Report misleading.
5. I have read National Instrument 43-101 and Form 43-101 F1. This report has been prepared in compliance with these documents to the best of my understanding.
6. I consent to the filing of the technical Report with any stock exchange and other regulatory authority and any publication by them for regulatory purposes, including electronic publication in the public company files on their web sites accessible by the public, of the Technical Report. I also consent to the use of extracts from, or a summary of, this Technical Report.

Dated this Day of July 6th 2018.

PROVINCE OF MANITOBA
M.A.F.
FEDIKOW
REGISTERED PROFESSIONAL GEOSCIENTIST

Mark Fedikow

Signed on Salt Spring Island, British Columbia~~, on this~~ this day of July 6th, 2018.

I, Scott Zelligan, B.Sc. (Honours), P.Geo., do hereby certify that:

1. I am currently a self-employed Consulting Geologist residing at:

   3357 Beechwood Drive, Coldwater, Ontario, L0K 1E0.

2. I graduated with an honours degree in Earth Sciences (B.Sc.) from Carleton University (Ottawa, Ontario) in 2008.
3. I am a member in good standing of the Association of Professional Engineers and Geoscientists of Ontario as a Professional Geoscientist (P. Geo.), License #2078.
4. I have practiced my profession as a geologist for a total of ten years since my graduation from university; as an employee of major and junior mining companies, as an employee of engineering consulting firms, and as an independent consultant, including: five months working underground in a producing gold mine; three years working in exploration for numerous commodities (including base, precious, and other minerals); and seven years modelling, estimating, and evaluating mineral properties of all types (including base, precious, and other minerals) throughout North America and occasionally globally. I have worked on numerous properties with similar mineralization styles to the Project.
5. I have read the definition of "qualified person" set out in National Instrument 43-101 ("NI 43-101") and certify that by reason of my education, affiliation with a professional association (as defined in NI 43-101) and past relevant work experience, I fulfil the requirements to be a "qualified person" for the purposes of NI 43-101.
6. I am responsible for the preparation of parts of section 12 and all of section 14 of the Technical Report titled "Zoro Lithium Project". These have been prepared according to NI43-101. The date and duration of the most recent inspection of the property was May 26, 2018.
7. I am not aware of any material fact or material change with respect to the subject matter of the Technical Report that is not reflected in the Technical

Report, the omission to disclose which makes the Technical Report misleading.

8. I have read National Instrument 43-101 and Form 43-101 F1. This report has been prepared in compliance with these documents to the best of my understanding.
9. I consent to the filing of the Technical Report with any stock exchange and other regulatory authority and any publication by them for regulatory purposes, including electronic publication in the public company files on their web sites accessible by the public, of the Technical Report. I also consent to the use of extracts from, or a summary of, this Technical Report.

Dated this Day of July 6th 2018.

Signed in Coldwater, Ontario, on this this day of July 6th 2018.

PROFESSIONAL GEOSCIENTIST
SCOTT T. ZELLIGAN
PRACTISING MEMBER
2078
ONTARIO